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东海西湖凹陷西斜坡和西次凹近致密-致密砂岩气藏的地质特征及发育前景

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提要: 基于沉积构造背景和大量钻测资料, 利用气样地球化学测试、铸体薄片、扫描电镜、激光扫描共聚焦显微镜和阴极发光等手段, 揭示了东海西湖凹陷西斜坡和西次凹近致密-致密砂岩气藏的地质特征, 并分析了发育前景。研究结果表明: 2个靶区近致密-致密砂岩气藏均分布在中深层, 且位于常规气藏之下, 但发育层位不同, 西斜坡在平湖组中下段, 西次凹在花港组下段。2个靶区的近致密-致密砂岩气均为煤型气, 来源于平湖组煤系源岩, 西次凹致密气藏属于“下生上储”型, 西斜坡属于“自生自储”型。西次凹属于典型的致密砂岩储层, 西斜坡属于近致密砂岩储层。强压实作用和硅质、自生伊利石及碳酸盐胶结是储层致密化的重要因素。溶孔和原生粒间孔是致密砂岩储层主要储集空间; 2个靶区的近致密-致密储层不同深度段发育“孔隙型”甜点, 为酸性溶蚀成因, 但西斜坡的溶蚀强度大于西次凹。另外, 西次凹深层发育部分裂缝, 形成该区“裂缝型”甜点。西次凹致密气藏具有弱超压特征, 西斜坡具有超压-强超压特征。平湖组广覆式分布高有机质丰度的煤系源岩全天候持续供气是靶区发育大规模近致密-致密砂岩气藏的物质基础, 大面积近致密-致密储层提供了良好的储集体, 因此发育前景好, 勘探潜力大。

关 键 词: 近致密-致密砂岩; 气藏; 地质特征; 发育前景; 西湖凹陷; 中西部

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Development prospects and geological characteristics of nearly tight and tight sandstone gas reservoirs in western secondary sag and western slope zone of Xihu depression, East China Sea

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Abstract: Based on tectonic and sedimentary background as well as a large number of drilling and logging data, the authors studied geological characteristics and development prospects of nearly tight and tight sandstone oil and gas reservoirs in western secondary sag and western slope zone of Xihu depression, East China Sea by using casting thin section, laser scanning confocal microscope, scanning electron microscope, cathodoluminescence, stable carbon isotope composition and light hydrocarbon components. The results indicate that nearly tight and tight sandstone gas reservoirs are distributed under the conventional gas reservoir, but their development horizon are different, belonging respectively to middle and lower Pinghu Formation and lower Huagang Formation. Nearly tight and tight sandstone gas is coal-type gas and came from coal measures source rock in Pinghu Formation. The reservoir is above source rock in western secondary sag, and multiple reservoir and source rock are superimposed upon each other in western slope zone. Western secondary sag reservoir belongs to tight sandstone and western slope zone reservoir is nearly tight sandstone. Strong compaction, siliceous cement, authigenic illite cementation and carbonate cementation are important factors for densification of the reservoir. The pore types of two zones mainly include intergranular dissolved pore, native intergranular pore and feldspar intragranular dissolved pore. There are relatively high porosity and high permeability reservoirs in different depths of the two zones, resulting from acid dissolution. However, western slope zone reservoir has larger corrosion scale and does not develop microfracture. There is a certain size of microfracture which is “microfracture” sweet spot in deep layer of western secondary sag. The former has weak overpressure and the latter has overpressure—strong overpressure. The source rocks which have high abundance of organic matter in Pinghu Formation are distributed widely in Xihu depression. The supply of continuous gas constitutes the material basis of large-scale development of nearly tight and tight sandstone gas reservoirs. Large-sized nearly tight and tight sandstone provides effective reservoirs. So, there are rich tight sandstone oil and gas resources in East China Sea.

Key words: tight sandstone gas; geological characteristics; development prospects; Xihu depression

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致密砂岩气藏是指赋存于低孔隙度($<12\%$)和低渗透率($1\times10^{-3}\mu\text{m}^2$)砂岩中的非常规天然气藏,其含气饱和度低($<60\%$),含水饱和度高($>40\%$)^[1-3]。目前,致密砂岩气产量呈日益增加的趋势,已占全国天然气总产量的四分之一^[4],因此致密砂岩气的勘探与开发已受到外界日益重视。

中国东海海域大规模油气勘探始于20世纪70年代初,迄今主要在西湖凹陷发现了丰富的天然气和凝析油,凹陷探明储量巨大而探明率低,具有很好的勘探前景。目前,随着东海盆地勘探程度的增加,发现西湖凹陷中深层普遍发育近致密-致密砂岩储层,且油气显示丰富。常规开采这些低渗致密气层,其自然产能较低甚至无自然产能。2008年首次尝试应用压裂技术和2011年应用水平井技术进行开发,获得了工业气流^[5],并陆续拉开了大规模勘探与开发东海近致密-致密砂岩油气资源的帷幕,受到了外界强烈关注。2012年最新油气资源评价显示,东海油气资源丰富,其中低渗致密油气资源占凹陷总量的80%以上^[6]。因此,应用多段压裂和多分支水平井技术开发致密砂岩气藏,是东海今后

最重要的方向。东海致密砂岩气勘探及研究程度尚处起步阶段,因此当务之急是弄清该区致密砂岩气藏的地质特征,分析其形成条件与前景,同时与国内其他盆地致密砂岩气藏相比,剖析其自身独特之处。西部斜坡带和西部次凹是西湖凹陷2个重要的构造带,本文以这2个重点靶区为例,在大量钻测资料的基础上,应用气样的地球化学特征、铸体薄片、冷阴极发光、激光扫描共聚焦显微镜和电子扫描显微镜,研究已钻遇层位低渗致密-致密砂岩气藏的地质特征并进行对比,同时分析其形成条件与前景,以期为更深层和其他区带此类油气藏的勘探提供一些思路。

1 地质背景

西湖凹陷位于东海陆架盆地东北部,面积约 $5.9\times10^4\text{ km}^2$,沉积厚度超过10 km,凹陷自西向东可划分为西部斜坡带(简称“西斜坡”)、西部次凹(简称“西次凹”)、中央反转构造带、东部次凹和东部断阶带4个二级构造单元;其中西次凹和西斜坡邻近相靠,位于凹陷中西部(图1)。地层自下而上分别由

