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## 江西省都昌县阳储岭钨钼矿年代学研究

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**提要:**江西省都昌县阳储岭钨钼矿床位于江南造山带东段,为准确厘定阳储岭钨钼矿床的成岩成矿时代,本次在详实的野外地质调查基础上,对阳储岭成矿花岗闪长斑进行锆石U-Pb定年工作,同时开展辉钼矿Re-Os同位素定年研究。LA-ICP-MS锆石U-Pb年龄为( $145.08\pm0.35$ )Ma(MSWD=0.51, n=7);辉钼矿Re-Os模式年龄为( $143.3\pm2.0$ )Ma~( $145.5\pm2.2$ )Ma,等时线年龄为( $145.4\pm1.0$ )Ma。阳储岭成矿花岗闪长斑岩锆石U-Pb年龄与辉钼矿Re-Os等时线年龄基本一致,表明阳储岭钨钼矿床成岩成矿时代约为145 Ma。辉钼矿中Re含量为 $16.62\times10^{-6}$ ~ $87.76\times10^{-6}$ ,平均值为 $44.68\times10^{-6}$ ,与壳幔混源岩浆热液矿床中Re的含量相似,指示阳储岭钨钼矿床成矿物质来源于壳幔混源。

**关 键 词:**阳储岭;锆石U-Pb;辉钼矿Re-Os;钨钼矿;江西

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## Geochronology of the Yangchuling tungsten-molybdenum deposit in Duchang County, Jiangxi Province

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**Abstract:** The Yangchuling tungsten-molybdenum deposit in Duchang County, Jiangxi Province, is located in the eastern part of the Jiangnan orogenic belt. In order to accurately define the diagenetic and mineralization ages of the Yangchuling W-Mo deposit, the authors carried out the zircon U-Pb dating of mineralization-associated granodiorite porphyry and Re-Os isotope dating of molybdenite in the Yangchuling deposit based on detailed field geological studies. The zircon LA-ICP-MS U-Pb age is ( $145.08\pm0.35$ )Ma (MSWD=0.03, n=22), the Re-Os model ages of molybdenite range from ( $143.3\pm2.0$ )Ma to ( $145.5\pm2.2$ )Ma, and the Re-Os isochron age is ( $145.4\pm1.0$ )Ma. The geochronological consistency between the ore-bearing porphyry and molybdenite indicates that the diagenesis and mineralization ages of the Yangchuling tungsten-molybdenum deposit are about 145 Ma. The  $\omega(\text{Re})$  values of molybdenite range from  $16.62\times10^{-6}$  to  $87.76\times10^{-6}$ , averaging  $44.68\times10^{-6}$ , similar to Re content in crust-mantle mixed source

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magmatic hydrothermal deposits, indicating that the ore-forming materials of the Yangchuling W-Mo deposit should have come from crust-mantle mixed source.

**Key words:** Yangchuling; zircon U-Pb; molybdenite Re-Os; tungsten-molybdenum deposit; Jiangxi

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## 1 引言

中国钨矿床成矿岩浆岩主要为S型花岗岩(徐克勤等,1982;顾连兴,1990;刘家军等,2007;张云政等,2009),而国外与钨矿成矿有关的花岗岩主要是I型花岗岩(Singoyi and Zaw, 2001)。赣北都昌县阳储岭大型斑岩型钨矿(白钨矿为主)是中国少有的成矿岩浆岩为I型花岗岩的钨矿床(迟实福,1987;满发胜等,1988)。前人针对阳储岭矿床的研究主要集中于矿体特征、岩石学、矿物学、成矿流体学、矿床成因(陈炳才等,1984;张玉学等,1985;迟实福等,1985;王小松等,1985;刘英俊等,1985)。满发胜等(1988)对阳储岭矿区成矿斑岩体进行过年代学研究,满发胜等(1988)测得阳储岭矿区含矿斑岩体全岩Rb-Sr年龄为( $140.6\pm0.6$ )Ma、含矿斑岩黑云母K-Ar年龄为( $134.0\pm4.7$ )Ma,而迟实福等(1990)测得含矿斑岩体的黑云母K-Ar年龄为175~157 Ma,两者测年结果相差较大。阳储岭成矿岩体为杂岩体,具有多次侵位特征(迟实福等,1985;李秉伦等,1985),而全岩Rb-Sr与黑云母K-Ar测年受外界干扰因素较大。因此,阳储岭矿床的成矿岩体年龄仍不确定,且前人没有进行成矿年代学研究,导致成岩与成矿关系不清楚,从而制约了对矿床成因的认识。本次拟通过阳储岭矿区的详细调查,采用封闭体系好、准确性高的锆石U-Pb测年与辉钼矿Re-Os同位素定年手段,重新厘定阳储岭矿床成岩、成矿时代,以期促进赣北地区钨矿勘查与研究工作。

