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雅鲁藏布江缝合带蛇绿岩中铬铁矿的前景讨论

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提要:蛇绿岩地幔橄榄岩中产出的豆荚状铬铁矿是铬的主要来源,是中国极缺的重要战略资源。开展豆荚状铬铁矿成矿作用及围岩地幔橄榄岩的研究,是进一步寻找铬铁矿床和缓解中国铬铁矿资源的瓶颈状态的必要手段。本文以西藏雅鲁藏布江蛇绿岩带内几个主要的地幔橄榄岩体及其中的铬铁矿体为研究主体。在野外地质调查的基础上,系统总结了蛇绿岩的组成、矿物成分、岩石地球化学成分和Re–Os同位素等特征,探讨铬铁矿和地幔橄榄岩的形成过程,取得以下进展和认识:(1)雅鲁藏布江缝合带各段的岩石组合存在较大差异,构造背景的演化过程也不同,佐证了特提斯洋演化过程的不均一性;(2)在雅江西段存在高铝型和高铬型两类铬铁矿矿体,其余都为高铬型铬铁矿,铬尖晶石的矿物化学特征记录了不同构造背景的痕迹;(3)地幔橄榄岩的矿物学和地球化学表明地幔橄榄岩及铬铁矿具有深海地幔橄榄岩和岛弧地幔橄榄岩两者的特点,是岩石/熔体反应和部分熔融作用叠加的结果;(4)提出豆荚状铬铁矿为多阶段形成的认识,经历了早期俯冲到地幔过渡带,在地幔柱/地幔对流驱动下,运移到过渡带顶部冷凝固结,在侵位过程和俯冲带环境,含水熔体与方辉橄榄岩反应的过程;(5)在雅鲁藏布江缝合带中金刚石等超高压矿物的普遍存在,西段的几个大型岩体与罗布莎存在较多相似之处,均经历了相同的构造背景和豆荚状铬铁矿的成矿作用,存在较大的找矿空间。

关 键 词:豆荚状铬铁矿; 蛇绿岩; 多阶段; 雅鲁藏布江缝合带

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The prospects of chromitite in ophiolite of Yarlung Zangbo Suture Zone, Tibet

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Abstract: Podiform chromitite is the main source of chromium, and is a kind of important strategic material of extreme shortage in China; therefore, the study of chromitite and peridotite has become the first way to solve the long-term dependence on imports. In this study, the authors focused on the main ultramafic body and chromitite from Yarlung Zangbo ophiolite belt. On the basis of field geological work, the authors summarized the sequence of ophiolite, mineral composition, geochemistry, Re–Os isotopic composition and other characteristics, so as to discuss the formation of chromite and mantle peridotite. Some conclusions have been reached: (1) The ophiolite in different segments of Yarlung Zangbo suture zone exhibits quite different rock assemblages, with different tectonic evolution backgrounds, which proves the heterogeneity of Tethys evolution; (2) There are two types of high aluminum and high chrome chromitite found in western YZSZ, and the mineral chemical characteristics of chrome spinel have recorded traces of different tectonic settings, mineralogy and geochemistry; (3) The mineralogy and geochemistry of mantle peridotites peridotite and chromite have both the features of deep mantle peridotite and the island arc mantle peridotite, suggesting multi-processes of rock/melt reaction and partial melting; (4) The podiform chromite formation experienced multiple stages, from the early subduction to the mantle transition zone; driven by plume / mantle convection, the materials migrated onto the top of the transition zone, where they condensed and consolidated, and reacted with harzburgite in the emplacement process under a subduction zone environment; (5) The diamond and other UHP minerals extensively exist in all of the Yarlung Zangbo suture zone, with the western part bearing similarity to Luobusha in that they both experienced the same tectonic setting and ore mineralization, thus having potential for further prospecting.