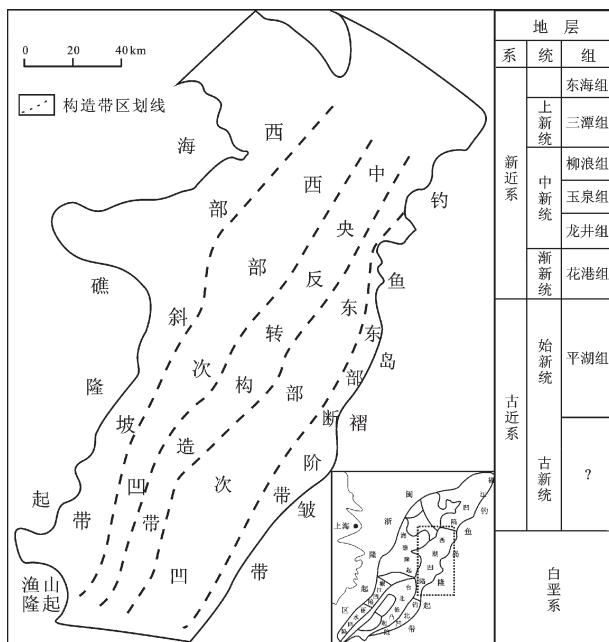


图1 西湖凹陷西斜坡和西次凹区域位置与地层
(引自张建培^[7],有修改)

Fig.1 Location, structural division and strata of Xihu depression

古新统(E_1)、下始新统宝石组(E_{2b})、中上始新统平湖组(E_{2p})、渐新统花港组(E_{3h})、中新统龙井组(N_1^1l)、玉泉组(N_1^2y)和柳浪组(N_1^3l)、上新统三潭组(N_2s)及第四系东海群(Qd)组成^[8-9],其中平湖组和花港组是主要目的层^[10]。

2 近致密-致密砂岩气藏分布特征

西湖凹陷西斜坡和西次凹早期勘探发现的中浅层油气藏均为常规油气藏,油气主要分布在断块以及反转形成的背斜、断背斜圈闭中。近年来,随着钻井深度增加,除发现中浅层存在常规油气藏外,深度3300~3500 m(11口井统计),钻遇低渗近致密-致密气层,常规开采自然产能很低,但利用压裂和水平井技术进行开采,获得高产气流。图2可以看出,西湖凹陷西次凹X气田气层与物性有关,3500 m以上发育常规油气层,其孔隙度大于12%,渗透率大于1 md,3500 m以下的气层需要通过压裂等储层改造技术才能获得气流,其孔隙度小于12%,渗透率小于1 md,为致密砂岩储层。因此,东海盆地西湖凹陷的近致密-致密砂岩气藏与上部常规油气藏呈纵向叠加分布,与鄂尔多斯盆地苏里格

致密气田不同^[11],并非单独孤立分布。西斜坡和西次凹近致密-致密砂岩气藏均具有如此规律,例如西次凹某井花港组自上而下划分为H1~H12油气层,而从H5到H12均为低渗致密层,西斜坡平湖油气田深层段P11和P12也是如此。但2个靶区致密储层发育层位不同,西斜坡近致密-致密砂岩储层多位于平湖组中下段,西次凹则多位于花港组下段。

3 近致密-致密砂岩气成因特征

笔者统计了目前西湖凹陷西斜坡和西次凹部分近致密-致密砂岩气层中20余个天然气气样的组分、碳同位素和轻烃数据。将气样的 $\delta^{13}\text{C}_1$ 与 $\text{C}_1/\text{C}_{2+3}$ 值投入Whiticar天然气成因鉴别图^[12](图3-a),发现2个靶区的近致密-致密砂岩气均为III型干酪根热成因气(煤型气),即气源来自煤或腐殖源岩。利用 $n\text{C}_7$ 、MCC₆和 $\sum \text{DMCC}_5$ 等轻烃参数可有效识别天然气母质类型^[13],从图3-b中可看出气样也主要分布在煤型气区。综上,与目前国内其他盆地(鄂尔多斯、四川和塔里木盆地等)一样,东海盆地的致密砂岩气均为煤型气。气源岩为平湖组煤系源岩,结合上述致密气藏分布层位,可知西次凹近致密-致密砂岩气藏的生储配置组特征为“下生上储”型,西斜坡为“自生自储”型。

4 近致密-致密砂岩储层特征

4.1 岩石学特征

据统计的西湖凹陷西斜坡12口代表性单井和西次凹4口代表性单井的中深层近致密-致密砂岩储层岩样的砂岩成分,西次凹近致密-致密砂岩储层的岩性主要为岩屑长石砂岩和长石岩屑砂岩,而西斜坡则主要为长石岩屑砂岩,总体上西次凹砂岩石英含量明显低于西斜坡(图4);前者储层颗粒主要为细砂,部分为中砂;后者则主要为细砂和极细砂。2个靶区碎屑颗粒分选中等-较好,磨圆度次棱角状-次圆状,颗粒接触关系均为线状-凹凸状接触,颗粒之间主要充填碳酸盐胶结、粘土矿物和硅质胶结,胶结类型为再生和孔隙式胶结。