## 2 区域地质

阳储岭钨钼矿床位于江南造山带东段九岭—鄣公山隆起带(图1)。隆起带广泛分布中元古界双桥山群浅变质岩和晋宁期九岭花岗岩基,极少地区

残留南华纪莲沱组地层。区内岩浆活动强烈,除晋宁期九岭花岗岩基外,主要产出燕山期壳源重熔型酸性侵入岩,部分为壳幔重熔型(如阳储岭岩体),形成九岭—鄣公山燕山期花岗岩带,岩体多为岩株状杂岩体,成岩年龄145~110 Ma(黄兰椿等,2012;胡正华等,2018a,2018b,2018c)。岩石类型主要为黑云母花岗岩—浅色花岗岩(如大湖塘、彭山、徐家尖等)。区内经历了长期多次构造运动,其中燕山运动陆内强烈造山作用,导致大规模岩浆侵入与金属成矿大爆发。构造形变主要表现为近东西向—北东东向构造与北北东向(新华夏系)构造复合特征。隆起带中次级东西向—北东东向片理带、断裂带、韧性剪切带十分发育,其与北北东向构造(如武宁—莲花、东至—龙南、鹅湖—宜黄等断裂带)复合,控制大湖塘、阳储岭、莲花山、鹅湖等钨、金矿集区及矿床展布。伴随构造演化和燕山期强烈岩浆活动,区内形成以与壳熔型花岗岩岩浆热液有关的钨、锡(金、钼、铅、锌、银)矿床成矿系列为主导、多系列多类型多矿种组合的金属及非金属矿床成带成片分布,如香炉山矿田、彭山矿集区、大湖塘矿集区、莲花山矿集区和鹅湖矿集区。阳储岭矿田为九岭—鄣公山隆起带内已发现的与壳幔重熔型花岗岩岩浆热液有关的钨、钼矿床。

## 3 矿床地质

矿区出露新元古界双桥山群(Pt,Sh)浅变质岩,岩性以角岩化粉砂岩、千枚岩、板岩为主;矿区发育东西向褶皱,背斜2个,向斜3个;断裂构造有东西向(F1、F2)、北东向(F3)和北北东向(F4、F5)三组控制着三角形岩体边界(图1)。阳储岭杂岩体呈北窄南宽三角形状岩株侵入于双桥山群板岩中(图1),杂岩体主要岩性及侵入顺序自早而晚为二长花岗斑岩→花岗闪长

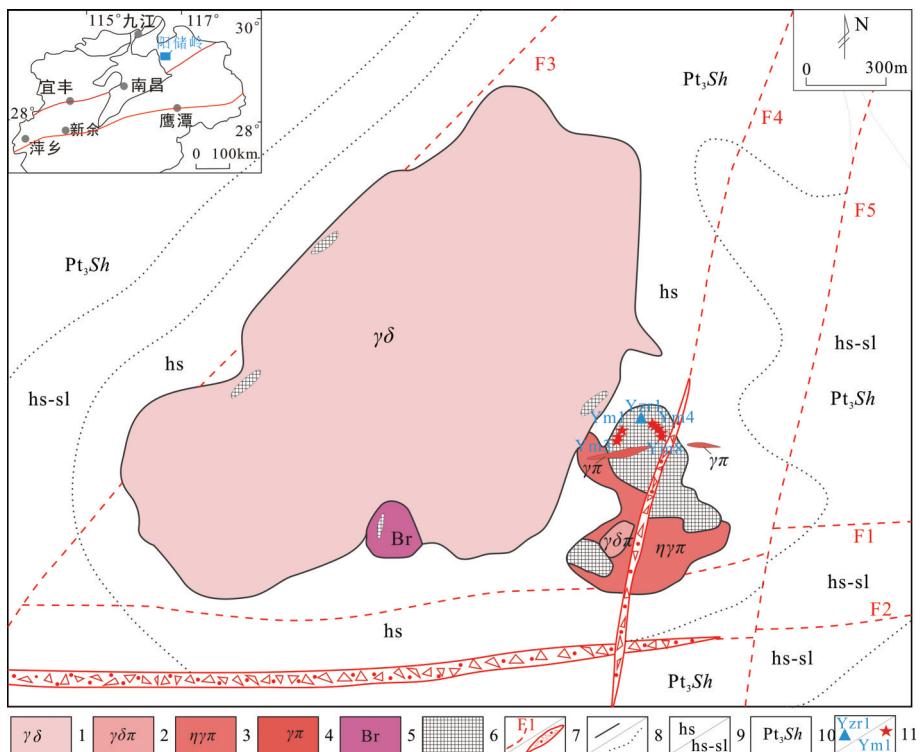


图1 阳储岭钨钼矿区地质略图(江西省地质矿产勘查开发区,2015)

