

# 中天山北缘华力西期造山作用 ——变质岩锆石 U-Pb 年代学限定

林彦蒿<sup>1</sup> 张泽明<sup>2</sup> 贺振宇<sup>2</sup> 董 昕<sup>2</sup> 于 飞<sup>1</sup>

(1. 中国地质大学(武汉)地球科学学院, 湖北 武汉 430074; 2. 中国地质科学院地质研究所, 北京 100037)

**摘要:**笔者对中天山微陆块北缘托克逊干沟地区角闪岩相变质岩中的锆石进行了 U-Pb 年代学研究, 结果证明变质沉积岩中的碎屑锆石记录了从太古宙至元古宙(3320~530 Ma)的源区岩浆热事件, 变质火成岩中的岩浆锆石记录了新元古代晚期(550 Ma)的岩浆作用, 而变质锆石记录了晚泥盆纪(385~360 Ma)的变质作用。这一定年结果表明, 中天山微陆块北缘的造山作用很可能发生在华力西期, 中天山微陆块形成于新元古代以前, 但并没有经历前寒武纪变质作用, 具有与塔里木克拉通明显不同的前寒武纪构造演化历史。因此, 中天山微陆块很可能是一个独立的块体, 并不支持其是从塔里木板块分离出来的观点。

**关键词:**中亚造山带; 中天山; 变质作用; 锆石 U-Pb 定年; 华力西期造山作用

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东西向绵延达 2500 km 的天山山脉夹持于准噶尔盆地与塔里木盆地之间, 是中亚地区最重要的造山带之一。中天山地块(或称伊犁—中天山地块)位于天山山脉的核部, 被认为是一个微陆块<sup>[1]</sup>, 其西宽东窄, 在库米什以东地区尖灭(图 1)。基于岩石学研究, 特别是蛇绿岩、高压-超高压变质岩以及沉积相分析, 研究人员推断沿中天山南、北两边存在两个古生代的俯冲-碰撞带, 正是这两次碰撞作用形成了天山造山带的基本构造格局<sup>[2-6]</sup>。同时, 中、新生代的构造运动导致中天山微陆块边界的再活化, 使其构造变得更加复杂。

天山造山带是巨型中亚造山带的一部分, 记录了从新元古代到新生代的长期增生造山历史。关于天山造山带的古生代构造演化已经取得了重要的研究进展<sup>[1,4,8-18]</sup>。但同时也存在许多争议, 如对于天山南部蛇绿岩的形成时代就有从晚泥盆到早石炭世、