Key words: podiform chromitite; ophiolite; multi-stage; Yarlung Zangbo Suture Zone

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雅鲁藏布江蛇绿岩带位于特提斯喜马拉雅与拉萨地块之间,岩带宽不足15 km,长度超过2000 km^[1],是青藏高原南部最重要的大地构造界线,印度板块与欧亚大陆的最后一条缝合带,代表新特提斯洋盆的残余。较早的研究始于20世纪的60—70年代,在80年代第一次中法国际合作后提出了雅鲁藏布江蛇绿岩带代表印度—欧亚大陆的俯冲—碰撞界线,其北侧为欧亚板块的中—新生代弧前盆地火山—沉积建造、冈底斯岩浆弧和拉萨地块,而其南侧的构造沉积单元则属于印度板块的被动大陆边缘沉积建造和结晶基底,两者之间在中生代时曾经存在数千米乃至近万千米宽的大洋新特提斯洋^[2-11](图1)。之后被认为雅鲁藏布江蛇绿岩带代表喜马拉雅陆弧与拉萨地块之间的狭长弧后裂陷盆地,新特提斯洋可能是弧后或者弧间小洋盆^[12-14]。

雅鲁藏布江蛇绿岩带大致沿雅鲁藏布江分布,根据蛇绿岩体的空间分布特点,将该岩带划分为东段(曲水—墨脱)、中段(昂仁—仁布)和西段(萨嘎以西至中印边境)三部分^[15]。西段又分为北亚带(达机翁—萨嘎蛇绿岩带)和南亚带(达巴—休古嘎布蛇绿

岩带)^[15]。带内蛇绿岩绝大多数已被构造肢解,结构和层序出露不全。东段岩体主要有泽当岩体、罗布莎岩体等。中段蛇绿岩体出露的岩体数量多、规模大,主要包括桑桑、昂仁、柳区、白朗、大竹卡、仁布等,有的岩体之间首尾相连,日喀则岩体规模最大出露面积达1000 km²;西段北亚带(达机翁—萨嘎蛇绿岩带)自萨嘎向西沿马攸木拉、冈仁波齐峰南坡和阿依拉山主脊一线断续展布,岩体规模不大,多呈长条状、透镜状平行于区域构造线和沿断裂带产出,受断裂制约,主要岩体有日康巴、门士、马牧木拉(公珠错北)岩体等。南亚带(达巴—休古嘎布蛇绿岩带),西起札达县达巴乡,经普兰县拉昂错,东至仲巴县帕羊乡休古嘎布一带,呈北西向展布长约400 km,宽10~35 km,由8个蛇绿岩体(群)组成。主要由规模较大的单一岩体组成,岩石新鲜,如普兰、东波、当穷岩体,这些岩体具有明显的岩相分带。

1 东段蛇绿岩

1.1 罗布莎蛇绿岩

该岩体位于拉萨市东南约200 km的雅鲁藏布

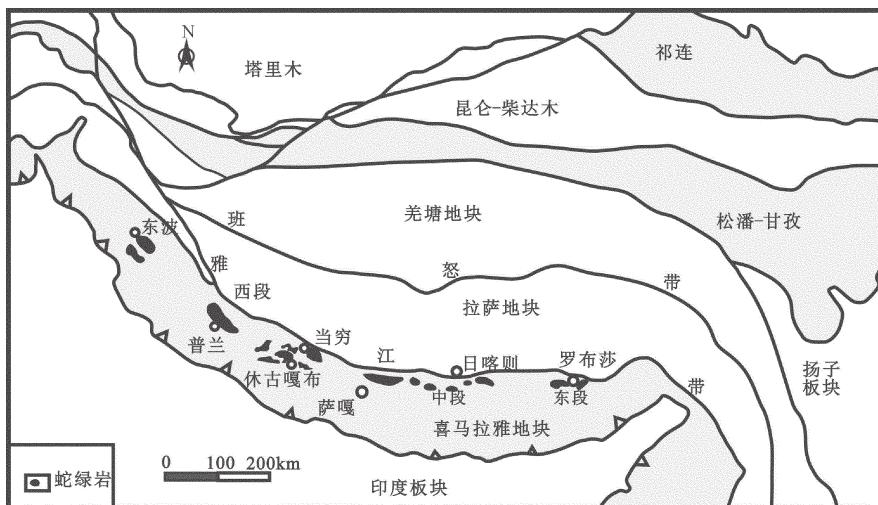


图1 研究区区域地质简图(据文献[16])