1—花岗闪长岩;2—花岗闪长斑岩;3—二长花岗斑岩;4—花岗斑岩;5—爆破角砾岩;6—矿体;7—断层及编号/破碎带;8—地质界线/蚀变界线;9—角岩/角岩化板岩;10—新元古界双桥山群;11—花岗闪长斑岩采样位置及样品编号/辉钼矿采样位置及样品编号

Fig. 1 Simplified geological map of the Yangchuling tungsten-molybdenum deposit(after Jiangxi Bureau of Geological and Mineral Exploration and Development, 2015)

1—Granodiorite; 2—Granodiorite porphyry; 3—Monzogranite porphyry; 4—Granite porphyry; 5—Blasting breccia; 6—Orebody; 7—Fault and number; 8—Geological boundary/alteration boundary; 9—Hornstone/ Hornfelsed slate; 10—Sampling location and its number of granodiorite/sampling location and its number of molybdenite

斑岩→花岗斑岩→花岗闪长岩→爆破角砾岩(李秉伦等,1985);杂岩体的主体为花岗闪长岩,二长花岗斑岩和花岗闪长斑岩为成矿岩浆岩,分布于花岗闪长岩东南外接触带中(莫名演,1988)。二长花岗斑岩和花岗闪长斑岩呈渐变关系,二长花岗斑岩由边缘向内部渐变为似斑岩二长花岗斑岩,并分布于花岗闪长斑岩外接触带中。二长花岗斑岩和花岗闪长斑岩内蚀变较强,自岩体中心往往外接触带常具有钾硅化带→石英绢(白)云母化带→绿泥石碳酸盐化带→石英云母角岩带→角岩化板岩带的蚀变分带特征(江西省地质矿产勘查开发局,2015),硅化叠加于各个蚀变带中,钾长石化、硅化与成矿密切相关。

矿区钨钼矿体主要呈近水平似层状、透镜状、网脉状产出,于二长花岗斑岩、花岗闪长斑岩,少量产出于花岗闪长岩、爆破角砾岩中,具上钨(白)下铜、上富下贫的矿化特征。区内钨钼矿体均主要呈细脉、浸

染状产出,于二长花岗斑岩、花岗闪长斑岩内有零星小矿体产出。二长花岗斑岩、花岗闪长斑岩体中可圈定出Ⅰ号、Ⅱ号2个矿带,矿带总体走向均为NW、倾向南东、倾角5~15°。Ⅰ号矿带为主矿带,走向延伸200~500m、宽100~300m,层厚数米至数十米、最厚为139m,WO<sub>3</sub>品位多介于0.15%~0.20%、平均0.19%,Mo品位多介于0.05%~0.06%、平均0.06%,伴生有Cu、Pb、Zn、Be、Au、Ag等有益元素。矿石中的金属矿物主要有白钨矿和辉钼矿,次为黄铁矿、磁黄铁矿,少量黑钨矿、锡石、辉铋矿、黄铜矿、方铅矿、闪锌矿等;非金属矿物主要有钾长石、绿泥石、石英、白云母、角闪石、黑云母、方解石,偶有绿帘石、黄玉、电气石、磷灰石、萤石等。白钨矿呈他形粒状、浸染状产出,嵌布于其他矿物粒间或呈较粗粒状或团块状集合体产出。辉钼矿以浸染状为主,呈叶片状、鳞片状集合体产出。