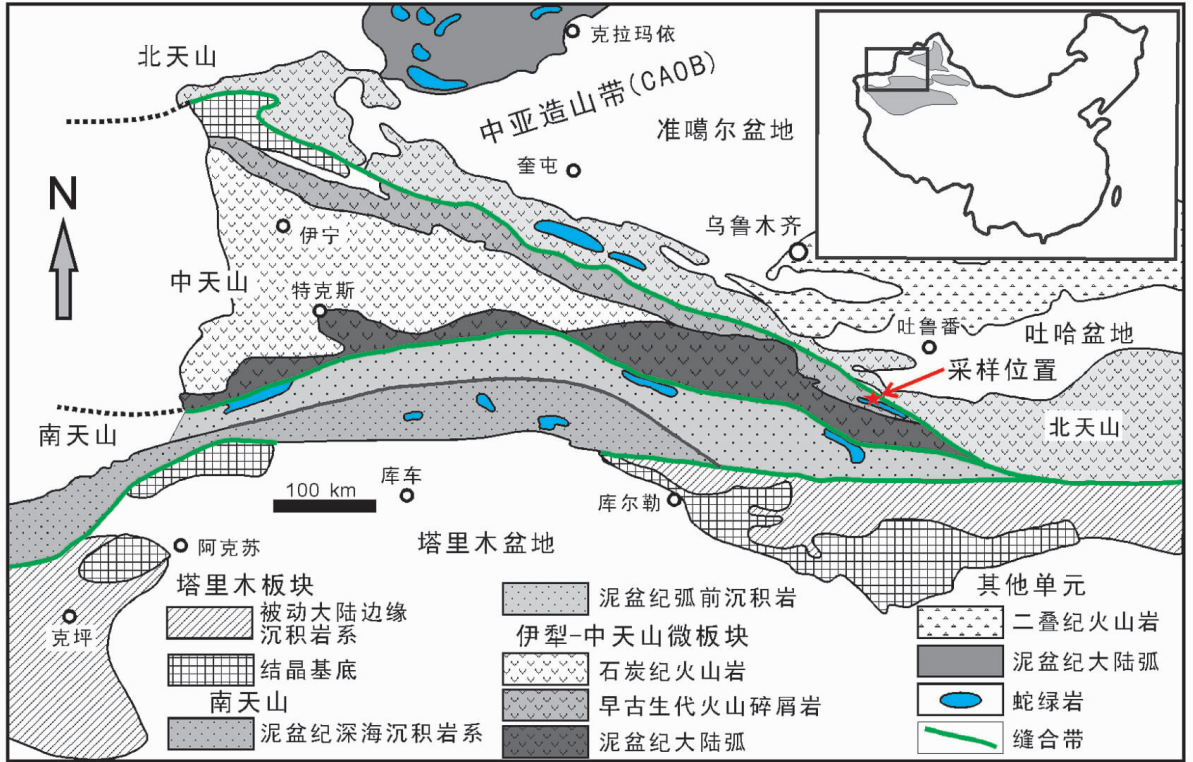
晚泥盆到早二叠世和晚二叠世等不同认识<sup>[8,15-16]</sup>。关于塔里木板块与中天山微陆块的碰撞时代也有不同认识, 如发生在早三叠世(~250 Ma)<sup>[19]</sup>或早二叠世<sup>[20]</sup>(~290 Ma)。王润三等<sup>[21]</sup>和 Shu et al.<sup>[22]</sup>认为库米什榆树沟蛇绿岩部分地叠加了高压麻粒岩相变质作用。周鼎武等<sup>[23]</sup>对榆树沟高压麻粒岩中锆石的 SHRIMP U-Pb 定年结果表明, 其变质作用峰期时代为 390 Ma。杨天南等<sup>[24-25]</sup>研究证明库米什地区的陆弧型花岗岩形成时代为 396 Ma, 并认为是由于榆树沟蛇绿岩套所代表的洋壳向北俯冲过程形成的, 地幔局部熔融及玄武岩浆上涌带来的热量引起中、下地壳的局部熔融, 分别形成玄武质侵入岩及重熔型花岗岩, 并使部分岩石发生麻粒岩相变质作用。这些研究成果为中天山微陆块北缘俯冲-碰撞作用发生在早泥盆世提供了重要的限定。但是, 到目前为止, 中天山微陆块南缘俯冲-碰撞带的形成时间还缺少必要的

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作者简介: 林彦蒿, 男, 1987 年生, 本科生, 岩石学专业; E-mail: linyanhao1987@sina.com。

通讯作者: 张泽明, 男, 1961 年生, 研究员, 岩石学与变质地质学专业; E-mail: zmzhang@cags.ac.cn。

图 1 天山造山带构造简图(据 Yang et al., 2007<sup>[7]</sup>)Fig.1 Sketch geological map of the Tianshan Mountains (after Yang et al., 2007<sup>[7]</sup>)

变质作用时代限定。因此,本文选取中天山微陆块南部碰撞带附近的变质岩进行了锆石 U-Pb 年代学研究,获得了泥盆纪的变质年龄,为中天山经历华力西期造山作用提供了重要的制约。

## 1 样品及岩相学特征

研究样品采自托克逊县城南约 30 km 的干沟(阿格布拉克沟)(图 1)。在 1:50 万乌鲁木齐幅区域地质调查报告中,研究区被划分为新元古界,岩性为结晶片岩、片麻岩、石英岩和大理岩。但在 1:20 万库米什幅区域地质调查报告中,这里又被划归为志留纪阿哈布拉克岩群,岩性包括黑云母石英片岩和大理岩。沿着中天山微陆块北缘分布有许多大小不等的超基性岩体,它们被认为是蛇绿岩套的一部分,其中研究区的超基性岩被称为米什沟蛇绿岩。

笔者所研究的样品为米什沟超基性岩体的围岩,经历了中压角闪岩变质作用,主要岩性是片麻岩和石榴石云母片岩。其中样品 07XT04-1 为十字石石榴石黑云母片岩,共生矿物是石榴石+十字石+黑云母+石

