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# 豫西崤山韩沟岩体锆石U-Pb定年、地球化学特征及地质意义

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**提要:**韩沟二长花岗斑岩岩体位于崤山北部,其成因研究不仅可对崤山北部构造演化提供重要约束,而且可为区域找矿提供新思路。对韩沟岩体中黑云母二长花岗斑岩进行了LA-ICP-MS锆石U-Pb定年,20个测点的<sup>206</sup>Pb/<sup>238</sup>U年龄为142~147 Ma,加权平均年龄为(145.1±0.7)Ma。韩沟岩体具有高硅、富碱高钾、贫镁低钙的特征,属于富钾钙碱性花岗岩类(KCG)。韩沟岩体的稀土配分模式具有右倾、轻稀土富集及重稀土亏损的特征,(La/Yb)<sub>N</sub>范围为10.41~11.51,显示弱正Eu异常。韩沟岩体样品富集大离子亲石元素,亏损高场强元素Nb、Ta,在微量元素蛛网图上显示明显的Nb、Ta及Ti负异常。韩沟岩体具有高Sr、低Y和低Yb的特征,属于埃达克岩,Na<sub>2</sub>O和K<sub>2</sub>O含量表明其为C型埃达克岩。韩沟岩体是加厚大陆下地壳的部分熔融形成的,部分熔融源区的残余相矿物包括石榴石和金红石,且无(富钙)斜长石残留。崤山北部至少经历了~145 Ma和~130 Ma两次规模不等的岩石圈拆沉作用,韩沟岩体为~145 Ma规模较小岩石圈拆沉作用的产物。韩沟岩体的形成时代与~145 Ma区域成岩成矿期相接近,在岩石圈拆沉作用这一灾变过程中,必然导致深部流体的大规模活动及其快速释放,表明韩沟岩体周缘及深部应具有较大的Au成矿潜力,其周缘11个Au矿床(点)就是最好的找矿证明。

**关 键 词:**韩沟岩体;LA-ICP-MS锆石定年;埃达克岩;岩石圈拆沉;成矿潜力;崤山

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## Zircon U-Pb dating and geochemical features of Hangou granitic body in Xiaoshan Mountain, western Henan Province, and its geologic implications

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**Abstract:** Hangou monzonitic granite porphyry body lies in northern Xiaoshan Mountain, western Henan Province. Petrogenetic

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research on Hangou rock body is important for recognizing tectonic evolution in northern Xiaoshan Mountain area, and it is also significant for estimating mineralization potential. The zircon U-Pb dating method of LA-ICP-MS was used to confirm the age of biotite monzonitic granite porphyry sample named LLHG01 from Hangou rock body. The results of 20 zircon spots are consistent with  $^{207}\text{Pb}/^{235}\text{U}$ - $^{206}\text{Pb}/^{238}\text{U}$  trend line, with  $^{206}\text{Pb}/^{238}\text{U}$  ages being from 142 Ma to 147 Ma, and weighted average age being  $(145.1 \pm 0.7)$  Ma. Hangou rock body is characterized by higher  $\text{SiO}_2$  and alkali, enrichment of  $\text{K}_2\text{O}$ , and lower  $\text{MgO}$  and  $\text{CaO}$ . Compositional spots of Hangou rock body fall into high-K calc-alkaline series in the  $\text{SiO}_2-\text{K}_2\text{O}$  diagram, and it belongs to K-rich calc-alkaline granite (KCG). In chondrite-normalized REE patterns, Hangou rock body samples have shown the characteristics of LREE-enrichment and HREE depletion with trivial Eu positive anomaly, whose  $(\text{La/Yb})_{\text{N}}$  ratios are from 10.41 to 11.51. Samples from Hangou rock body are enriched in large ion lithophile elements and depleted in high field-strength elements, which has formed troughs of Nb, Ta and Ti in spider diagram. Samples of Hangou rock body have higher Sr and lower Y and Yb, suggesting that Hangou rock body is adakite. According to the content of  $\text{Na}_2\text{O}$  and  $\text{K}_2\text{O}$ , Hangou rock body is C-type adakite. Hangou rock body originated from partial melting of the thickened lower crust, and its residual phases of partial melting source included garnet and rutile, with no plagioclase. North Xiaoshan Mountain area underwent the process of lithosphere delamination of different scales in ~145 Ma and ~130 Ma, and Hangou rock body is a product of smaller scale lithosphere delamination that occurred at ~145 Ma. The formation age of Hangou rock body falls into the time interval of ~145 Ma stage of acid magmatism and endogenetic mineralization in western Henan Province. In the disaster process of lithosphere delamination in Hangou area, huge deep fluids were driven and discharged rapidly, suggesting that Hangou rock body and its concealed part have large mineralization potential. There are 11 Au deposits and Au mineralization spots around Hangou rock body, which constitute the best exploration evidence.

**Key words:** Hougou monzonitic granite porphyry; LA-ICP-MS zircon dating; adakite; lithosphere delamination; mineralization potential; Xiaoshan Mountain

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## 1 引言

内生金属矿床的形成与岩浆及深部流体关系密切,是否存在岩浆作用的产物被视为评价其成矿潜力的要素之一。豫西小秦岭—崤山—熊耳山地区是河南省重要的燕山期内生金属矿产地之一,其内广泛出露与成矿作用密切相关的燕山期中-酸性侵入岩(罗铭玖等,2000)。不容忽视的是,小秦岭—崤山—熊耳山成矿带内成矿作用出现了明显的分布不均,小秦岭地区和熊耳山地区成矿作用强烈、矿床规模大,位于它们之间的崤山地区已探知的成矿效应明显较弱。崤山北部出露龙卧沟、后河、白石崖、小妹河和韩沟等燕山期中-酸性小岩体,除韩沟岩体的K-Ar年龄为 $(143.8 \pm 2.1)$  Ma<sup>①</sup>以外,其余4个岩体被认为是~130 Ma崤山北部岩石圈拆沉作用的产物,且它们的深部及周缘均具有较大的成矿潜力(卢仁等,2013,2014; Liang et al.,

2013; 梁涛和卢仁,2015)。因此,精确限定韩沟岩体的形成时代和构筑其成因模型,不仅可为崤山北部燕山期深部构造演化提供新约束,而且可为崤山区域找矿提供新的思路。

## 2 地质特征

崤山位于华北克拉通南缘(图1a),主要出露太古宇太华群变质岩系、中元古界熊耳群火山岩系及中元古界官道口群碳酸盐岩(图1b)。崤山北部形成以太华群为核部、熊耳群为两翼的穹隆背斜,中部和南部分别出露杜关向斜和夜长坪背斜;断裂构造以近东西向为主,如川口—宫前断裂、潘河—马超营断裂和银家沟—南坪断裂,其次为近北东向和近北西向(河南省地质矿产局,1989)。