常与白钨矿共生。矿石构造以致密块状、脉状、浸染状构造为主,次为角砾状构造;矿石结构以交代结构为主。

## 4 样品与分析测试

### 4.1 样品采集

采集阳储岭矿区含矿花岗闪长斑岩1件(用于挑选锆石定年样品)与8件辉钼矿样品,所有样品均来自露天开采断面(图1)。锆石U-Pb定年样编号为YZr1。花岗闪长斑岩呈灰白—浅灰色(图2a),斑状结构、块状构造,斑晶( $64\% \pm$ )包括斜长石( $28\% \pm$ )、石英( $17\% \pm$ )、钾长石( $15\% \pm$ )、黑云母( $4\% \pm$ ,图2b)。斜长石斑晶,半自形板状为主,粒径 $0.5\sim 2$  mm,常见聚片双晶构造(图2b,局部具绢云母化;石英斑晶,他形粒状,粒径 $0.5\sim 2$  mm;钾长石斑晶,他形粒状,粒径 $0.5\sim 1.5$  mm,常见高岭石化;黑云母斑晶,褐色,多色性一般为浅黄褐—黄褐色,片径一般为 $0.5\sim 2$  mm,解理缝或边缘可见绿泥石化。基质,显微粒状结构,由长英质及少量黑云母构成,粒径一般为 $0.05\sim 2$  mm,长石类可见绢云母化、高岭石化。用于分析测试辉钼矿Re-Os同位素样品编号为Ym1~Ym8(图1)。样品中辉钼矿呈细粒状或鳞片状,以辉钼矿+石英细脉形式产出在花岗闪长斑岩中(图2a,c,d)。

### 4.2 分析测试

锆石U-Pb定年样品破碎后采用浮选和磁选方法分选出锆石,双目镜下挑选晶形、色泽好,透明度高的锆石颗粒。挑选出的锆石与标准锆石一起用环氧树脂充分固定、制靶,后磨制抛光。锆石样品经反射光和透射光照射后,在阴极发光(CL)图像观察锆石内部结构,选择环带清晰的岩浆岩锆石作为

待测点。LA-ICP-MS锆石U-Pb年龄的测试工作在中国地质科学院矿产资源研究所LA-ICP-MS实验室完成,测试仪器为Agilent 7500a型ICP-MS与激光剥蚀系统ComPex 102 Excimer。锆石年龄计算以国际标准锆石91500为外标,测试结果用ICPMS DataCal 4.3程序处理,并进行校正。锆石年龄谐和图用Isoplot 3.0程序处理,具体分析测试过程详见文献柳小明等(2002)。

辉钼矿Re-Os同位素定年样品破碎后经重力和电磁分离,在显微镜下挑选出纯度大于99%的辉钼矿单矿物用以分析测试。样品分析测试由国家地质实验测试中心Re-Os同位素实验室完成,分析仪器为电感耦合等离子体质谱仪TJA X-series ICP-MS。样品处理流程和质谱测定技术参考相关文献(杜安道等,2007;李超等,2012)。本次试验空白水平远低于所测试样品中Re、Os含量,测试结果准确可信。

## 5 测试结果

### 5.1 锆石U-Pb年代学

成矿花岗闪长斑岩内锆石无色,透明,主要呈长柱状或短柱状,可见少数呈细粒浑圆状,部分锆石内见细小的包裹体及裂纹(图3),锆石粒径为 $(40\sim 80)\mu\text{m} \times (120\sim 230)\mu\text{m}$ 。阴极发光图像显示锆石成分含量比较均匀,主要以振荡环带为主。本次22个测试点的Th、U、Pb含量范围分别为 $217\sim 792\ \mu\text{g/g}$ 、 $408\sim 2519\ \mu\text{g/g}$ 、 $13\sim 61\ \mu\text{g/g}$ , Th/U比值介于0.19~0.53(表1),显示出岩浆锆石的特征。经测定22个有效锆石测点的 $^{207}\text{Pb}/^{235}\text{U}$ 、 $^{206}\text{Pb}/^{238}\text{U}$ 分析结果在误差范围内比值基本一致,其中 $^{206}\text{Pb}/^{238}\text{U}$ 年龄的加权平均值为 $(145.08 \pm 0.35)\text{Ma}$ (MSWD=0.51, n=7)(表1,图4),代表花岗长斑岩的结晶年龄。