Fig.1 Geological sketch map of the study area (after reference [16])

江缝合带东段。罗布莎岩体沿雅鲁藏布江谷地展布,东西延伸约42 km,最宽处约3.7 km,面积达70 km²,平面略呈一平置的反“S”形。该岩体的南侧为上三叠统复理石沉积,与岩体呈断层接触关系,北侧被古近-新近系罗布莎群不整合覆盖。罗布莎蛇绿岩岩块主要由地幔橄榄岩和堆晶岩组成,以及少

量被肢解的火山岩和硅质岩构成的混杂岩出露在堆晶岩的北侧^[1]。地幔橄榄岩主要为方辉橄榄岩和纯橄榄岩,少量的二辉橄榄岩;堆晶岩主要由异剥橄榄岩、辉石岩、纯橄榄岩和辉长岩组成(图2)^[17]。罗布莎铬铁矿床分布于纯橄榄岩-方辉橄榄岩内,为典型的阿尔卑斯型铬铁矿,即豆英状铬铁矿床^[17]。矿体

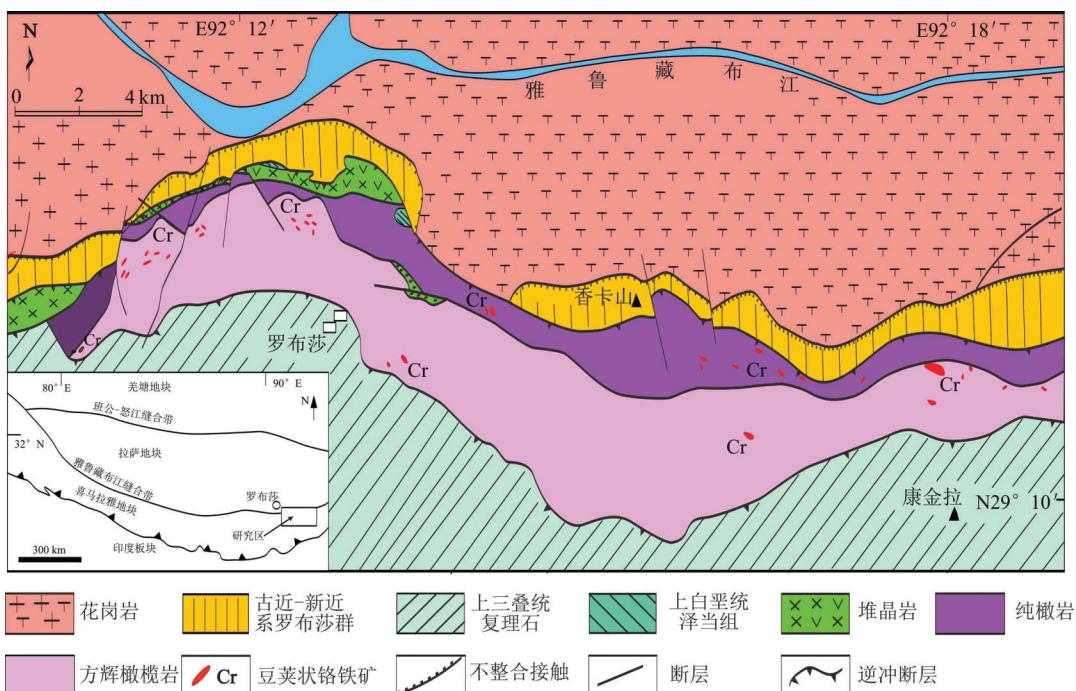


图2 西藏罗布莎岩体及铬铁矿地质简图(据文献[19])

Fig.2 Geological map of the Luobusa ophiolite, Tibet

边缘一般具有一层薄的纯橄岩外壳,矿体和纯橄岩之间不属断层接触,但岩性界线清楚^[18]。铬铁矿矿体多呈似板状、透镜状以及其他不规则形态,并具成群分布,成带集中,侧向延伸有限的特征。根据铬铁矿的开采和地形等因素,罗布莎铬铁矿床自西向东分为罗布莎、香卡山、康金拉3个矿区。到目前为止,罗布莎铬铁矿床共编排15个矿群,除香卡山矿区XI矿群产于堆晶结构纯橄岩相带中外,主要工业矿体多集中产于蛇绿岩壳-幔边界以下方辉橄榄岩相与堆晶状纯橄岩相接触界面之下的方辉橄榄岩相带一侧。