崤山地区除广泛出露中元古界熊耳群火山岩以外,还分布少量的中元古代辉绿岩、新元古代晋宁期闪长岩、古生代闪长岩和燕山期(石英)闪长

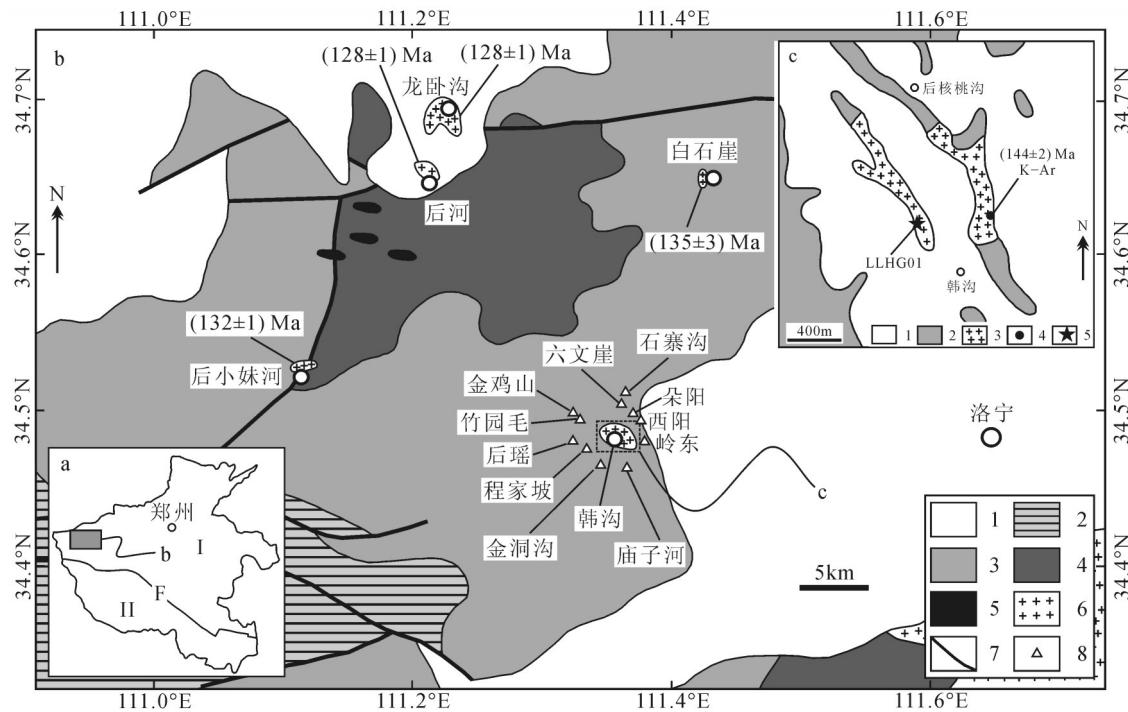


图1豫西崤山韩沟岩体地质简图

a—河南省构造分区简图(据河南省地质矿产局,1989简化);I—华北克拉通,II—秦岭造山带,F—栾川—确山—固始深大断裂带;b—豫西崤山北部地质简图(据河南省地质矿产局,1989简化):1—新生界,2—元古宇官道口群,3—元古宇熊耳群,4—太古宇太华群,5—元古宙辉绿岩,6—中生代花岗岩,7—断层,8—Au矿床(点)(据刘传权,2000简化修改),图中龙卧沟岩体和后河岩体的形成时代据卢仁等(2013,2014),白石崖岩体的形成时代据Liang et al.,(2013),小妹河岩体的形成时代据梁涛和卢仁(2015);c—韩沟岩体地质简图(据文献①简化):1—新生界,2—元古界熊耳群火山岩系,3—中生代花岗岩,4—韩沟岩体K-Ar法定年样品位置及结果,5—样品LLHG01取样位置

Fig. 1 Simplified geological map of Hangou rock body in Xiaoshan Mountain, western Henan Province

a—Simplified tectonic geological map of Henan Province (after Bureau of Geology and Mineral Resources of Henan Province, 1989): I—North China craton, II—Qinling orogen belt, F—Luanchuan—Queshan—Gushi huge fault belt; b—Simplified geological map of north Xiaoshan Mountain, western Henan Province (after Bureau of Geology and Mineral Resources of Henan Province, 1989): 1—Cenozoic, 2—Guandaokou Group (Pt<sub>2</sub>), 3—Xiong'er Group (Pt<sub>2</sub>), 4—Taihua Group (Ar), 5—Diabase (Pt<sub>2</sub>), 6—Mesozoic granite, 7—Fault, 8—Au deposit (after Liu, 2000), ages of Longwugou body and Houhe rock body after Lu et al., (2013, 2014), Baishiya after Liang et al., (2013) and Xiaomeihe after Liang and Lu (2015); c—Simplified geological map of Hangou rock body (after reference ①): 1—Cenozoic, 2—Xiong'er Group (Pt<sub>2</sub>), 3—Mesozoic granite, 4—Sampling location and result of K-Ar dating of Hangou rock body, 5—Sampling location of LLHG01

岩、花岗(斑)岩小岩体及正长斑岩岩脉(河南省地质矿产局,1989),其中在崤山北部出露有龙卧沟、后河、小妹河、白石崖和韩沟等燕山期酸性小岩珠(图1b)。崤山北部主要形成石英脉型和构造蚀变岩型Ag、Au、Pb矿床,如位于龙卧沟岩体和后河岩体周缘的半宽Au(Pb)矿和申家窑Au矿,韩沟岩体周缘的石寨沟、岭东、后瑶等Au矿床(点)(图1b)。

韩沟岩体覆盖严重,呈不规则带状出露于韩沟两侧,面积约1.5 km<sup>2</sup>,与熊耳群火山岩系呈侵入接触关系,其余为新生界不整合覆盖(图1c)。韩沟岩体以二长花岗斑岩为主,新鲜面呈浅灰白—浅灰红色,块状构造,斑状结构(图2a和2c)。斑晶含量为15%~20%,以石英、斜长石和碱性长石为主,石英斑晶多呈浑圆状,斜长石斑晶发育聚片双晶(图2c),

部分可见斜长石环带,发育一定程度的绢云母化和黏土化蚀变,碱性长石(以钾长石为主)斑晶发育卡氏双晶,发育一定程度的黏土化蚀变。基质为细粒半自形粒状结构,含量为80%~85%,以石英(20%~25%)、斜长石(25%~35%)、碱性长石(30%~40%)为主,可见黑云母及角闪石(~5%),副矿物以磁铁矿、磷灰石、锆石、黄铁矿等为主。韩沟岩体的局部地段出现钾长石巨斑晶(图2b),含量10%~15%,长轴长度以2~5 cm者居多,发育卡氏双晶。

### 3 锆石LA-ICP-MS测试

#### 3.1 样品特征

韩沟岩体锆石定年样品LLHG01的岩性为黑云母二长花岗斑岩,呈灰红—灰白色,块状构造,斑状

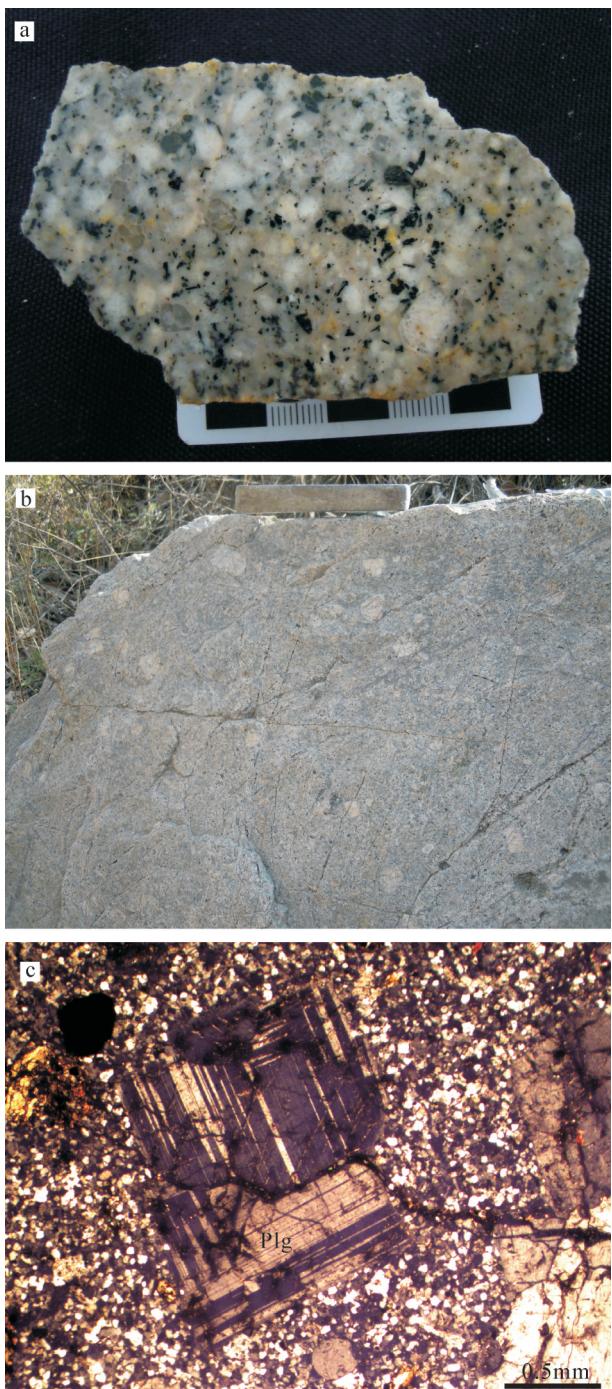


图2 韩沟岩体岩石照片

a—黑云母二长花岗斑岩;b—钾长石巨斑晶;c—斜长石斑晶,发育聚片双晶(正交偏光)