4.2 物性特征

西次凹和西斜坡16口代表性单井的近致密-致密砂岩气层,共200余个样品点的岩心孔隙度、渗透率和孔径的测试统计结果表明,2个靶区致密储层

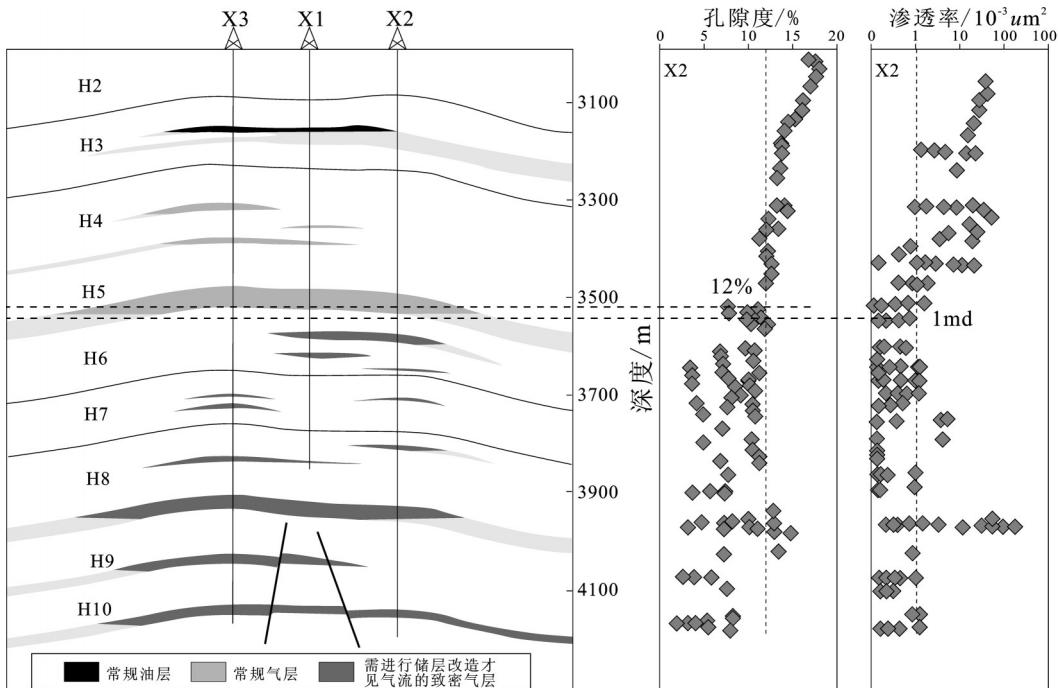


图2 西次凹X气田油气藏剖面与X2井物性剖面
Fig.2 Reservoir section in X gas field, western secondary sag and property section in X2 well

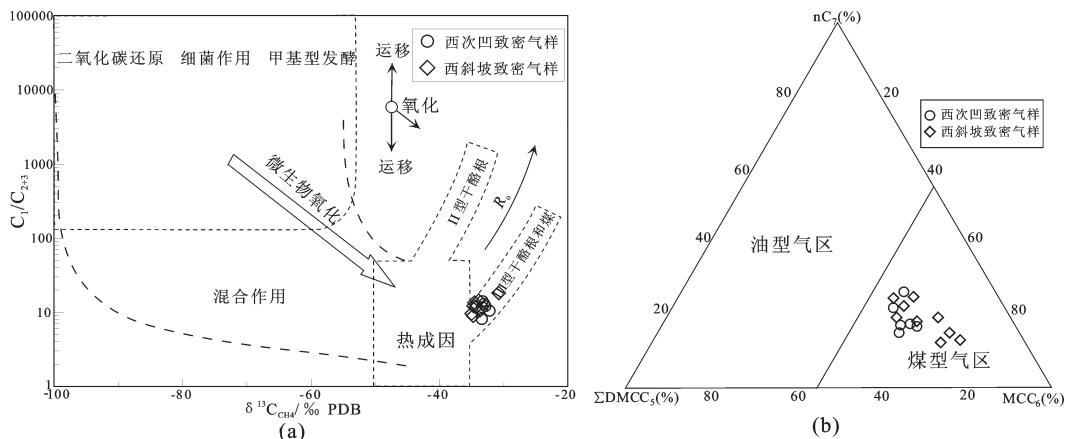


图3 西湖凹陷西次凹和西斜坡近致密 – 致密砂岩气成因识别
Fig.3 Genetic identification of nearly tight and tight sandstone gas in western secondary sag and western slope zone of Xihu depression

物性有差异,西次凹花港组下段孔隙度主要分布在5%~11%,平均孔隙度为8.6%,渗透率主要分布在0.1~10 md,平均为1 md左右(图5-a),依据中国石油天然气行业标准,为典型的致密砂岩储层;西斜坡平湖组中下段平均孔隙度为10.5%,渗透率多介于1~50 md(图5-b),为近致密砂岩储层。压汞和激光共聚焦显微镜直接测试了储层的孔喉直径,其中西次凹致密砂岩储层孔喉直径介于0.08~8 μm,

西斜坡近致密砂岩储层孔喉直径分布较广,但多集中在1~10 μm。与国内外8个盆地致密砂岩储层的孔喉直径并进行对比(图5-c),国内外盆地致密砂岩储层孔径绝大部分小于2.2 μm,目前统计的西次凹致密砂岩的孔径稍大于这个值,但主体仍小于2.2 μm,西斜坡平湖组中下段储层孔径大些,为近致密砂岩的级别。结合渗透率、孔隙度和孔径,综合认为西次凹花港组下段为致密砂岩储层,而西斜坡

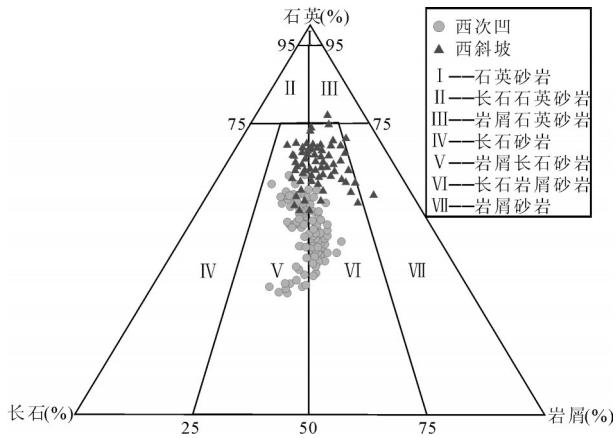


图4 西次凹和西斜坡近致密-致密砂岩储层砂岩成分分类三角图

Fig.4 The composition classification triangle of nearly tight and tight sandstone reservoirs in western secondary sag and western slope zone