1.2 泽当蛇绿岩

此岩体位于雅鲁藏布江缝合带东段,地处泽当县城以西20 km。该岩体主体呈北西-南东向产出,地势上西北低、东南高。岩体最宽处可达4 km,总出露面积45 km²左右(图3)。

泽当蛇绿岩金鲁剖面以方辉橄榄岩为主,并出露有纯橄岩、二辉橄榄岩-含单斜辉石方辉橄榄岩、辉石岩、辉长岩以及辉绿岩^[20]。纯橄岩以透镜体的形式与方辉橄榄岩呈断层接触关系。辉石岩、辉长岩以及辉绿岩常以岩脉或团块状的形式就位于方辉橄榄岩的内部,显示出较方辉橄榄岩晚期的特点,其中辉长岩根据粒度可明显分为粗粒和细粒的两期。二辉橄榄岩-含单斜辉石方辉橄榄岩与方辉橄榄岩之间的界限在野外出露不明显,整个泽当蛇绿岩的蛇纹石化为轻到中等。方辉橄榄岩是泽

当蛇绿岩的主体,占泽当蛇绿岩总出露面积的70%以上。泽当蛇绿岩南侧为姐德秀组(T_3-K_1)的晚三叠统复理石相粉砂质板岩夹细砂岩,北侧为岛弧火成岩组合,蛇绿岩与它们都呈断层接触。

2 中段蛇绿岩

日喀则-白朗蛇绿岩相比典型的彭罗斯蛇绿岩,是出露层序最好的地区,主要包括冲堆-夏鲁、德几、群让、柳曲等蛇绿岩。从上到下包括:枕状和块状熔岩、席状岩墙群、堆晶杂岩、地幔橄榄岩,并密切共生硅质岩和泥页岩等海相沉积物,其产出类型跟大竹卡蛇绿岩类似(图4)。放射虫硅质岩直接覆盖在枕状熔岩之上;枕状熔岩与下伏席状岩墙群之间没有截然的界线,枕状熔岩下部随着席状岩墙的逐渐增多而过渡为下伏的席状岩墙群,发育明显的冷凝边,可见岩墙侵入枕状玄武岩中。席状岩墙群堆积岩通常是由辉长岩和辉绿岩组成的一套杂岩,其底部不整合于地幔橄榄岩之上。地幔橄榄岩由纯橄岩、斜辉橄榄岩、含透辉石斜辉橄榄岩和二辉橄榄岩组成^[17]。

日喀则-白朗蛇绿岩北侧出露冈底斯中新生代的中-酸性侵入岩和日喀则群(K_2)弧前盆地沉积及错江顶群(E_{1-2})残余弧前盆地滨海相磨拉石,秋乌组(E_2)残余弧前盆地相磨拉石冈底斯组(E_3-N_1)冈底斯相磨拉石;南侧为蛇绿岩、蛇绿混杂岩及嘎学群(J_3-K_1)海相复理石、硅质岩及基性熔岩所组