矿物代号:Plg—斜长石

Fig.2 Petrographs of Hangou rock body

a—Biotite monzonitic granite porphyry;b—K-feldspar megacrysts in Hangou intrusion;c—Plagioclase and its polysynthetic twin; Plg—Plagioclase

结构。斑晶含量约25%,以石英(约20%)、斜长石(约35%)和碱性长石(约40%)为主,石英斑晶多呈浑圆状,斜长石斑晶发育聚片双晶,可见其环带构造,发育一定程度的绢云母化和黏土化蚀变,碱性长石(以钾长石为主)斑晶发育卡氏双晶,发育一定程度的黏土化蚀变,可见2颗钾长石巨斑晶,粒径分别为 $1.5\text{ cm} \times 1.0\text{ cm}$ 和 $2.5\text{ cm} \times 1.0\text{ cm}$ 。基质含量约70%,由斜长石(约30%)、石英(约20%)、碱性长石(约35%)和黑云母(约5%)组成,具细粒半自形粒状结构。副矿物以磁铁矿、磷灰石、锆石等为主。

从样品LLHG01中分选获得的锆石数量大于1000粒,大部分无色透明,个别锆石呈浅淡黄色,部分锆石可见裂纹和暗色包裹体。大多数锆石晶型完整,多呈柱状,可见近等轴状,锆石长轴粒径以 $100\sim 150\text{ }\mu\text{m}$ 为主,短轴粒径以 $50\sim 100\text{ }\mu\text{m}$ 为主,大多数锆石发育清晰、致密、均匀的韵律震荡环带(图3),为典型岩浆锆石。

### 3.2 分析方法

样品LLHG01的锆石分选由河北省廊坊区域地质调查所实验室完成,将分选后的锆石粘贴在环氧树脂表面打磨抛光制靶,对样品靶中的锆石颗粒进行反射光、透射光及阴极发光(CL)照相,综合设计最恰当的测点。

样品LLHG01的LA-ICP-MS锆石定年测试在中国地质科学院矿产资源所LA-ICP-MS实验室完成。剥蚀系统的激光发射器使用美国New Wave公司生产的UP213型激光器,使用高纯He气作为载气将激光剥蚀物送入等离子质谱仪,质谱仪型号为德国Finnigan公司的Neptune型多接收等离子体质谱仪。定年测试中,标样GJ-1、Plesovice和样品锆石有序穿插进行测试。数据处理使用ICPMSDataCal程序(Liu et al., 2006),U-Pb协和图绘制和加权平均年龄计算使用Isoplot程序(Ludwing, 2001)。

### 3.3 锆石U-Pb定年结果

样品LLHG01的U-Pb锆石30个测点的LA-ICP-MS定年结果见表1和图4。

测点LLHG01-26的U-Pb同位素比值明显偏离 $^{207}\text{Pb}/^{235}\text{U}-^{206}\text{Pb}/^{238}\text{U}$ 一致线,测点LLHG01-11,-12,-13,-16,-20,-21,-27,-29和-30的 $^{207}\text{Pb}/^{235}\text{U}$ 年龄的误差限超过11 Ma(图4a),所以在计算加权平均年龄时,它们未考虑在内。其余20个测点的测

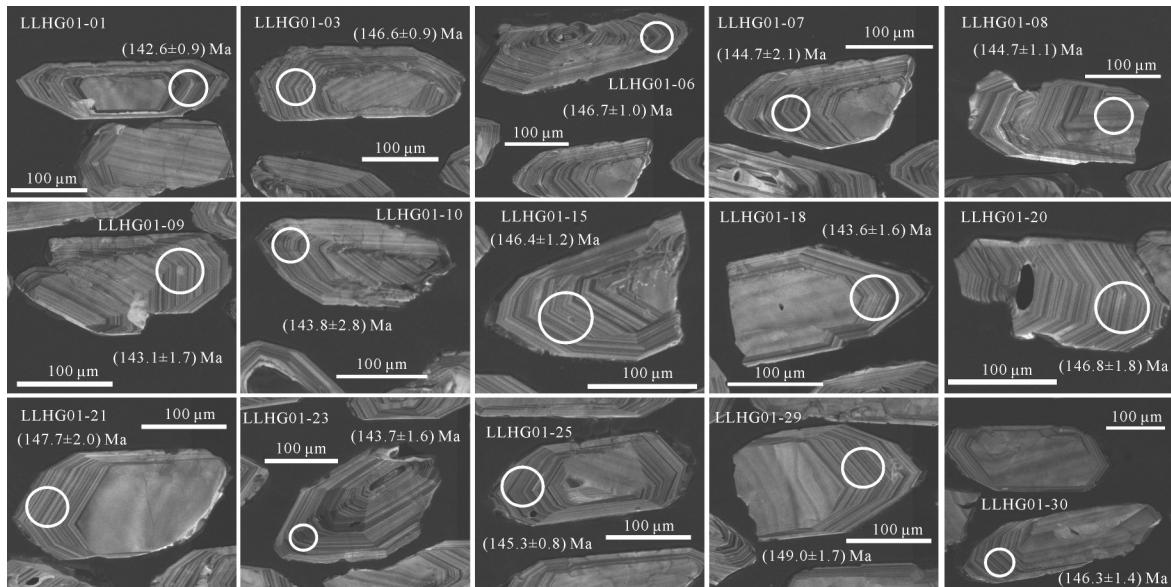


图3 韩沟岩体样品LLHG01锆石阴极发光图像

Fig.3 Cathodoluminescence images of zircons in LLHG01 from Hangou rock body

定值均位于 $^{207}\text{Pb}/^{235}\text{U}-^{206}\text{Pb}/^{238}\text{U}$ 一致线上,其 $^{206}\text{Pb}/^{238}\text{U}$ 年龄介于142~147 Ma, Th/U比值介于0.17~0.36,它们的 $^{206}\text{Pb}/^{238}\text{U}$ 加权平均年龄为(145.1±0.7) Ma(95%置信度),MSWD=1.9(图4b)。结合上述样品锆石CL图像特征,(145.1±0.7) Ma可以解释为岩浆锆石的结晶年龄。

#### 4 元素地球化学

对韩沟岩体的8件样品进行了主量元素和微量元素分析,测试均由河南省有色金属地质勘查总院检测中心完成。它们的主量元素及Ba、Rb、Sr和Zr使用ZSX Primus II型X-射线荧光光谱仪分析测定,FeO含量使用湿化学方法单独测定,其余微量元素使用Thermo Fisher X Series 2等离子体质谱仪(ICP-MS)测定。分析过程中采用国家一级标样控制准确度及精密度,用重复性密码分析及异常点抽检来验证其可靠性,其质量控制参数均合格。

##### 4.1 主量元素

8件韩沟岩体样品的主量元素含量测试结果见表2。

韩沟岩体8件样品的 $\text{SiO}_2$ 含量介于67.70%~70.46%, $\text{Al}_2\text{O}_3$ 含量的最高值和最低值分别为15.80%和13.94%, $\text{MgO}$ 含量介于0.47%~0.82%, $\text{CaO}$ 含量范围是2.27%~2.79%, $\text{Na}_2\text{O}$ 和 $\text{K}_2\text{O}$ 的含量