英+斜长石。另一个样品 07XT06-1 为白云斜长片麻岩,共生矿物为斜长石+钾长石+白云母+石英。

## 2 锆石 U-Pb 定年结果

锆石 U-Pb 同位素定年和微量元素分析在中国地质大学(武汉)地质过程与矿产资源国家重点实验室(GPMR)利用 LA-ICP-MS 分析完成。激光剥蚀系统为 GeoLas 2005, ICP-MS 为 Agilent 7500 a。激光剥蚀过程中采用氦气作载气、氩气为补偿气以调节灵敏度,二者在进入 ICP 之前通过一个 T 型接头混合。在等离子体中心气流(Ar+He)中加入了少量氮气,以提高仪器灵敏度、降低检出限和改善分析精度。每个时间分辨分析数据包括 20~30 s 的空白信号和 50 s 的样品信号。对分析数据的离线处理(包括对样品和空白信号的选择、仪器灵敏度漂移校正、元素含量及 U-Th-Pb 同位素比值和年龄计算)采用软件 ICPMSDataCal。详细的仪器操作条件和数据处理方法同 Liu et al.<sup>[26-28]</sup>。

锆石微量元素含量利用多个 USGS 参考玻璃

(BCR-2G, BIR-1G)作为多外标、Si作内标的方法进行定量计算。U-Pb同位素定年中采用锆石标准91500作外标进行同位素分馏校正,每分析5个样品点,分析2次91500。对于与分析时间有关的U-Th-Pb同位素比值漂移,利用91500的变化采用线性内插的方式进行了校正<sup>[27]</sup>。锆石样品的U-Pb年龄谐和图绘制和年龄权重平均计算均采用Isoplot/Ex\_ver3完成。

十字石榴石黑云母片岩(07XT04-1)中的锆石呈半自形短柱状,部分为他形椭圆状。阴极发光图像显示,大部分锆石具有核-边结构,其核部发育有明显的同心韵律状环带,而较窄的边部不具环带。白云斜长片麻岩(07XT06-1)中的锆石呈半自形短柱

状或他形椭圆状。阴极发光分析揭示,锆石具有核-边结构,其核部具有同心韵律状环带,边缘不具环带。已有研究表明,具有同心韵律状环带的锆石是岩浆结晶形成的,而不具环带的锆石是在变质作用中生长的<sup>[29-35]</sup>。因此,所研究的片岩和片麻岩中的锆石核部是早期岩浆结晶形成的,而边部是后来变质成因的。

LA-ICP-MS U-Pb原位分析结果表明,十字石榴石黑云母片岩(07XT04-1)中锆石岩浆核的11分析点给出了类似的<sup>206</sup>Pb/<sup>238</sup>U年龄,在517~588 Ma(表1和图2-a),所获得的峰值年龄在~550 Ma(图2-b)。一个锆石变质边分析点给出了357 Ma的近谐和年龄(表1,图2-a)。分析结果也表明,锆石