平湖组中下段为近致密砂岩储层。

薄片镜下观察发现,西次凹和西斜坡近致密-致密砂岩储层压实作用较为强烈,颗粒之间几乎均为线接触和凹凸接触(6-a),而且石英次生加大较为明显(图6-b),从扫描电镜也发现了自生石英充填在溶孔之间,堵塞了孔隙空间(图6-c)。除此之外,自生伊利石充填孔隙的现象也较为常见,丝网状自生伊利石搭接在颗粒之间,不仅充填孔隙,更严重地降低了渗透率(图6-d)。西湖凹陷中央反转构造带储层的碳酸盐胶结物含量高,阴极光发光强烈^[18],但西斜坡和西次凹的致密层岩样阴极光较弱,出现部分暗红色阴极光(图6-e),说明2个靶区仅少量碳酸盐胶结物充填颗粒,且含铁量较高。总体上,2个靶区致密砂岩储层致密化因素较为一致。因此,强压实作用、硅质胶结、自生伊利石胶结和碳

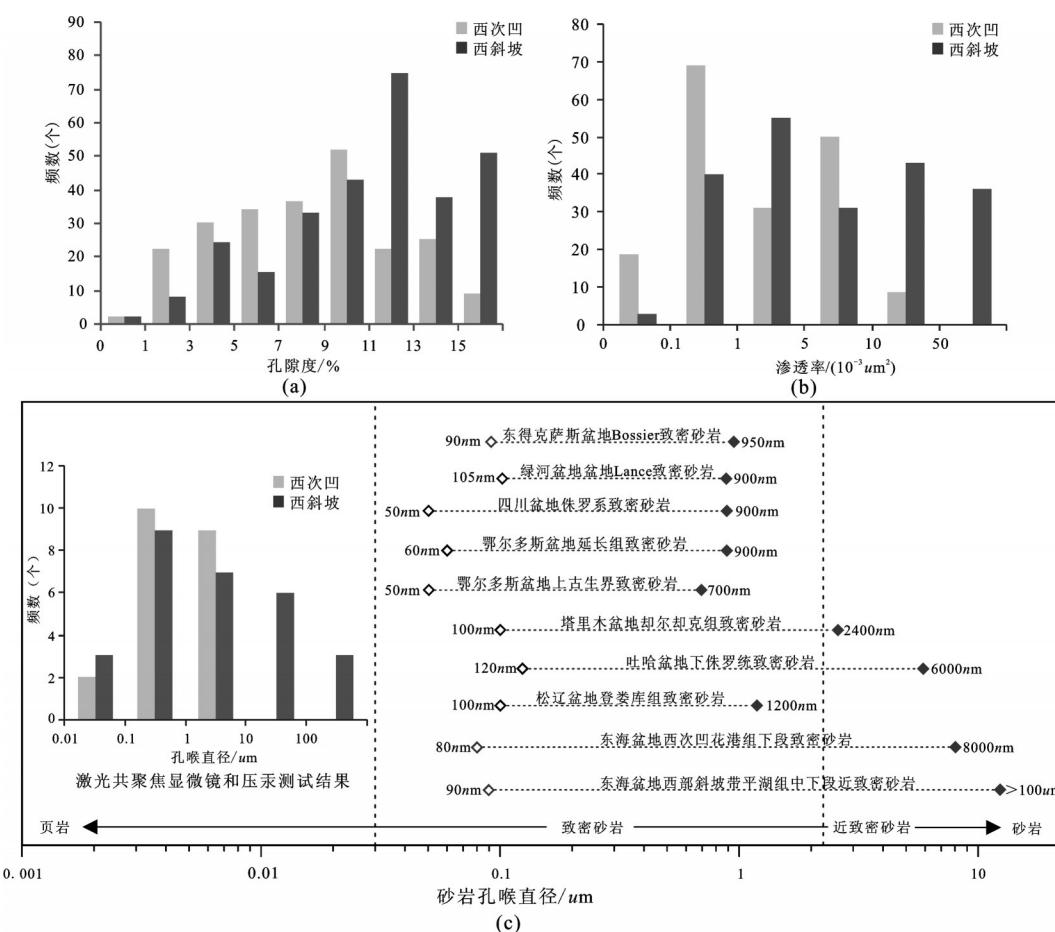


图5 西次凹和西斜坡近致密-致密砂岩储层孔隙度、渗透率和孔喉直径对比图(图中部分数据见文献[14-17])

Fig.5 Comparison of reservoir porosity, permeability and pore throat diameter of nearly tight and tight sandstone gas reservoirs in western secondary sag and western slope zone