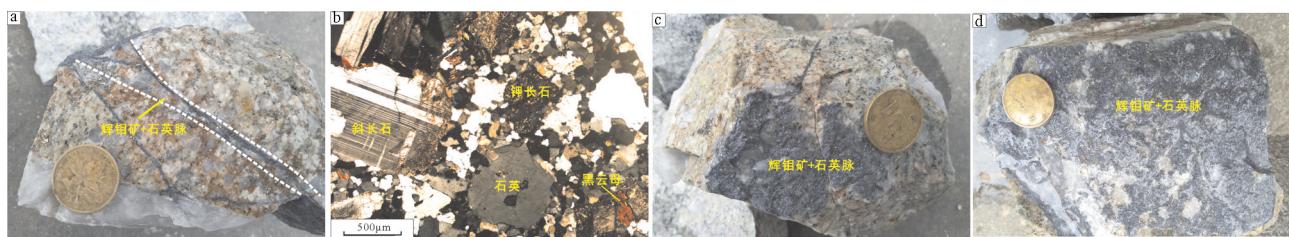


图2 阳储岭矿区锆石U-Pb与辉钼矿Re-Os样品照片

a—含矿花岗闪长斑岩照片;b—花岗闪长斑岩镜下照片;c—花岗闪长斑岩内脉状辉钼矿;d—花岗闪长斑岩外接触带角岩内辉钼矿脉

Fig. 2 Photographs of zircon U-Pb and molybdenite Re-Os samples in Yangchuling deposit

a—Photographs of ore-bearing granodiorite porphyry; b—Microscopic photos of granodiorite porphyry; c—Vein molybdenite in granodiorite porphyry; d—Molybdenite veins in hornfels at the outer contact zone of granodiorite porphyry

## 5.2 辉钼矿 Re-Os 定年

矿区 8 件辉钼矿样品的 Re-Os 同位素测试结果见表 2 和图 5。辉钼矿 Re-Os 模式年龄( $143.3 \pm 2.0$ ) Ma~( $145.5 \pm 2.2$ ) Ma, 加权平均年龄为( $144.89 \pm 0.7$ ) Ma, MSWD=0.49(图 5), 与单个样品的模式年龄相差 0~1.69 Ma。采用 ISOPLOT 软件对 8 件辉钼矿 Re-Os 数据进行等时线拟合, 获得其 Re-Os 等时线年龄为( $145.4 \pm 1.0$ ) Ma, MSWD=0.84(图 5)。

## 6 讨论

### 6.1 成岩成矿时代

矿区成矿花岗闪长斑岩体的锆石 U-Pb 年龄( $145.08 \pm 0.35$ ) Ma(MSWD=0.51,  $n=7$ )与辉钼矿 Re-Os 模式年龄( $143.3 \pm 2.0$ ) Ma~( $145.5 \pm 2.2$ ) Ma、等时线年龄( $145.4 \pm 1.0$ ) Ma 基本一致, 表明阳储岭钨钼矿床的成矿花岗闪长斑岩的结晶年龄为( $145.08 \pm 0.35$ ) Ma, 成矿年龄为( $145.4 \pm 1.0$ ) Ma。

### 6.2 成矿物质来源

辉钼矿中 Re-Os 同位素体系对于成矿物质来源的示踪作用研究由来已久, Stein et al. (2001)指出变质流体参与成矿的矿床中辉钼矿的 Re 含量低于  $20 \times 10^{-6}$ 。Mao et al. (1999)通过对国内多个矿床辉

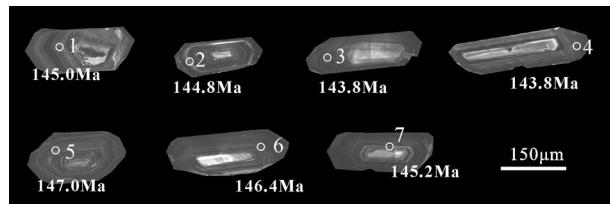


图3 阳储岭矿区黑云母花岗岩锆石阴极发光及测点位置图  
Fig.3 Zircon cathodoluminescence photos of biotite granite in Yangchuling mining area and the position of measuring point

钼矿 Re 含量与成矿物质来源的关系研究提出, 从幔源→壳幔混源→壳源辉钼矿中 Re 的含量具有规律性的变化, 并总结为  $n \times 10^{-4} \rightarrow n \times 10^{-5} \rightarrow n \times 10^{-6}$  逐渐降低的变化规律(Mao et al., 1999)。本文测得的阳储岭钨钼矿 8 件辉钼矿中 Re 含量为  $16.62 \times 10^{-6}$ ~ $87.76 \times 10^{-6}$ , 平均值为  $44.68 \times 10^{-6}$ , 与壳幔混源岩浆热液矿床中 Re 的含量相似, 指示阳储岭钨钼矿床成矿物质应来源于壳幔混源。阳储岭钨钼矿含矿斑岩  $^{87}\text{Sr}/^{86}\text{Sr}$  初始值为 0.70769~0.71001(李秉伦等, 1985)、0.7108~0.7120(满发胜等, 1988)、0.70862(莫名演, 1990), 表明杂岩体具壳幔混源同熔型岩浆岩特征。阳储岭矿区金属硫化物  $\delta^{34}\text{S}$  变化范围为 +1.3‰~+4.01‰, 平均 +3.21‰, 硫同位素组成呈塔形, 表明硫与岩浆岩同源(迟实福等, 1985)。本次