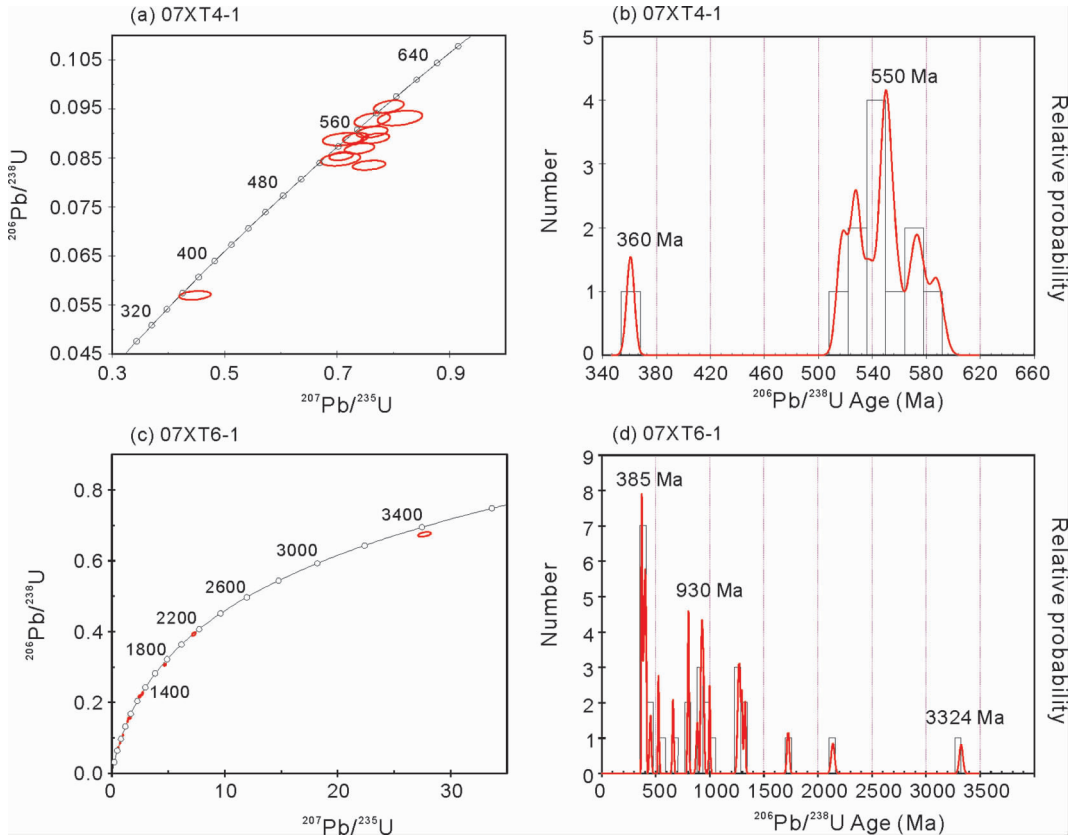


图2 干沟变质岩锆石U-Pb年龄谐和图和频率图

(a)和(b)为十字石榴石黑云母片岩(07XT4-1);(c)和(d)为白云母片麻岩(07XT6-1)

Fig.2 Zircon U-Pb age concordia and probability diagrams of Gangou metamorphic rocks (a) and (b) are staurolite-garnet-biotite schist (07XT4-1); (c) and (d) are muscovite-schist (07XT6-1)





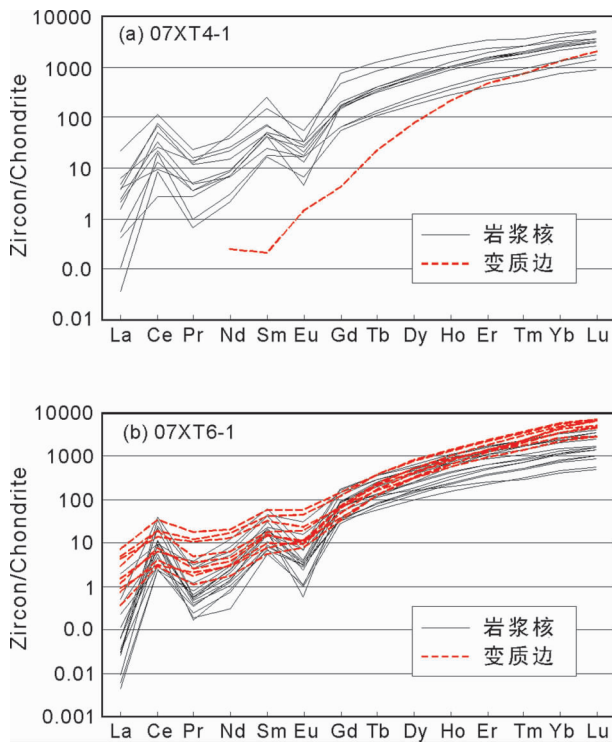


图3 球粒陨石标准化的锆石 REE 型式

Fig.3 Chondrite-normalized REE patterns of zircons

岩浆核具较高的稀土元素 (REE) 含量, 并具有 LREE 亏损, HREE 富集和负 Eu 异常的 REE 配分型式(图 3-a)。同时, 核部分析点也具有较高的 Th/U 比值(0.305~1.208)。这些成分特征也表明, 锆石核部是岩浆结晶形成的。相反, 锆石的边缘具有较低的 REE 含量, 特别是非常较低的 LREE 含量, 也不具有 Eu 异常, Th/U 比值也低到 0.003(图 3, 表 1), 是变质锆石的典型特征<sup>[29-30,32-35]</sup>。因此, 笔者认为锆石核部获得的~550 Ma 年龄代表变质岩原岩的结晶年龄, 而~360 Ma 很可能代表角闪岩相变质年龄。