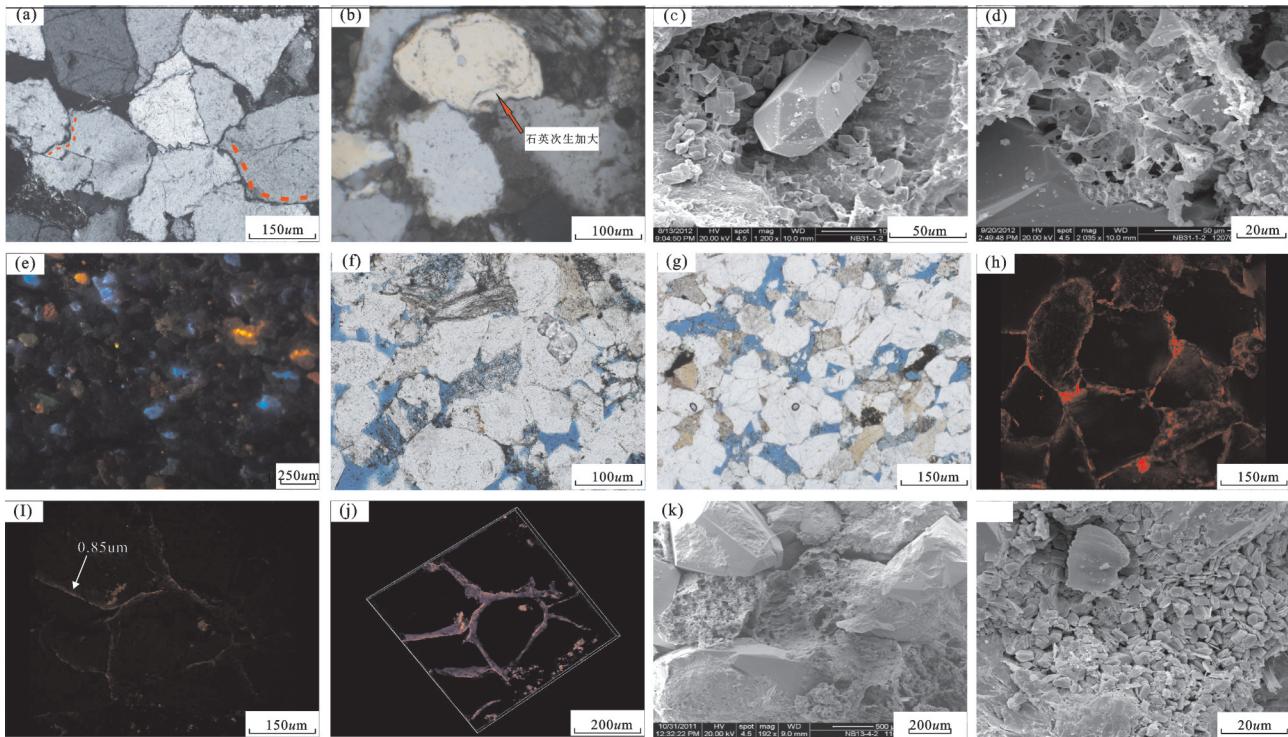


图6 西次凹和西斜坡近致密-致密砂岩储层正交偏光、阴极光、扫描电镜和激光共聚焦的显微照片
(测试仪器为Leica DM5500型激光共聚焦显微镜,测试电压为800V;扫描电镜型号为S4800冷场发射扫描电子显微镜;冷阴极发光型号为RELIION CL 6,测试条件为10kV和500 μA)

a—西斜坡,A1井,4109.86m,颗粒线状—凹凸接触,正交偏光;b—西斜坡,A2井,3685m,石英次生加大胶结充填,正交偏光;c—西次凹,B1井,4426.93 m,溶蚀孔中自生石英占据孔隙,扫描电镜SEM;d—西次凹,B1井,4434.83 m,粒间孔中分布伊利石占据孔隙,堵塞喉道,SEM;e—西次凹,B2井,3982.7 m,碳酸盐胶结物发暗红色阴极光,阴极发光;f—西斜坡,A3井,4197.67 m,长石溶蚀形成粒内孔和粒间孔,蓝色铸体;g—西次凹,B1井,3970 m,次生溶孔较发育,蓝色铸体;h—西斜坡,A3井,4208.67 m,烃类主要存储在粒间孔和粒间溶孔,激光共聚焦照片;i—西次凹,B2井,3982.7 m,烃类主要存储在粒间孔和粒间溶孔,测试孔径为0.85 μm,激光共聚焦照片;j—图6-i的烃类存储的三维模型,激光共聚焦照片;k—西斜坡,A1井,4109.16 m,次生溶蚀孔,扫描电镜;l—西斜坡,A4井,4197.67 m,叠片状高岭石分布在粒间孔,扫描电镜

Fig. 6 Microphotographs of PLM, CL, SEM and CMSL of nearly tight and tight sandstone gas reservoirs in western secondary sag and western slope zone

a—Western slope zone, A1 well, 4109.86m, particle linear—bump contact, PLM; b—Western slope zone, A2 well, 3685m, Pores filled with quartz overgrowth, PLM; c—Western secondary sag, B1 well, 4426.93 m, Authigenic quartz in dissolved pore, SEM; d—Western secondary sag, B1 well, 4434.83 m, illite in intergranular hole plugging the pore and throat, SEM; e—Western secondary sag, B2 well, 3982.7 m, luminescence of carbonate cements is dark red, CL; f—Western slope zone A3 well, 4197.67 m, feldspar corrosion produces intergranular dissolved pore and intragranular dissolved pore, blue casting; g—Western secondary sag, B1 well, 3970 m, secondary solution pores, Blue casting; h—Western slope zone A3 well, 4208.67 m, hydrocarbon stored in intergranular pores and intergranular dissolution pores, CMSL; i—Western secondary sag, B2 well, 3982.7 m, hydrocarbon stored in intergranular pores and intergranular dissolution pores, pore diameter is 0.85 μm, CMSL; j—3D model of hydrocarbon storage in Fig6-i, CMSL; k—Western slope zone, A1 well, 4109.16 m, Secondary dissolution pores, SEM; l—Western slope zone, A4 well, 4197.67 m, laminated sheet kaolinite is distributed in intergranular pores, SEM

酸盐胶结是导致西次凹和西斜坡发育近致密-致密砂岩储层重要的因素。

4.3 储集空间特征

激光共聚焦显微镜精度高,其放大倍数可达10000倍^[19],分辨率比一般显微镜高1.4倍,可直观观察孔隙、孔隙结构和裂缝。利用激光共聚焦显微镜、铸体和扫描电镜进行联合观测研究区典型井近致密-致密砂岩储层岩石薄片。结果表明:2个靶区

近致密-致密砂岩储层主要的储集类型为溶孔和原生粒间孔(图6-f,g),高岭石晶间孔次之;其中溶孔主要为粒间溶孔和长石、岩屑粒内溶孔。激光共聚焦显微镜清晰显示了油气在近致密-致密砂岩储层颗粒之间的赋存状态,图6-h,i可知西斜坡和西次凹致密气主要存储在颗粒粒间溶孔和原生粒间孔,部分在粒内溶孔,而且明显看出西次凹存储物性要差于西斜坡。激光共聚焦显微镜还原了油气存储