分别介于2.38%~3.65%和4.14%~7.07%(表2)。

在Harker图解中(图5),韩沟岩体样品 $\text{MgO}$ 、 $\text{Al}_2\text{O}_3$ 、 $\text{TFe}_2\text{O}_3$ 、 $\text{MnO}$ 、 $\text{CaO}$ 和 $\text{Na}_2\text{O}$ 含量投点显示了一定程度的负相关关系, $\text{TiO}_2$ 和 $\text{P}_2\text{O}_5$ 含量投点整体趋于发散,它们的 $\text{K}_2\text{O}$ 含量投点总体趋势为正相关关系。在 $\text{SiO}_2-(\text{Na}_2\text{O}+\text{K}_2\text{O})$ 图解中(图6a),韩沟岩体样品投点均落入5区(花岗闪长岩)、6区(花岗岩)和11区(石英二长岩)相接区域附近,属于亚碱性系列岩石。在 $\text{SiO}_2-\text{K}_2\text{O}$ 图解中(图6b),样品投点落入高钾钙碱性系列区域内,并且它们具有近垂直增涨的趋势。

##### 4.2 微量元素

8件韩沟岩体样品的微量元素含量测试结果见表2。它们的稀土元素总量介于 $95.95 \times 10^{-6}$ ~ $124.52 \times 10^{-6}$ , $(\text{La}/\text{Yb})_{\text{N}}$ 范围为10.41~11.51。韩沟岩体样品的稀土配分模式具有右倾、轻稀土富集及重稀土亏损的特征(图7a),它们显示了弱正Eu异常, $\delta\text{Eu}$ 介于1.16~1.60。

韩沟岩体样品富集大离子亲石元素(LILE),如Rb含量 $66.33 \times 10^{-6}$ ~ $98.97 \times 10^{-6}$ ,Th含量介于 $3.86 \times 10^{-6}$ ~ $4.81 \times 10^{-6}$ ,亏损高场强元素Nb、Ta,在微量元素蜘蛛图上显示明显的Nb、Ta及Ti负异常(图7b)。

韩沟岩体8件样品具有高Sr( $790 \times 10^{-6}$ ~ $1050 \times 10^{-6}$ )、低Y( $11.68 \times 10^{-6}$ ~ $15.33 \times 10^{-6}$ )和低Yb( $1.38 \times 10^{-6}$ ~ $1.71 \times 10^{-6}$ )的特征,在Y-Sr/Y图解中样品点均

表1 韩沟岩体黑云母二长花岗斑岩样品LLHG01锆石LA-ICP-MS定年结果

Table 1 LA-ICP-MS dating results of zircons in LLHG01 biotite monzonitic granite porphyry of from Hangou rock body

测点	含量/ $10^{-6}$			同位素比值						表面年龄/Ma					
	Th	U	Th/U	$^{207}\text{Pb}/^{206}\text{Pb}$	$1\sigma$	$^{207}\text{Pb}/^{235}\text{U}$	$1\sigma$	$^{206}\text{Pb}/^{238}\text{U}$	$1\sigma$	$^{207}\text{Pb}/^{206}\text{Pb}$	$1\sigma$	$^{207}\text{Pb}/^{235}\text{U}$	$1\sigma$	$^{206}\text{Pb}/^{238}\text{U}$	$1\sigma$
LLHG01-1	196.69	546.05	0.36	0.0502	0.0014	0.1546	0.0043	0.0224	0.0001	205.6	95.4	146.0	3.8	142.6	0.9
LLHG01-2	199.31	638.16	0.31	0.0491	0.0012	0.1525	0.0039	0.0225	0.0001	150.1	52.8	144.1	3.4	143.7	0.9
LLHG01-3	59.81	261.25	0.23	0.0506	0.0023	0.1610	0.0074	0.0230	0.0001	233.4	100.9	151.6	6.5	146.6	0.9
LLHG01-4	113.10	502.67	0.23	0.0492	0.0026	0.1566	0.0085	0.0230	0.0002	166.8	121.3	147.7	7.5	146.9	1.0
LLHG01-5	97.25	418.95	0.23	0.0490	0.0014	0.1537	0.0046	0.0228	0.0001	146.4	70.4	145.2	4.1	145.2	0.8
LLHG01-6	104.57	433.36	0.24	0.0517	0.0020	0.1638	0.0062	0.0230	0.0002	272.3	91.7	154.0	5.4	146.7	1.0
LLHG01-7	150.97	500.94	0.30	0.0489	0.0031	0.1540	0.0114	0.0227	0.0003	146.4	140.7	145.5	10.0	144.7	2.1
LLHG01-8	73.24	299.08	0.24	0.0513	0.0030	0.1609	0.0096	0.0227	0.0002	253.8	130.5	151.5	8.4	144.7	1.1
LLHG01-9	52.78	272.94	0.19	0.0498	0.0038	0.1550	0.0121	0.0224	0.0003	187.1	166.6	146.3	10.6	143.1	1.7
LLHG01-10	96.56	388.69	0.25	0.0513	0.0038	0.1597	0.0121	0.0226	0.0004	253.8	168.5	150.5	10.6	143.8	2.3
LLHG01-11	47.78	279.76	0.17	0.0501	0.0047	0.1547	0.0150	0.0224	0.0004	211.2	194.4	146.1	13.2	143.0	2.7
LLHG01-12	65.00	362.10	0.18	0.0514	0.0080	0.1569	0.0218	0.0223	0.0005	261.2	322.2	148.0	19.2	142.0	3.1
LLHG01-13	83.93	356.63	0.24	0.0492	0.0045	0.1528	0.0145	0.0225	0.0003	153.8	212.9	144.4	12.7	143.3	2.0
LLHG01-14	148.60	575.86	0.26	0.0491	0.0036	0.1542	0.0115	0.0228	0.0001	153.8	168.5	145.6	10.1	145.1	0.6
LLHG01-15	40.90	247.74	0.17	0.0484	0.0022	0.1529	0.0071	0.0230	0.0002	120.5	112.0	144.5	6.3	146.4	1.2
LLHG01-16	69.60	301.16	0.23	0.0507	0.0061	0.1594	0.0216	0.0225	0.0004	227.8	259.2	150.2	19.0	143.4	2.7
LLHG01-17	62.57	289.72	0.22	0.0510	0.0034	0.1576	0.0111	0.0223	0.0003	242.7	151.8	148.6	9.8	142.0	2.1
LLHG01-18	127.79	453.83	0.28	0.0520	0.0021	0.1602	0.0063	0.0225	0.0002	283.4	90.7	150.9	5.6	143.6	1.6
LLHG01-19	69.28	302.23	0.23	0.0514	0.0017	0.1614	0.0056	0.0227	0.0002	261.2	74.1	151.9	4.9	144.6	1.3
LLHG01-20	42.68	236.90	0.18	0.0487	0.0069	0.1565	0.0223	0.0230	0.0003	200.1	235.2	147.6	19.6	146.8	1.8
LLHG01-21	43.66	248.24	0.18	0.0512	0.0053	0.1646	0.0177	0.0232	0.0003	250.1	222.2	154.7	15.5	147.7	2.0
LLHG01-22	65.58	280.83	0.23	0.0496	0.0033	0.1529	0.0104	0.0224	0.0002	189.0	157.4	144.4	9.2	142.6	1.2
LLHG01-23	138.82	533.58	0.26	0.0490	0.0021	0.1521	0.0066	0.0225	0.0002	146.4	100.0	143.7	5.8	143.7	1.6
LLHG01-24	52.20	245.08	0.21	0.0500	0.0023	0.1578	0.0075	0.0228	0.0002	194.5	107.4	148.8	6.6	145.2	1.2
LLHG01-25	63.48	299.58	0.21	0.0513	0.0019	0.1616	0.0061	0.0228	0.0001	257.5	85.2	152.1	5.3	145.3	0.8
LLHG01-26	54.28	294.01	0.18	0.0572	0.0018	0.1810	0.0061	0.0229	0.0001	498.2	70.4	168.9	5.2	145.7	0.8
LLHG01-27	41.56	238.38	0.17	0.0488	0.0060	0.1549	0.0193	0.0230	0.0003	200.1	205.5	146.2	17.0	146.9	1.7
LLHG01-28	44.76	214.64	0.21	0.0511	0.0029	0.1637	0.0096	0.0231	0.0001	242.7	131.5	154.0	8.3	147.0	0.9
LLHG01-29	32.05	171.39	0.19	0.0492	0.0088	0.1578	0.0280	0.0232	0.0003	166.8	361.1	148.7	24.6	148.0	1.7
LLHG01-30	67.17	309.74	0.22	0.0509	0.0043	0.1616	0.0135	0.0230	0.0002	239.0	194.4	152.1	11.8	146.3	1.4