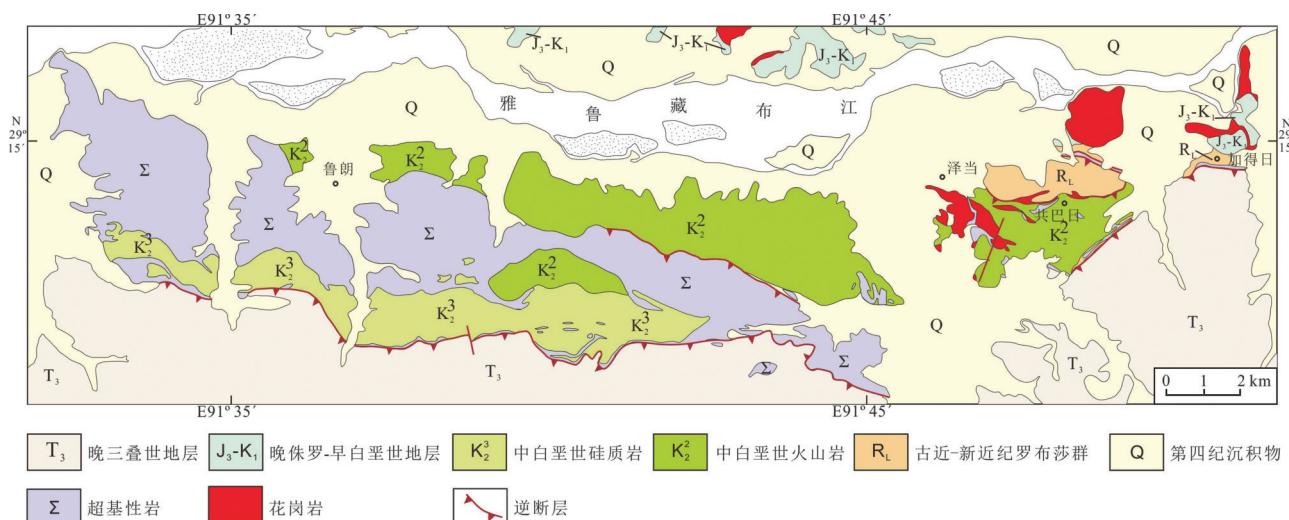


图3 西藏泽当岩体地质简图(据文献[20])
Fig.3 Geological map of the Zetang ophiolite, Tibet(after reference [20])

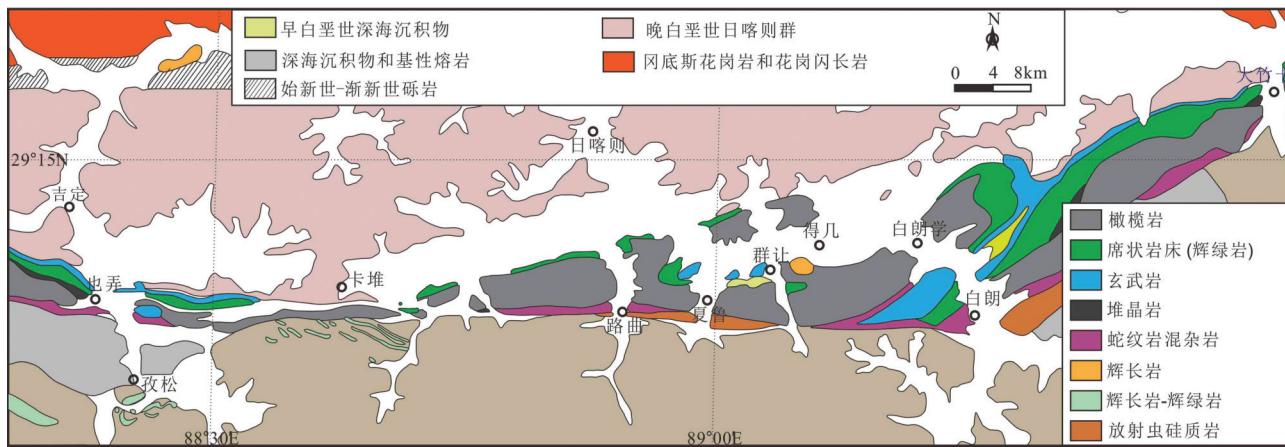


图4 西藏日喀则岩体地质简图
Fig.4 Geological map of the Xigaze ophiolite, Tibet

成的构造单元^[15]。

3 西段蛇绿岩

雅鲁藏布江缝合带地段位于喜马拉雅地块内，地质构造复杂，主体构造线走向为北西西向，平行于雅鲁藏布江缝合带。依据前人的地质构造单元的划分方案^[9,21]，将雅鲁藏布江缝合带西段从北到南划分为冈底斯、阿依拉—仲巴、达巴—休古嘎布、喜