的三维形态,图6-j中可看出,即使在此如此致密的情况下,致密砂岩中油气连通性仍较好。

统计研究区16口单井发现,西次凹和西斜坡近致密-致密储层段均发育深度不一的次生孔隙带,即在低孔低渗储层中发育了相对高孔高渗的储集体(“甜点”)。例如西次凹X2井大于在4000 m左右出现异常高孔带,即孔隙度最高可达到14%,渗透率可达100 md(图2)。甜点深度段次生溶孔特别发育,即溶蚀作用不仅使得孔隙度得到很大的改善,孔喉也得到一定改善(图6-k)。粒间溶孔常常伴随着书叠状分布的自生高岭石充填(图6-l),说明溶蚀孔与高岭石产出有关,应为长石、岩屑酸性溶蚀导致。随机抽取部分铸体,统计得到西斜坡的次生面孔率/总面孔率为77.4%,而西次凹为58.2%,说明西斜坡近致密-致密储层相比于西次凹经历了较强的溶蚀作用,同时对应西斜坡自生高岭石含量为2.0%和西次凹为0.2%,这也进一步证明了该区溶蚀作用主要为酸性溶蚀。大量薄片观察统计,西斜坡岩石薄片均未发现明显裂缝,而西次凹仅部分样品局部发育裂缝,不具代表性。从孔渗关系可以看出,2个靶区总体孔渗关系较好,但西次凹孔渗曲线存在异常段,孔隙度从6%减小至2%,而渗透率维持在0.1~0.04 md(图7-a),也表明西次凹致密储层下部可能发育了部分裂隙。因此,裂缝不是东海西湖凹陷西次凹和西斜坡近致密-致密砂岩储层主要的有效储集空间和流体渗流通道,不具典型性,与吐哈盆地柯柯亚地区致密砂岩气藏的高“裂缝型”甜

点不同。因此从油气储集类型角度来看,西斜坡几乎不存在“裂缝型”甜点,几乎全为“孔隙型”甜点,西次凹主要以“孔隙型”甜点为主,深层局部发育“裂缝型”甜点。

5 压力发育特征

西次凹6口单井共26个钻杆测试(DST)压力数据显示,该区3500 m以下(花港组下段)地层压力系数逐渐增加,到达4300 m左右时,压力系数达到了1.38,平均压力系数为1.29,地层主要在3500 m以下逐渐发育弱超压。西次凹代表性单井A井的声波时差、视电阻率和深度的关系图(图8)可知,深度在3400~3500 m,声波时差偏离正常演化趋势线,出现异常偏高区,而视电阻率则偏离正常演化趋势线,出现异常偏低区,也指示此深度段以下发育超压,与DST指示一致。前已述及,该区也大约在3500 m段以上发育常规气藏,之下发育低渗致密砂岩气藏,气藏类型与压力发育具有明显对应关系,即西次凹致密砂岩气藏具有弱超压的特征。西部斜坡带10口单井的43个DST压力数据与B井的声波时差、视电阻率与深度分布关系表明,西斜坡地层在3550 m左右出现超压,到4400 m左右时,最大压力系数可达1.73,达到了强超压的标准,因此西斜坡近致密-致密砂岩气藏表现出超压-强超压的特征。综上,西湖凹陷2个靶区都发育超压,但程度不同。因而西湖凹陷致密砂岩气藏与川西致密砂岩气藏一样^[20],发育异常高压,与鄂尔多斯盆地上古生界致

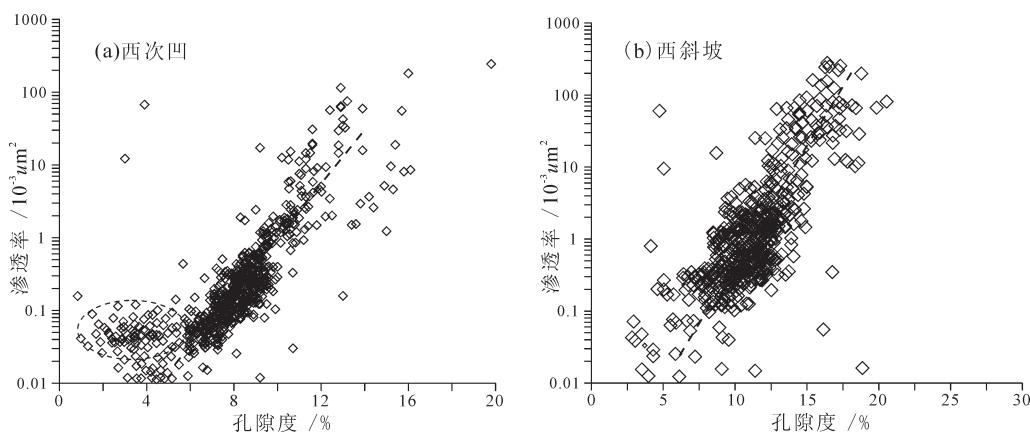


图7 西次凹和西斜坡近致密-致密砂岩储层孔隙度-渗透率关系

Fig.7 The relationship between porosity permeability of nearly tight and tight sandstone gas reservoirs in western secondary sag and western slope zone

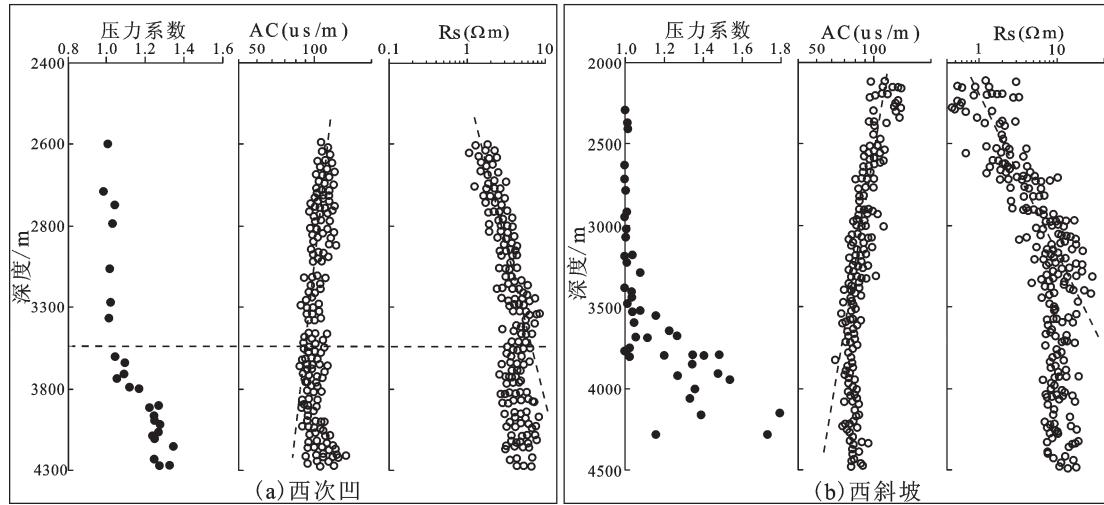


图8 西次凹和西斜坡实测地层压力及声波、视电阻率与深度的关系

Fig.8 Diagram of measured stratigraphic pressure, AC and Rs versus depth in western secondary sag and western slope zone