落入埃达克岩区域内(图8a),表明其具有埃达克质岩石的属性。

## 5 讨论与结论

### 5.1 韩沟岩体形成时代

韩沟岩体LA-ICP-MS定年样品LLHG01为黑云母二长花岗斑岩,具有块状构造,从中获得的锆

石具有典型的岩浆锆石特征,且所有测点的Th/U比值大于0.1,20个锆石测点的加权平均 $^{206}\text{Pb}/^{238}\text{U}$ 年龄为( $145.1\pm0.7$ ) Ma(95%置信度)。韩沟岩体的K-Ar同位素年龄( $143.8\pm2.1$ ) Ma<sup>①</sup>,与锆石U-Pb定年结果相近,表明韩沟岩体形成于145.1 Ma。

崤山北部的龙卧沟、后河、白石崖和小妹河等岩体的LA-ICP-MS锆石U-Pb年龄分别为( $128\pm1$ )

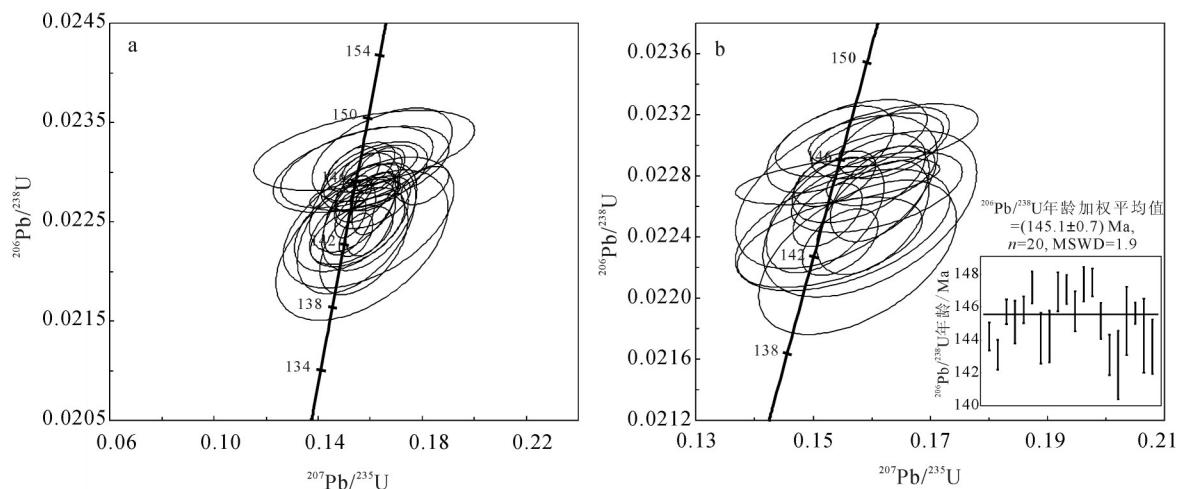


图4 韩沟岩体样品LLHG01锆石U-Pb谐和曲线及定年结果  
Fig.4 Concordia diagrams for U-Pb age of zircons in LLHG01 from Hangou rock body

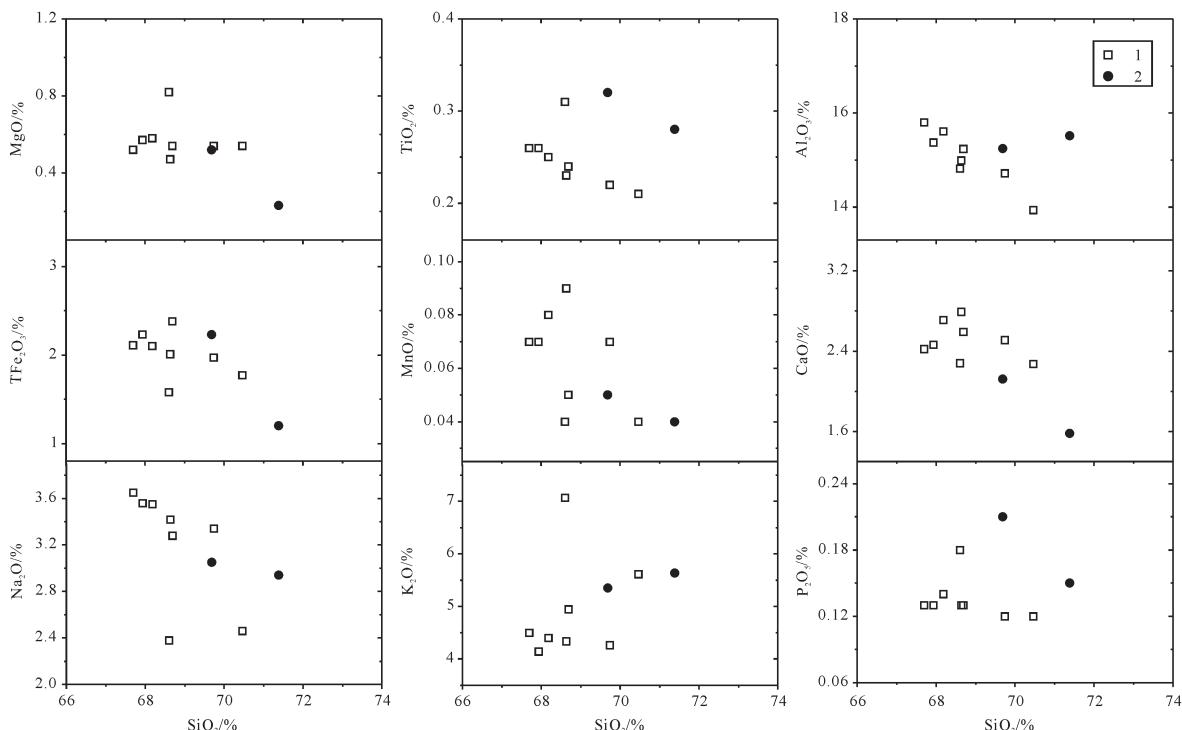


图5 韩沟岩体Harker图解  
数据1—本次研究样品;2—据李磊等,2013  
Fig.5 Harker diagram of Hangou rock body  
Data 1 are the results of the samples of this study; data 2 are from Li et al., 2013

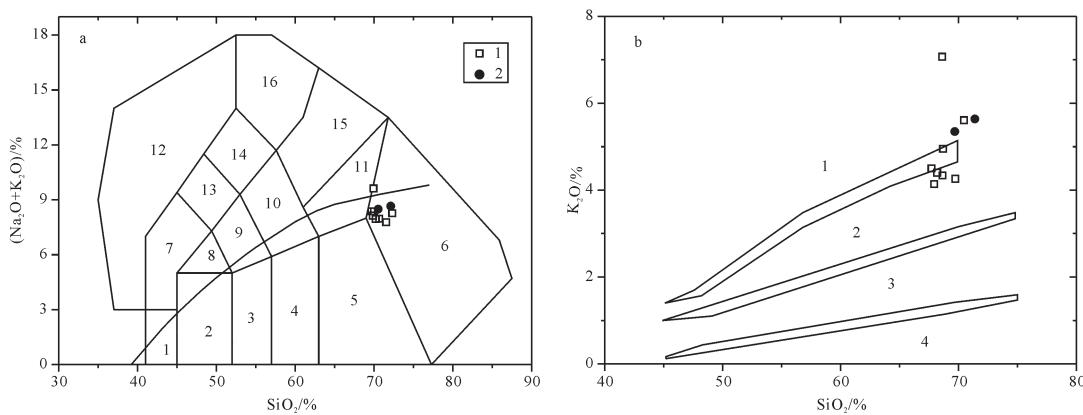
Ma、(128±1) Ma、(135±3) Ma 和(132±1) Ma(卢仁等, 2013, 2014; Liang et al., 2013; 梁涛和卢仁, 2015), 表明崤山北部存在~130 Ma 和~145 Ma 两期酸性岩浆活动, 韩沟岩体是较早一期岩浆作用的产物。