白云斜长片麻岩(07XT06-1)中锆石核部的 18 个分析点给出了分散的  $^{206}\text{Pb}/^{238}\text{U}$  年龄, 从 3324~527 Ma, 而且不具有明显的年龄峰值(图 2-c, d, 表 1)。相反, 9 个锆石变质边分析点给出了相近的  $^{206}\text{Pb}/^{238}\text{U}$  年龄, 为 372~451 Ma, 其峰值年龄在 385 Ma(图 2-c, d, 表 1)。分析结果也显示, 锆石岩浆核部与变质边部的 REE 含量类似, 部分变质边的 REE 含量甚至更高(图 3-b 和表 2), 这很可能是由于角闪岩相变质过程中没有形成富 REE 矿物(如石榴石), 导

致岩石中的 REE 富集到了变质锆石之中。但是, 岩浆锆石核具有明显的负 Eu 异常, 而变质锆石的 Eu 异常则不明显(图 3-b, 表 2)。而且, 锆石岩浆核具有较高的 Th/U 比值(0.09~1.015), 变质边具有较低的 Th/U 比值(0.003~0.033)。这些成分特征表明, 锆石的核部是继承的碎屑岩浆锆石, 而边部是变质成因的。因此, 可以认为白云母片麻岩的原岩为沉积岩, 其物质源区记录了从太古宙至元古宙的岩浆热事件, 而其角闪岩相变质作用发生在~385 Ma。

基于上述定年结果, 笔者认为中天山微陆块北缘经历了新元古代晚期的岩浆作用和古生代华力西期的变质作用。

### 3 构造意义

作为中亚造山带的重要组成部分, 天山造山带是研究不同陆块碰撞拼合、大陆地壳生长和成矿作用的天然实验室<sup>[7-8,15-16,36-37]</sup>。但正如前面谈到的, 到目前为止, 有关中天山北缘俯冲-碰撞带的形成时代还存在争议。研究区的蛇绿岩是天山洋的一部分, 冯益民等<sup>[38]</sup>认为该洋盆闭合的时间发生在 400 Ma 之前。朱宝清等<sup>[39]</sup>对干沟一带前陆盆地碎屑岩中的陆源锆石(~460 Ma)及侵入该碎屑岩的花岗闪长岩中的岩浆锆石(~390 Ma)进行的研究表明, 早古生代中天山北缘洋盆俯冲消减开始于奥陶纪, 结束在 390 Ma 之前。石玉若等<sup>[40]</sup>在研究区获得了同碰撞花岗岩的形成年龄在 430 Ma。但也有研究者认为, 该洋盆在晚泥盆世到早石炭世仍然存在<sup>[15,41]</sup>, 中天山与准噶尔板块的碰撞发生在石炭纪到二叠纪<sup>[7-8,41-43]</sup>或在奥陶纪<sup>[44]</sup>。但是, 本文所获得的 360~385 Ma 的变质年龄很可能说明, 中天山与准噶尔板块的碰撞发生在晚泥盆纪, 中天山微陆块北缘经历了华力西期造山作用。该变质年龄与榆树沟麻粒岩的变质年龄(390 Ma)<sup>[23]</sup>基本相同, 这很可能说明中天山微陆块南、北边缘经历了同时代的造山作用。

本研究表明, 中天山微陆块经历了新元古代晚期的岩浆事件, 因此中天山陆块的基底形成时代可能更早。但是, 所获的继承锆石年龄并不能证明中天山微陆块经历了前寒武纪变质作用。现有的研究表明, 塔里木克拉通经历了多期前寒武纪构造热事件, 如 2.5 Ga 的岩浆事件, 1.9 Ga 的变质事件和 0.9 Ga 的岩浆与变质事件等<sup>[45-48]</sup>。因此, 笔者的初步研究结果并不支持中天山微陆块是从塔里木板块分离出来