密砂岩气藏^[21]具有异常低压不同。

6 近致密-致密砂岩气藏发育前景

6.1 广覆式分布富有机质的煤系源岩

平湖组早期,西湖凹陷在继承了始新统早期断陷的基础上,继续拉张裂陷,凹陷与外海连通,发生了一定海侵,直到平湖组晚期,东部钓鱼岛隆升,凹陷逐渐与外海割断,海水逐渐退去,期间发育了大量海陆过渡相地层。与此同时,当时古气候为亚热带湿热气候^[22],植物大量发育,而且凹陷总体上呈缓慢沉降。因此,平缓古构造地形、频繁海进海退和温湿古气候有利于平湖组泥炭沼泽环境的发育,从而形成了广泛分布的沼泽相煤系源岩,为西斜坡和西次凹近致密-致密砂岩气藏提供气源。图9-a可看出,整个凹陷几乎均有一定厚度的煤层和炭质泥岩,西次凹和西斜坡中部其煤层和炭质泥岩较为富集。

凹陷煤系源岩热演化程度 Ro 基本均大于 0.6%,已经进入大规模生气门限(图9-a)。盆地模拟显示西次凹平湖组煤系源岩于中新世中晚期开始进入大规模排烃阶段^[23],西斜坡平湖组源岩也于上新世早期进入大规模排烃阶段,因此该区源岩均发生了煤生成气和排烃作用,具有广覆式生气的特征。

有机显微组分三角投点图(图9-b)显示,煤系源岩有机质主要分布在腐殖型区,仅有极个别分布在腐泥型区,生烃组分热模拟实验表明腐殖型干酪根全阶段产气。腐殖母质相较于腐泥母质,具有全

阶段持续生气的优势,因此平湖组煤系源岩作为气源岩,能够保证“全天候”连续不断的供气,这样才有利于形成大规模致密气藏。即使因构造运动导致致密气藏发生泄露,但足量天然气能维持充注与逸散平衡,以至于气藏不会发生破坏。

钻井取心发现平湖组煤系源岩主要由煤岩、炭质泥岩和暗色泥岩组成,即巨厚的深灰及灰黑色泥岩、粉砂质泥岩为主夹多层薄煤层的煤系地层,共同组成了平湖组富有机质的煤系烃源岩。岩石热解法测试该区煤、炭质泥岩和暗色泥岩生烃潜力。煤岩产烃潜量(S_1+S_2)平均为 200 mg/g;有机碳含量(TOC)平均为 55.1%;氢指数(HI)平均为 335 mg/g。煤岩有机质丰度最好,炭质泥岩次之,暗色泥岩最低(图9-c)。总体上,西湖凹陷煤系源岩属于“中等-好”源岩。

综上,西湖凹陷西斜坡和西次凹平湖组煤系源岩分布广泛,也进入了大规模生排气阶段,而且煤系具有持续生气的特点,同时有机质丰度高,生气潜力大,这些为 2 个靶区提供了充足、持续的气源,是形成大规模近致密-致密砂岩气藏的物质基础。

6.2 大面积近致密-致密储层发育

西湖凹陷中深层发育大片致密砂岩,其中西次凹花港组下段主要储集体为三角洲前缘和三角洲平原,西斜坡平湖组中下段则主要为潮控三角洲和潮汐海岸。凹陷内广泛发育的多个三角洲前缘砂相互交汇,形成大片连续的致密砂体。西湖凹陷部

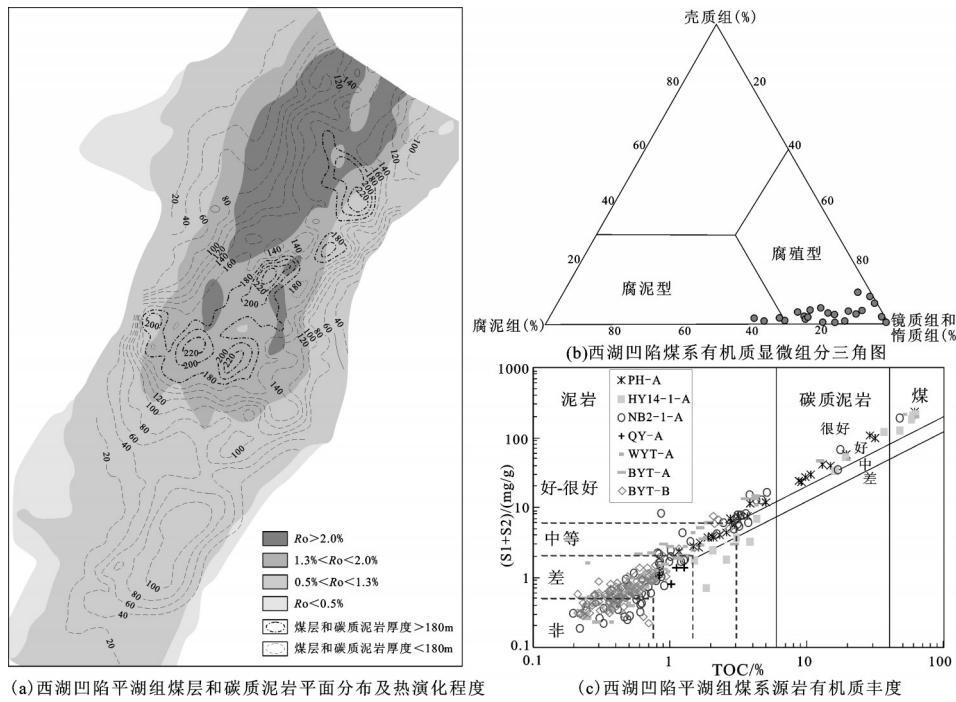


图9 西湖凹陷平湖组煤系烃源岩特征

Fig.9 The characteristics of coal-measure source rock of Pinghu Group in Xihu depression