出露于崤山南部的燕山期酸性侵入岩也具有~130 Ma 和~145 Ma 两期活动的特征(表3), 这与崤山北部燕山期岩浆活动期次相一致, 进而表明崤山地区燕山期岩浆活动至少存在~130 Ma 和~145 Ma

表2 韩沟岩体二长花岗斑岩样品主量元素(%)和微量元素( $10^{-6}$ )分析结果Table 2 Major elements (%) and trace elements ( $10^{-6}$ ) analyses of monzonitic granite porphyry of Hangou rock body

样号	LLHG201	LLHG203	LLHG204	LLHG205	LLHG206	LLHG207	LLHG208	LLHG209
SiO <sub>2</sub>	68.61	68.69	70.46	68.18	67.93	69.74	67.70	68.64
TiO <sub>2</sub>	0.31	0.24	0.21	0.25	0.26	0.22	0.26	0.23
Al <sub>2</sub> O <sub>3</sub>	14.82	15.24	13.94	15.61	15.37	14.72	15.80	14.99
Fe <sub>2</sub> O <sub>3</sub>	0.16	1.19	0.94	1.03	1.28	1.38	0.80	0.94
FeO	1.29	1.08	0.75	0.97	0.86	0.54	1.19	0.97
MnO	0.04	0.05	0.04	0.08	0.07	0.07	0.07	0.09
MgO	0.82	0.54	0.54	0.58	0.57	0.54	0.52	0.47
CaO	2.28	2.59	2.27	2.71	2.46	2.51	2.42	2.79
Na <sub>2</sub> O	2.38	3.28	2.46	3.55	3.56	3.34	3.65	3.42
K <sub>2</sub> O	7.07	4.95	5.61	4.40	4.14	4.26	4.50	4.34
P <sub>2</sub> O <sub>5</sub>	0.18	0.13	0.12	0.14	0.13	0.12	0.13	0.13
烧失量	1.91	1.52	1.94	1.57	1.37	1.27	0.89	1.02
总和	99.87	99.49	99.30	99.06	98.00	98.72	97.91	98.02
V	26.87	34.68	41.24	44.65	55.91	57.78	49.44	50.6
Co	2.05	2.4	2.05	3.43	3.15	2.82	2.7	2.64
Ni	2.61	2.27	2.22	2.73	2.56	3.03	3.02	3.41
Cr	13.92	16.10	15.83	18.27	19.23	19.01	12.86	19.35
Ga	19.31	20.34	17.28	20.15	18.49	17.77	19.48	20.23
Ba	2683	1596	2365	1370	1863	1436	2235	1706
Rb	98.97	79.66	79.86	69.41	74.94	77.30	66.33	68.31
Th	3.92	4.20	3.86	4.53	4.29	4.20	4.81	4.45
Nb	9.90	10.34	9.25	11.05	10.43	10.21	11.93	11.10
Ta	0.46	0.45	0.39	0.49	0.46	0.46	0.59	0.50
Sr	833	911	790	959	964	907	1050	962
Zr	110.21	103.78	101.93	100.76	93.48	97.29	85.30	93.89
Hf	5.08	4.15	3.68	4.07	3.11	3.86	3.54	3.76
Y	15.33	13.54	11.68	13.74	13.95	13.64	14.88	13.94
La	25.70	24.54	21.61	27.14	24.71	26.77	28.89	25.93
Ce	50.98	43.84	39.20	47.18	43.40	45.71	51.03	46.04
Pr	6.24	5.09	4.56	5.47	5.11	5.30	5.84	5.45
Nd	24.07	19.46	17.32	20.51	19.56	19.84	22.25	20.47
Sm	4.01	3.35	2.91	3.46	3.31	3.29	3.73	3.47
Eu	1.86	1.32	1.51	1.29	1.44	1.28	1.63	1.40
Gd	3.74	3.18	2.78	3.23	3.16	3.07	3.52	3.27
Tb	0.53	0.45	0.39	0.45	0.45	0.45	0.49	0.46
Dy	2.94	2.50	2.16	2.51	2.54	2.50	2.74	2.55
Ho	0.58	0.49	0.43	0.50	0.50	0.49	0.53	0.50
Er	1.70	1.48	1.28	1.49	1.53	1.48	1.63	1.52
Tm	0.25	0.23	0.20	0.23	0.24	0.23	0.25	0.23
Yb	1.66	1.55	1.38	1.59	1.60	1.62	1.71	1.57
Lu	0.26	0.25	0.22	0.25	0.26	0.25	0.27	0.25
Sr/Y	54.32	67.28	67.65	69.81	69.10	66.48	70.58	68.98
$\Sigma$ REE	124.52	107.73	95.95	115.30	107.81	112.28	124.51	113.11
(La/Yb) <sub>N</sub>	10.44	10.67	10.56	11.51	10.41	11.14	11.39	11.13
$\delta$ Eu	1.45	1.22	1.60	1.16	1.34	1.21	1.36	1.25

注:  $\delta$  Eu =  $2 \times \text{Eu}_{\text{N}} / (\text{Sm}_{\text{N}} + \text{Gd}_{\text{N}})$ 。

图6 韩沟岩体SiO<sub>2</sub>-(Na<sub>2</sub>O+K<sub>2</sub>O)和SiO<sub>2</sub>-K<sub>2</sub>O图解

a—SiO<sub>2</sub>-(Na<sub>2</sub>O+K<sub>2</sub>O)图解(据 Middlemost, 1989):1—橄榄辉长岩,2—辉长岩,3—辉长闪长岩,4—闪长岩,5—花岗闪长岩,6—花岗岩,7—副长石辉长岩,8—二长辉长岩,9—二长闪长岩,10—二长岩,11—石英二长岩,12—副长石岩,13—副长石二长闪长岩,14—副长石二长正长岩,15—正长岩,16—副长石正长岩,图中碱性与亚碱性系列界线据 Irvine and Baragar, 1971; b—SiO<sub>2</sub>-K<sub>2</sub>O图解(底图据 Rollison, 1993):

1—钾玄岩系列,2—高钾钙碱性系列,3—钙碱性系列,4—低钾拉班玄武系列。图中1—本次研究样品数据,2—据李磊等,2013

Fig.6 Diagrams of SiO<sub>2</sub>-(Na<sub>2</sub>O+K<sub>2</sub>O), and SiO<sub>2</sub>-K<sub>2</sub>O for Hangou rock body

a—Diagram of SiO<sub>2</sub>-(Na<sub>2</sub>O+K<sub>2</sub>O) (base map after Middlemost, 1989), 1—Peridotite gabbro, 2—Gabbro, 3—Gabbroic diorite, 4—Diorite, 5—Granodiorite, 6—Granite, 7—Foid monzo-gabbro, 8—Monzo-gabbro, 9—Monzo-diorite, 10—Monzonite, 11—Quartz Monzonite, 12—Foidolite, 13—Foid Monzo-diorite, 14—Foid Monzo-syenite, 15—Syenite, 16—Foid syenite. The boundary of alkaline and sub-alkaline after Irvine and Baragar, 1971; b—Diagram of SiO<sub>2</sub>-K<sub>2</sub>O (base map after Rollison, 1993), 1—Shoshonite series, 2—High-K calc-alkaline series, 3—Calc-alkaline series, 4—Low-K tholeiite series. In the map, data 1 are the results of the samples of this study; data 2 are from Li et al., 2013