表 2 锆石的稀土元素含量 ( $10^{-6}$ )

Table 2 Rare earth element contents of zircons

样品	位置	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Tm	Yb	Lu
07XT04-1-04	变质边		0.040		0.148	0.041	0.109	1.26	1.06	24.3	15.3	96.1	24.4	266	64.6
07XT04-1-16	岩浆核	1.95	21.4	1.69	13.5	13.7	2.05	48.6	15.4	173	61.6	260	50.7	445	84.4
07XT04-1-13	岩浆核	0.734	26.7	0.597	4.09	4.90	1.25	19.2	6.44	79.6	30.4	137	29.5	276	55.5
07XT04-1-10	岩浆核	6.89	90.7	2.86	23.8	30.2	3.93	141	40.1	411	131	493	84.6	673	118
07XT04-1-15	岩浆核	0.132	2.16	0.330	4.31	8.73	0.340	50.7	18.3	230	90.5	407	85.2	798	158
07XT04-1-07	岩浆核	0.011	6.94	0.082	1.31	3.28	0.508	16.3	5.11	58.0	20.3	85.4	17.0	151	28.5
07XT04-1-02	岩浆核	0.672	17.6	0.440	4.39	8.16	1.88	48.0	16.1	199	74.7	317	60.6	524	97.7
07XT04-1-03	岩浆核	0.478	42.1	1.54	27.2	48.3	2.48	215	59.3	596	188	697	120	934	162
07XT04-1-19	岩浆核	0.175	10.9	0.435	5.02	10.1	2.38	58.8	18.5	212	76.0	314	59.8	520	96.9
07XT04-1-06	岩浆核	1.28	7.76	0.623	5.50	9.42	0.980	49.7	17.1	206	74.4	321	62.8	556	105
07XT04-1-01	岩浆核	0.034	16.1	0.121	1.82	3.37	1.22	18.6	5.77	70.2	25.5	115	23.5	220	43.9
07XT04-1-14	岩浆核	1.19	63.7	1.95	16.0	14.3	1.50	52.3	16.7	192	69.2	291	57.8	507	96.0
07XT06-1-07	变质边	1.59	15.9	1.47	10.2	8.34	3.27	36.3	18.1	268	104	494	120	1239	237
07XT06-1-24	变质边	2.22	28.2	2.17	12.3	11.6	4.22	42.6	18.7	255	96.6	463	112	1176	228
07XT06-1-16	变质边	0.112	2.37	0.141	1.09	1.12	0.564	9.82	5.96	107	47.2	254	66.3	720	143
07XT06-1-11	变质边	0.229	3.35	0.222	1.88	2.04	0.721	14.7	8.36	137	57.0	298	79.6	893	172
07XT06-1-22	变质边	1.33	11.2	1.36	7.82	6.34	1.71	19.2	7.71	104	40.9	200	47.1	473	92.2
07XT06-1-04	变质边	0.490	5.21	0.445	2.73	3.03	0.766	18.7	10.9	169	63.6	280	60.6	557	96.1
07XT06-1-19	变质边	0.280	2.58	0.269	1.74	1.55	0.869	11.0	6.04	100	46.5	267	78.0	959	211
07XT06-1-05	变质边	0.897	15.4	0.619	3.90	3.61	1.42	20.8	11.6	178	74.8	383	97.4	1049	213
07XT06-1-01	变质边	0.374	6.53	0.321	2.30	2.89	0.830	19.5	10.1	146	61.8	303	73.1	760	153
07XT06-1-02	岩浆核	0.009	29.0	0.070	1.01	2.14	0.269	9.93	2.85	32.2	11.1	49.8	10.4	97.5	19.0
07XT06-1-14	岩浆核	0.113	5.36	0.458	7.23	11.9	0.650	54.1	16.9	188	68.1	284	56.6	508	92.6
07XT06-1-23	岩浆核	0.612	7.88	0.390	3.95	8.33	0.523	50.4	17.6	215	80.9	357	74.2	677	129
07XT06-1-26	岩浆核	0.000	6.96	0.063	1.26	3.40	0.077	19.7	6.58	83.2	30.9	136	27.7	248	46.7
07XT06-1-18	岩浆核	0.001	2.01	0.099	1.80	4.45	0.222	26.8	10.3	132	54.9	262	56.7	548	112
07XT06-1-20	岩浆核	0.003	7.68	0.020	0.512	1.47	0.071	8.54	3.30	41.7	16.6	77.3	16.6	156	29.9
07XT06-1-08	岩浆核	0.021	3.67	0.045	0.772	1.59	0.229	9.49	4.04	50.3	18.1	79.5	17.9	173	33.3
07XT06-1-10	岩浆核	0.009	2.34	0.079	1.43	4.41	0.181	22.6	6.16	53.5	14.6	53.7	9.87	84.9	15.9
07XT06-1-17	岩浆核	0.021	8.96	0.051	0.610	1.31	0.233	10.2	3.99	52.5	22.2	106	24.1	245	50.9
07XT06-1-09	岩浆核	0.071	2.25	0.150	1.95	5.94	0.306	37.6	14.0	149	49.2	191	36.3	309	55.7
07XT06-1-12	岩浆核	0.163	32.5	0.317	3.16	4.51	1.19	20.9	6.69	79.3	29.9	137	28.5	274	53.4
07XT06-1-27	岩浆核	0.000	2.12	0.030	0.452	1.75	0.208	11.2	3.94	47.8	17.8	80.0	17.1	164	33.6
07XT06-1-15	岩浆核	0.002	2.03	0.042	0.859	3.04	0.277	26.2	11.7	167	73.3	356	78.9	759	152
07XT06-1-21	岩浆核	0.020	18.3	0.070	1.54	3.96	0.243	25.5	9.63	122	48.6	223	47.2	431	82.7
07XT06-1-25	岩浆核	0.010	9.57	0.060	0.811	3.80	0.042	27.3	12.2	174	73.7	348	75.5	700	132
07XT06-1-03	岩浆核	0.008	21.0	0.326	5.17	8.70	2.22	43.2	14.2	161	59.9	267	54.9	513	102
07XT06-1-13	岩浆核	0.011	7.05	0.065	1.59	3.17	0.530	16.0	5.19	61.4	24.3	112	24.2	233	47.1
07XT06-1-06	岩浆核	0.036	4.68	0.024	0.179	1.11	0.080	11.0	5.54	84.0	41.2	228	55.6	552	111