分连井沉积相(图10)可知,平湖组中下段大面积分布近致密砂岩储层与煤系源岩相互接触或互层叠置分布,有利于煤系源岩生排气直接进入近致密储层,属于砂岩-泥岩互层型储层;西次凹储层则直接接触并位于源岩上方,气源岩向上方储层进行供气。西斜坡和西次凹发育的近致密-致密储层为天然气存储的有效载体。

6.3 发育前景

综合而言,西湖凹陷西斜坡和西次凹发育的优质煤系源岩提供了源源不断的气源,与此同时2个靶区发育的近致密-致密储层为天然气充注提供了储集体,而且还均存在物性较好的甜点,因此2个靶区具有发育大规模近致密-致密砂岩气藏的前景,勘探潜力大。值得注意的是,受钻井数和钻深的限制,目前研究区揭示的致密气层仅为冰山一角,推测在花港组和平湖组更深的储层可能发育数量更多、更典型的致密储层;例如目前钻遇的西斜坡平湖组主要为近致密砂岩气藏,可能属于常规气藏向典型致密气藏的过渡型,随着埋深的增加,下部将会发育大片典型的致密砂岩气藏,因此深层是重要的勘探方向。另外,西湖凹陷中央构造带中北部以往被认为油气潜力差于南部,事实上从图9-a中可知,中北部煤层更厚,

而且热演化程度更高, Ro 可达2.0%以上,因此气源更为充足,中北部致密气的资源潜力可能要高于南部,也应为更重要的勘探目标。

7 结 论

(1)2个靶区近致密-致密砂岩气藏均分布在中深层(约在3500 m以下),且位于常规气藏之下,但发育层位不同,西斜坡在平湖组中下段,西次凹在花港组下段。

(2)西斜坡和西次凹近致密-致密砂岩气均为煤型气,来源于平湖组煤系源岩,西次凹致密气藏属于“下生上储型”,西斜坡属于“自生自储”型。

(3)西次凹花下段为较为典型的致密砂岩储层;西斜坡平湖组中下段为近致密砂岩储层。强压实作用和硅质、自生伊利石及碳酸盐胶结是储层致密化的重要因素。溶孔和原生粒间孔是致密砂岩储层主要的储集空间;西次凹和西斜坡不同深度段几乎均存在“孔隙型”甜点,为酸性溶蚀成因,但西斜坡的酸性溶蚀强度大于西次凹。西斜坡较少发育裂缝,而西次凹深层发育部分裂缝,形成该区“裂缝型”甜点。

(4)西次凹致密气藏的压力系数介于1.1~1.39,

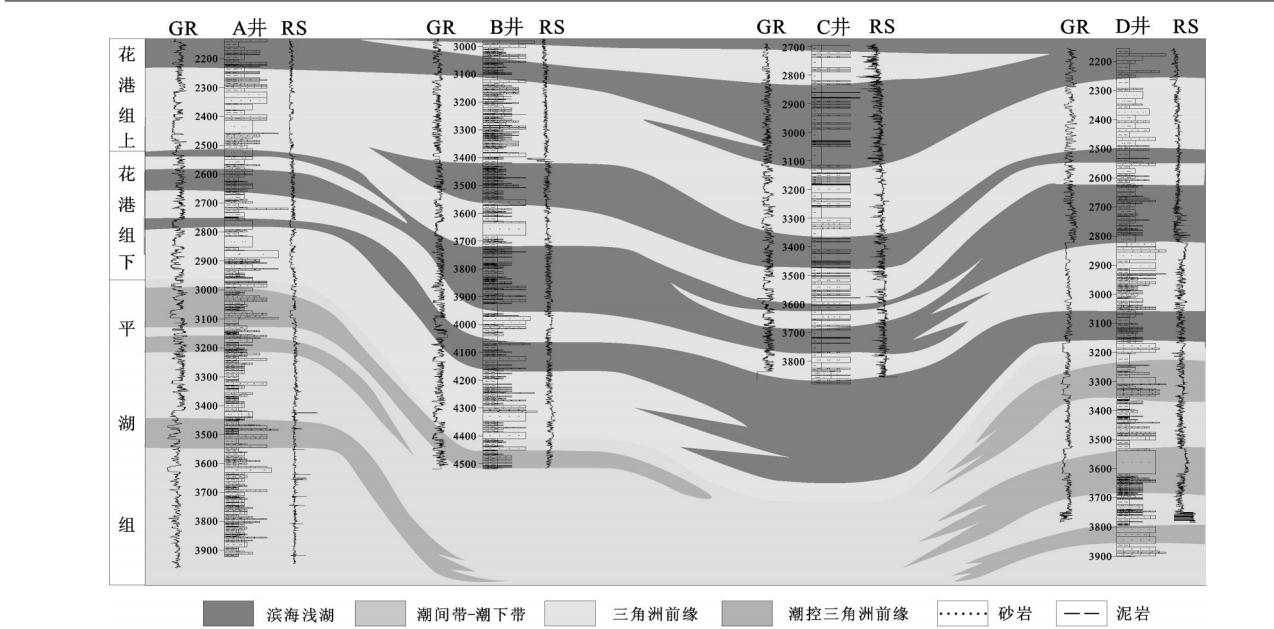


图10 西湖凹陷部分连井沉积相图

Fig.10 Section of well-tied sedimentary facies in Xihu depression

具有弱超压特征;西斜坡近致密气藏的压力系数介于1.1~1.73,具有超压-强超压特征。

(5)平湖组广覆式分布的煤系源岩,进入了大规模生排气阶段,具有持续生气的特点,同时有机质丰度高,生气潜力大,全天候持续供气是2个靶区发育大规模近致密-致密砂岩气藏的物质基础,大面积近致密-致密储层为致密砂岩气藏形成提供良好的储集体,因此发育前景好,勘探潜力大。

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