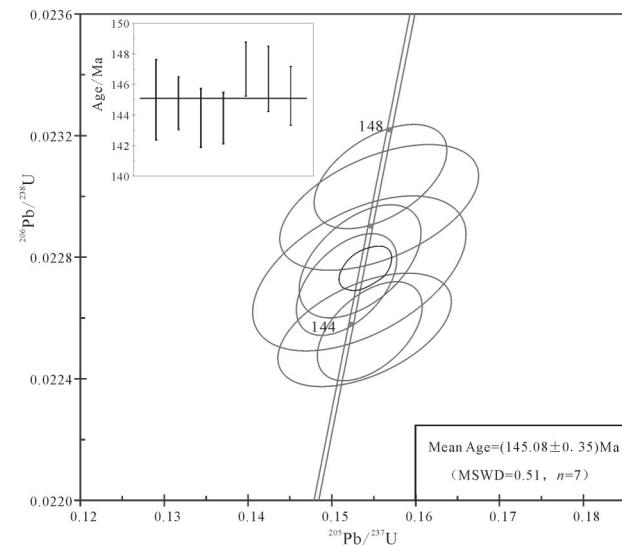


图4 阳储岭黑云母花岗岩 LA-ICP-MS 锆石 U-Pb 年龄谐和图  
Fig.4 LA-ICP-MS zircon U-Pb Age concordia diagram of biotite granite in Yangchuling mining area

表1 阳储岭矿区 LA-ICP-MS 锆石 U-Pb 同位素测年  
Table 1 LA-ICP-MS zircon U-Pb Isotope chronology of Yangchuling mining area

测点 编号	含量/(μg/g)			同位素比值						同位素年龄/Ma						
	Pb	Th	U	Th/U	$^{207}\text{Pb}/^{206}\text{Pb}$	$1\sigma$	$^{207}\text{Pb}/^{235}\text{U}$	$1\sigma$	$^{206}\text{Pb}/^{238}\text{U}$	$1\sigma$	$^{207}\text{Pb}/^{206}\text{Pb}$	$1\sigma$	$^{207}\text{Pb}/^{235}\text{U}$	$1\sigma$	$^{206}\text{Pb}/^{238}\text{U}$	$1\sigma$
1	61	792	2519	0.31	0.0475	0.0025	0.1533	0.0104	0.0227	0.0001	76.0	118.5	144.8	9.2	145.0	1.3
2	27	247	1285	0.19	0.0488	0.0012	0.1518	0.0049	0.0227	0.0001	200.1	57.4	143.5	4.3	144.8	0.9
3	28	301	1337	0.22	0.0489	0.0021	0.1539	0.0085	0.0226	0.0002	142.7	100.0	145.4	7.4	143.8	1.0
4	43	445	1997	0.22	0.0497	0.0013	0.1545	0.0051	0.0226	0.0001	189.0	59.2	145.9	4.5	143.8	0.8
5	13	217	408	0.53	0.0489	0.0017	0.1558	0.0065	0.0231	0.0001	142.7	88.0	147.0	5.7	147.0	0.9
6	29	353	1476	0.24	0.0478	0.0023	0.1557	0.0097	0.023	0.0002	87.1	111.1	146.9	8.5	146.4	1.1
7	46	619	2265	0.27	0.0481	0.0015	0.1534	0.0059	0.0228	0.0002	105.6	72.2	144.9	5.2	145.2	1.0

表2 阳储岭钨钼矿床辉钼矿Re-Os同位素测试成果表

Table 2 Re-Os isotope results of molybdenite in Yangchuling tungsten-molybdenum deposit

样号	样重/g	Re/(ng/g)		普 Os/(ng/g)		$^{187}\text{Re}/(\mu\text{g/g})$		$^{187}\text{Os}/(\mu\text{g/g})$		模式年龄/Ma	
		测定值	$2\sigma$	测定值	$2\sigma$	测定值	$2\sigma$	测定值	$2\sigma$	测定值	$2\sigma$
Ym1	0.01007	18592	153	0.0091	0.1853	11686	96	28.18	0.2	144.6	2.1
Ym2	0.0105	23747	181	0.0048	0.0122	14926	114	36.22	0.29	145.5	2.2
Ym3	0.05318	87760	574	0.0173	0.0099	55159	361	133.8	0.9	145.4	2
Ym4	0.01022	16618	113	0.0545	0.0407	10445	71	25.27	0.18	145	2
Ym5	0.01124	20067	159	0.0009	0.0743	12613	100	30.15	0.2	143.3	2
Ym6	0.01049	64778	422	0.0028	0.0648	40715	265	98.76	0.68	145.4	2
Ym7	0.01008	66792	554	0.2601	0.3157	41980	348	101.4	0.7	144.9	2.1
Ym8	0.01021	59089	405	0.3145	0.0385	37139	255	89.83	0.6	145	2