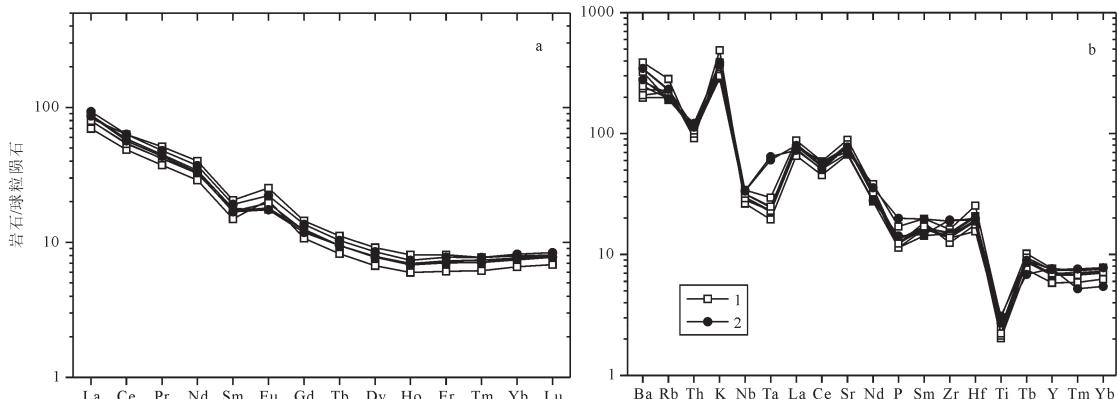


图7 韩沟岩体稀土元素配分模式图和微量元素蛛网图

a—韩沟岩体球粒陨石标准化稀土元素配分模式图,球粒陨石标准化数据据 Boynton, 1984;b—韩沟岩体微量元素蛛网图,标准化数据据 Thompson, 1982;图中1—本次研究样品数据,2—据李磊等,2013

Fig.7 Diagrams of REE distribution pattern and spider diagram for Hangou rock body

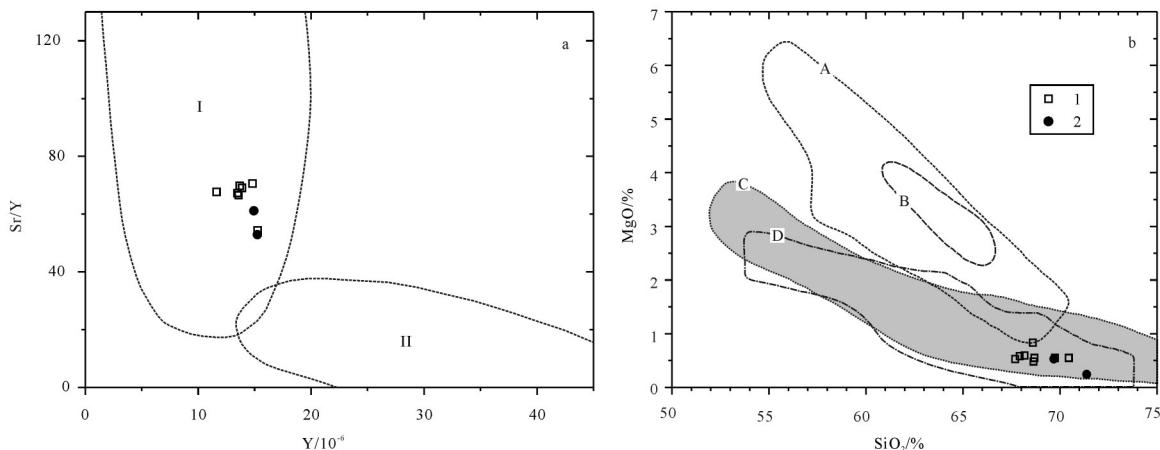
a—Chondrite-normalized REE distribution patterns of Hangou rock body, chondrite values after Boynton, 1984; b—Chondrite-normalized spider diagram of Hangou rock body, and chondrite values after Thompson, 1982. In the map, data 1 are the results of the samples of this study; data 2 are from Li et al., 2013

两期,较早一期的酸性侵入岩多分布于崤山南部,而~130 Ma的酸性岩浆活动则主要位于崤山北部。

## 5.2 部分熔融源区特征

通常认为,埃达克岩有O型和C型两种成因类型,分别与板块的俯冲作用有关和与加厚地壳发生

拆沉作用相关(张旗等,2001;罗照华等,2002;朱弟成等,2003)。韩沟岩体的SiO<sub>2</sub>、Al<sub>2</sub>O<sub>3</sub>、MgO、Sr和Yb含量特征显示其具有埃达克质岩的属性(表2和图8a),其Na<sub>2</sub>O和K<sub>2</sub>O含量又明显不同于典型的O型埃达克岩,符合C型埃达克岩的特征,表明其成因

图8 韩沟岩体的Y-Sr/Y图解和SiO<sub>2</sub>-MgO图解

a—韩沟岩体的Y-Sr/Y图解,I—埃达克岩,II—正常岛弧安山岩-英安岩-流纹岩(底图Defant et al., 2002);b—韩沟岩体的SiO<sub>2</sub>-MgO图解,A—俯冲洋壳起源的埃达克岩,B—与橄榄岩混合的变玄武岩和榴辉岩的实验熔体,C—变玄武岩和榴辉岩的实验熔体(1.0~4.0 GPa),D—加厚下地壳部分熔融起源的埃达克质岩石(底图据Wang et al., 2006),图中数据1—据本文,2—李磊等,2013

Fig.8 Diagrams of Y-Sr/Y and SiO<sub>2</sub>-MgO for Hangou rock body

a—Y-Sr/Y diagram of Hangou rock body, I—Adakites, II—Arc andesites, dacites and rhyolite (base map after Defant et al., 2002); b—Diagram of SiO<sub>2</sub>-MgO for Hangou rock body, A—Subducted oceanic crust-derived adakites, B—Metabasaltic and eclogite experimental melts hybridized by peridotite, C—Metabasaltic and eclogite experimental melts (1.0~4.0 GPa), D—Thick lower crust-derived adakitic rocks (base map after Wang et al., 2006).

In the map, data 1 are the results of the samples of this study; data 2 are from Li et al., 2013

与板块俯冲无直接联系。秦岭造山带及相邻的华北板块和扬子板块早在印支末期实现了碰撞拼接,区域构造演化进入陆(板)内阶段(卢欣祥,2000;张国伟等,2001),即崤山北部在~145 Ma不存在洋壳俯冲,再次表明韩沟岩体与板块俯冲之间不具有直接成因联系。

韩沟的SiO<sub>2</sub>含量(64.23%~66.94%)、MgO含量(0.62%~0.92%)及Mg<sup>#</sup>值(27.82~32.24)等的特征排除了通过玄武质岩浆的AFC演化形成埃达克质韩沟岩体的可能性,也否定了拆沉下地壳部分熔融形成韩沟埃达克质岩浆的成因模型。在SiO<sub>2</sub>-MgO图解中,韩沟岩体样品投点均落入加厚下地壳部分熔融起源的埃达克质岩石区域内(图8b),说明韩沟埃达克质岩浆应是加厚下地壳部分熔融形成的。

韩沟岩体的HREE亏损和低Y含量表明其部分熔融源区石榴石作为残余相出现,暗示部分熔融源区深度至少大于50 km(张旗等,2008)。Nb、Ta在金红石/长英质熔体体系中是强烈的相容元素(Xiong et al., 2006;熊小林等,2007),金红石本身也是富Ti矿物相,韩沟岩体显示了明显的Nb、Ta及Ti负异常亏损,说明其部分熔融源区的残余相内包含

金红石。斜长石是主要的富Al、Sr及控制Eu含量的重要矿物,韩沟岩体的高Al<sub>2</sub>O<sub>3</sub>和Sr含量及弱Eu正异常,这表明其源区无(富钙)斜长石残留。

所以,埃达克质的韩沟岩体是加厚下地壳的部分熔融形成的,部分熔融源区的残余相矿物包括石榴石和金红石,且无(富钙)斜长石残留。

### 5.3 深部构造背景

韩沟岩体内出现钾长石巨斑晶,副矿物以磁铁矿、磷灰石、锆石、黄铁矿等为主,且属于高钾钙碱性系列岩石,符合富钾钙碱性花岗岩类(KCG类)的特征(Barbarin, 1999),表明韩沟二长花岗斑岩属于KCG类花岗岩。此类花岗岩形成于构造体制转化阶段,它是壳、幔混合来源的钙碱性花岗岩类,富K<sub>2</sub>O贫CaO的地壳端元为主要组成部分(Barbarin, 1999;肖庆辉和卢欣祥,2002)。