的观点<sup>[8]</sup>,它更可能是一个独立的块体<sup>[15-16,41,49]</sup>。

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## Variscan orogeny of Central Tianshan Mountains: Constrains from zircon U–Pb chronology of high-grade metamorphic rocks

LIN Yan-hao<sup>1</sup>, ZHANG Ze-ming<sup>2</sup>, HE Zhen-yu<sup>2</sup>, DONG Xin<sup>2</sup>, YU Fei<sup>1</sup>

(1. Faculty of Earth Science, China University of Geosciences, Wuhan 430074, Hubei China;

2. Institute of Geology, Chinese Academy of Geological Sciences, Beijing 100037, China)

**Abstract:** In this paper, the authors have studied zircon U–Pb chronology of the high-grade metamorphic rocks on the northern margin of the Central Tianshan Microplate. The dating results show that the detrital zircons from the meta-sedimentary rocks have recorded the variable magmatic thermal events from Archean to Neoproterozoic (3320 to 530 Ma), the magmatic zircons of the meta-igneous rocks were crystallized in late Neoproterozoic of 550 Ma, whereas the metamorphic zircons were formed in Devonian of 385–360 Ma. These data demonstrate that the Central Tianshan block was probably formed in Precambrian, and subsequently subjected to Variscan orogeny. Unlike the Tarim Craton with multi-stage Proterozoic tectono-thermal events, the Central Tianshan Microplate which occurred as a single continent did not experience the Precambrian metamorphism.

**Key words:** Central Asian orogenic belt; Central Tianshan Mountains; metamorphism; zircon U–Pb dating; Variscan orogeny

**About the first author:** LIN Yan-hao, male, born in 1987, undergraduate, mainly engages in the study of petrology; E-mail: linyanhao1987@sina.com.

**About the Corresponding author:** ZHANG Ze-ming, male, born in 1961, senior researcher, engages in the study of petrology and metamorphic geology; E-mail: zmzhang@cags.ac.cn.