注:1. 普 Os 是根据原子量表和同位素丰度表,通过  $^{192}\text{Os}/^{190}\text{Os}$  测量比计算得出,  $^{187}\text{Os}$  是  $^{187}\text{Os}$  同位素总量;2. Re、Os 含量的不确定度包括样品和稀释剂的称量误差、稀释剂的标定误差、质谱测量的分馏校正误差、待分析样品同位素比值测量误差,置信水平 95%;3. 因为辉钼矿铼含量较高,几乎不含非放射成因的  $^{187}\text{Os}$ ,故用样品的铼、锇含量按照下列公式直接计算模式年龄( $t$ ): $t=1/\lambda[\ln(1+^{187}\text{Os}/^{187}\text{Re})]$ ,其中  $\lambda$ ( $^{187}\text{Re}$  衰变常数)= $1.666 \times 10^{-11} \text{ a}^{-1}$ ,模式年龄的不确定度还包括衰变常数的不确定度(1.02%),置信水平 95%;4. 在计算模式年龄或作  $^{187}\text{Re}-^{187}\text{Os}$  等时线时  $^{187}\text{Re}$  和  $^{187}\text{Os}$  的单位应该是质量摩尔浓度,即 mol/g,为了直观,实际上采用了质量分数,即 ng/g,这是因为  $^{187}\text{Re}$  相对原子质量 186.955765 和  $^{187}\text{Os}$  相对原子质量 186.955762 非常接近,无论采用什么单位得到的模式年龄或等时线年龄的差别都将小于千万分之一,远远小于目前年龄测定的不确定度范围 2%;5. 测试单位:国家地质实验测试中心 Re-Os 同位素实验室;6. 测试人:李超。

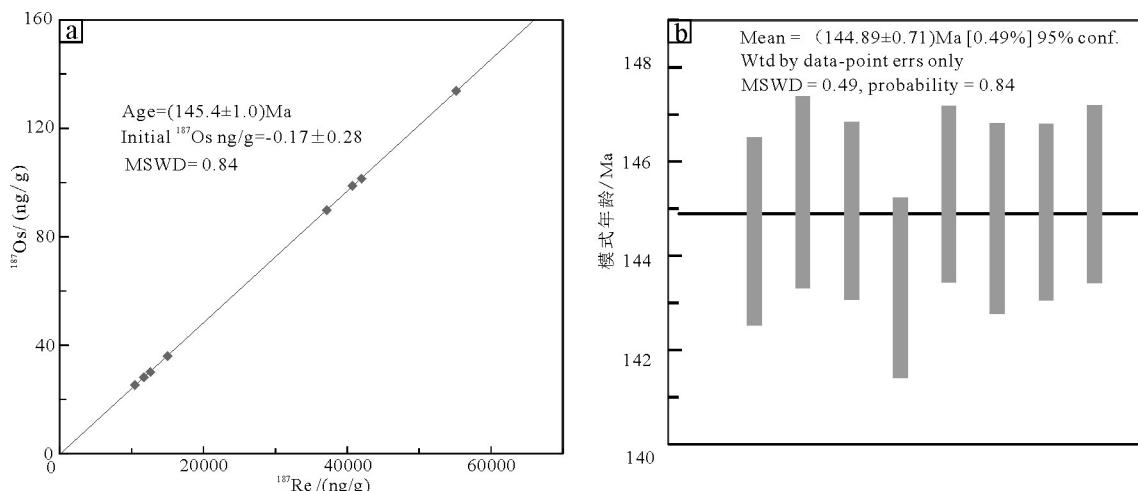


图5 阳储岭钨钼矿Re-Os加权平均年龄和等时线年龄  
(a—等时线年龄; b—加权平均年龄)

Fig.5 Re-Os weighted average age and isochron age of Yangchuling tungsten-molybdenum deposit  
(a—Isochrone age; b—Weighted average age)