印支期末期,华北板块和扬子板块实现了碰撞拼接(卢欣祥,2000;张国伟等,2001),这表明在~145 Ma时崤山北部的区域构造演化进入陆(板)内阶段。邓晋福等(1996, 1999, 2002, 2004, 2007)认为在碰撞过程中完成大陆拼接之后,岩石圈的汇聚仍在持续中,致使陆内块体间的碰撞作用或者陆内

碰撞造山过程得以显现,之后的深部构造过程为岩石圈垮塌。由各种块体经长期反复拼合而成的大陆板块具有复杂的岩石圈不均一性和重力不稳定性,在统一的地幔对流体制下,这种不均一性将导致各个组成块体产生不同的深部响应(罗照华等,2007)。

韩沟埃达克质岩的部分熔融源区深度大于50 km,表明~145 Ma时韩沟地区地壳厚度也应不小于50 km,现如今崤山北部莫霍面埋深约35 km(周国藩,1992;袁学诚,1996),二者之间至少15 km的差距足以表明韩沟地区在~145 Ma经历了岩石圈拆沉作用,这致使形成于加厚地壳中的韩沟埃达克质岩浆得以释放。崤山北部在~130 Ma经历了岩石圈拆沉作用(卢仁等,2013,2014;Liang et al., 2013;梁涛和卢仁,2015),表明崤山北部至少经历了~145 Ma和~130 Ma两次岩石圈拆沉作用过程,考虑到崤山北部出露的~145 Ma酸性侵入岩在数量和规模上小

于~130 Ma的,认为~145 Ma崤山北部地区的岩石圈拆沉规模弱于~130 Ma的岩石圈拆沉规模,前者可能是~130 Ma崤山北部地区大规模岩石圈拆沉作用的前奏。

#### 5.4 成矿潜力

崤山南部不仅出露与韩沟岩体形成时代相近的酸性岩浆活动,而且还出露与之近同时的内生金属成矿作用(表3)。银家沟大型斑岩-矽卡岩型硫铁多金属矿中辉钼矿的Re-Os模式年龄介于(142.9±2.1) Ma ~ (143.4±2.3) Ma, 加权平均年龄为(143.4±0.9) Ma, 绢云母的<sup>40</sup>Ar/<sup>39</sup>Ar等时线年龄为(143.0±2.0) Ma(武广等,2013)。夜长坪钼矿为大型斑岩-矽卡岩型钼矿床,它的辉钼矿Re-Os模式年龄介于(143.6±2.4) Ma ~ (145.4±2.4) Ma(毛冰等,2011)和(144.2±2.0) Ma~(147.2±2.5) Ma(晏国龙等,2012), 加权平均年龄分别为(144.4±1.0) Ma 和(145.3±0.8) Ma。

**表3 嵩山地区酸性侵入岩年龄**  
**Table 3 Ages of acid intrusions in Xiaoshan Mountain area**

序号	出露区域	岩体名称	定年方法	测试结果	数据来源
1	崤山北部	韩沟	锆石 LA-ICP-MS	(145.1±0.7) Ma	本文
2	崤山北部	龙卧沟	锆石 LA-ICP-MS	(128±1) Ma	卢仁等, 2014
3	崤山北部	后河	锆石 LA-ICP-MS	(128±1) Ma	卢仁等, 2013
4	崤山北部	白石崖	锆石 LA-ICP-MS	(145±2) Ma 和(135±3) Ma	Liang et al., 2013
5	崤山北部	小妹河	锆石 LA-ICP-MS	(131.5±0.9) Ma	梁涛和卢仁, 2015
6	崤山南部	蒲阵沟	锆石 LA-ICP-MS	(149±1) Ma	胡浩等, 2011
				(142.0±2.0) Ma (花岗闪长斑岩)	
			锆石 SHRIMP	(147.5±2.1) Ma (二长花岗斑岩)	李铁刚等, 2013
7	崤山南部	银家沟		(147.8±1.6) Ma (钾长花岗斑岩)	
			辉钼矿 Re-Os	(143.4±0.9) Ma	武广等, 2013
			绢云母 Ar-Ar	(143.0±2.0) Ma	武广等, 2013
8	崤山南部	圪老湾	锆石 LA-ICP-MS	(148±1) Ma	胡浩等, 2011
				(145±2) Ma (4个测点)	
			锆石 LA-ICP-MS*	(158±2) Ma (11个测点)	胡浩等, 2011
9	崤山南部	夜长坪		(170.0±4.6) Ma (3个测点)	
			辉钼矿 Re-Os	(145.3±4.4) Ma	毛冰等, 2011
				(144.9±1.0) Ma	晏国龙等, 2012
10	崤山南部	八宝山	锆石 LA-ICP-MS	(145.9±1.9) Ma 和(146.6±1.6) Ma	曾令君等, 2013
11	崤山南部	后蟠峪	锆石 LA-ICP-MS	(131±1) Ma	胡浩等, 2011
12	崤山南部	柳关	锆石 LA-ICP-MS	(132±1) Ma	胡浩等, 2011

注:夜长坪岩体样品YCP1累计分析了18测点,年龄集中于3组,本文认为应以最晚1组锆石年龄限定岩体形成时代,即年龄介于(143±1) Ma ~ (148±2) Ma的4颗锆石,它们的加权平均年龄为(145±2) Ma。

沙沟大型Ag多金属矿床位于韩沟岩体西南约25 km处,其绢云母和铬云母的 $^{40}\text{Ar}/^{39}\text{Ar}$ 等时线年龄分别为( $145.2\pm2.5$ ) Ma和( $147.6\pm2.3$ ) Ma(毛景文等,2006)。此外,南泥湖岩体的锆石U-Pb年龄为( $149.6\pm0.4$ ) Ma(包志伟等,2009)和( $146.7\pm1.2$ ) Ma(向君峰等,2012),花岗斑岩脉的LA-ICP-MS锆石U-Pb年龄为( $145.2\pm1.5$ ) Ma和( $145.7\pm1.2$ ) Ma(向君峰等,2012)。南泥湖—三道庄钼钨矿的辉钼矿Re-Os模式年龄介于( $141.8\pm2.1$ ) Ma~( $145.4\pm2.0$ ) Ma(李永峰等,2003)和( $143.4\pm2.0$ ) Ma~( $146.5\pm2.5$ ) Ma(向君峰等,2012),后者的加权平均年龄为( $145.0\pm0.7$ ) Ma。

由此可见,~145 Ma是豫西地区重要的成岩成矿期之一,韩沟岩体的形成时代(( $145.1\pm0.7$ ) Ma)正好与之重合,表明存在与韩沟岩体密切相关的成矿作用,庙子河金矿(刘传权,2000)和石寨沟金矿(赖勇等,2008;陈书中和喻广建,2010;刘军等,2012;叶杰和李滨,2014)就是例证。埃达克质韩沟岩体的深部构造背景为岩石圈拆沉作用,此灾变过程必然导致深部流体的大规模活动及其快速释放(罗照华等,2008,2009),绕韩沟岩体呈环状出露石寨沟等11个Au矿床(点)(图1b),这些都表明韩沟岩体周缘及深部应具有较大的Au成矿潜力,应视其为崤山北部有望获得找矿突破的地段区之一。

综上所述,获得以下结论:

(1)韩沟岩体黑云母二长花岗斑岩样品的LA-ICP-MS锆石 $^{206}\text{Pb}/^{238}\text{U}$ 年龄为( $145.1\pm0.7$ ) Ma。

(2)韩沟岩体具有埃达克质岩的特征,它是加厚下地壳的部分熔融形成的,部分熔融源区的残余相矿物包括石榴石和金红石,且无(富钙)斜长石残留。

(3)韩沟岩体是崤山北部~145 Ma小规模岩石圈拆沉作用的产物,这利于深部流体的快速释放,其形成时代也与区域上~145 Ma成岩成矿期相重合,表明韩沟岩体周缘及深部应具有较大的Au成矿潜力。

## 注释

① 河南省地质矿产厅. 1995. 长水幅1:50000地质图说明书[R]. 1~52.

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