通过辉钼矿中 Re-Os 同位素体系中 Re 的含量对成矿物质来源的指示,与前人含矿斑岩体全岩 Sr 同位素,金属硫化物中 S 同位素的分析结论一致(李秉伦等,1985; 满发胜等,1988; 莫名演,1990),进一步表明了阳储岭矿区成矿物质来自壳幔混源。

### 6.3 赣北地区钨成矿作用

赣北地区主要经历了晋宁、加里东、燕山 3 期次强烈的造山运动,出现了一系列重大成矿事件,每

个阶段内都有强烈的构造运动和成岩成矿作用,并形成相应类型的矿床,其中与钨成矿作用密切相关的为燕山期(胡正华等,2018a、2018b、2018c; Hu et al., 2018)。在早中侏罗世江西已有较多成矿期前花岗岩分布,但早中侏罗世成矿局限于雩山成矿带南部的盘古山一大吉山一线(181~178 Ma),属成矿初始期;大规模成矿时期为晚侏罗世—早白垩世(165~95 Ma),属“成矿大爆炸”期。赣北地区塔前

铜钨矿床主成矿岩体锆石U-Pb年龄值为(160.9±2.5) Ma(胡正华等,2015a)、辉钼矿等时线年龄为(162±2) Ma(黄安杰等,2013)。朱溪钨铜矿成矿黑云母花岗岩锆石U-Pb年龄(146.9±0.97)Ma~(152.9±1.7) Ma(王先广等,2015;胡正华,2015b);辉钼矿Re-Os与白钨矿Sm-Nd同位素年龄为(144±5) Ma~(153±3) Ma(刘善宝等,2017)。石门寺、大湖塘钨矿成矿似斑状白云母花岗岩锆石U-Pb年龄141~144 Ma(黄兰椿等,2012),辉钼矿Re-Os等时线年龄141~144 Ma(丰成友等,2012;Mao et al., 2013)。香炉山成矿二云母花岗岩锆石U-Pb年龄(123±1) Ma(胡正华等,2018c),白钨矿Sm-Nd等时线年龄为121 Ma、共生石英Rb-Sr等时线年龄128 Ma(张家菁等,2008),成岩成矿时代基本一致为(125±3) Ma。东坪成矿黑云母花岗岩锆石U-Pb年龄为(128.79±0.39) Ma~(132.9±1.4) Ma(胡正华等,2018a)。茅棚店、八字脑等成矿黑云母花岗岩锆石年龄120.5~125.0 Ma(胡正华,2018c)。通江岭铜钨矿成矿花岗闪长斑岩锆石U-Pb年龄为(143.31±2.70) Ma(胡正华,2018b),塔前、朱溪、石门寺、大湖塘、通江岭以及东坪、香炉山、阳储岭钨钼矿床成岩成矿时代指示出赣北地区存在(161±2) Ma、150~145 Ma、135~125 Ma等3期钨成矿事件,并且自南至北成矿时代逐渐年轻。赣南地区与钨成矿作用相关的岩浆岩均为S型花岗岩,而赣北地区则I、S型花岗岩均有。赣北地区朱溪、大湖塘、石门寺、东坪、香炉山等钨钼矿床的成矿岩浆岩均为S型花岗岩,而阳储岭成矿岩浆岩与九瑞矿集的通江岭、武山等矿床类似,均为I型花岗岩,形成时代均为150~145 Ma(李秉伦等,1985;满发胜等,1988;莫名演,1990;胡正华,2015c,2018c),进一步说明150~145 Ma为赣北地区钨多金属矿的主要成矿期。

## 7 结 论

(1)阳储岭钨钼矿床成矿花岗闪长斑LA-ICP-MS锆石U-Pb年龄为(145.08±0.35)Ma(MSWD=0.51,n=7),代表岩体结晶年龄;辉钼矿Re-Os模式年龄为(143.3±2.0)Ma~(145.5±2.2)Ma,等时线年龄为(145.4±1.0) Ma。阳储岭成矿花岗闪长斑岩锆石U-Pb年龄与辉钼矿Re-Os等时线年龄基本一致,表明阳储岭钨钼矿床成矿岩成矿时代约为145 Ma。

(2)辉钼矿中Re含量为 $16.62 \times 10^{-6}$ ~ $87.76 \times 10^{-6}$ ,平均值为 $44.68 \times 10^{-6}$ ,与壳幔混源岩浆热液矿床中Re的含量相似,指示阳储岭钨钼矿床成矿物质应来源于壳幔混源。

(3)赣北地区存在3期钨成矿事件:(161±2) Ma、150~145 Ma、135~125 Ma,并且自南至北成矿时代逐渐年轻。

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