马拉雅4个分区(图5)。

冈底斯地质分区，地层主要为新生界—中生界的白垩系，以火山岩为主，少量碎屑岩或碳酸盐岩，另有少量的石炭系一下二叠统零星分布于中冈底斯。该区南界为达机翁—马攸木断裂带(即雅鲁藏布江缝合带西段)，沿该断裂带有蛇绿岩块体断续分布，同时有上三叠统修康群和侏罗—白垩系海相复理石建造断续分布，夹有大量中基性火山岩，在西

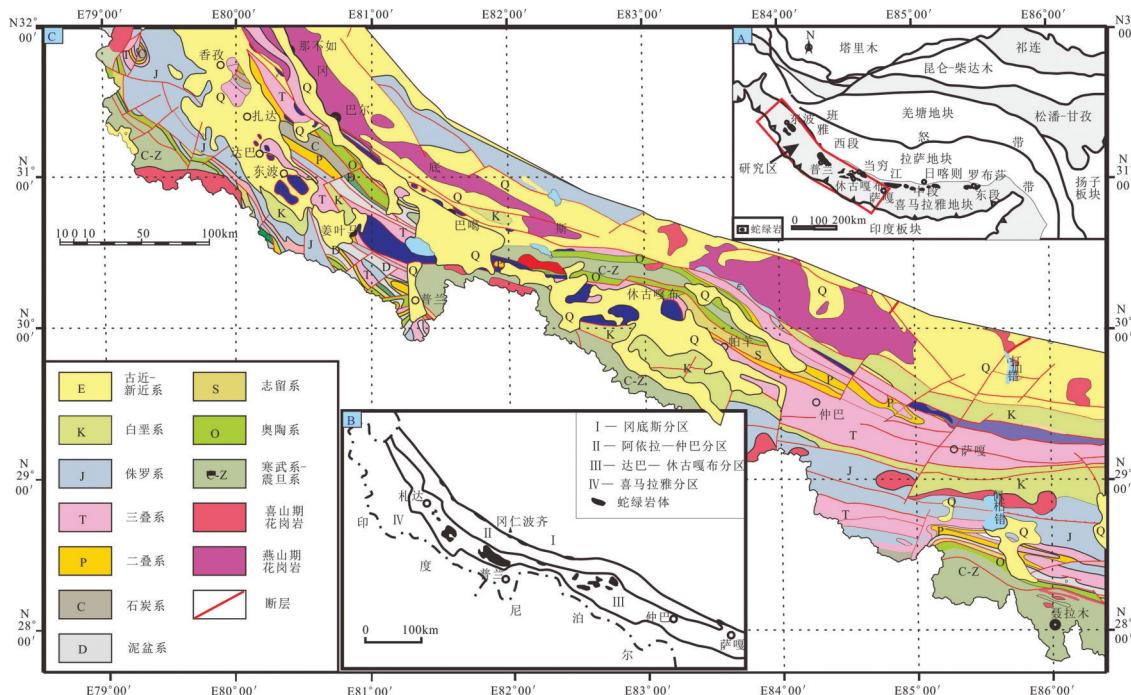


图5 雅鲁藏布缝合带西段岩体地质简图
Fig.5 Geological map of the ultramafic rocks, Tibet

部达机翁一带与阿依拉—仲巴地质分区的古生代地层断层接触^[21];阿依拉—仲巴地质分区,呈北西西向展布,主要为震旦系—古生界地块,也有称之为曲龙—吉拉地体^[21]。包括震旦系—寒武系齐吾贡巴群,主要是绿片岩相的浅变质岩,该套岩石构成了本区的构造基底。奥陶系—三叠系海相稳定型沉积,岩性主要为滨海相砂板岩、碳酸盐岩沉积,共同构成了本区的沉积盖层。除缺失下石炭统外,其余地层较为连续,呈北西西向至近东西向展布,并发育巨型多穹隆复式背斜构造^[22],碎屑锆石年龄谱和Hf同位素证据表明仲巴地块可能是一个独立的微陆块,与拉萨地块并无亲缘性^[23];达巴—休古嘎布主要分布中生界的上三叠统一白垩系,是一套海相碳酸盐、碎屑岩沉积,局部夹少量中基性火山岩,含有二叠系外来岩块。它们被第四系覆盖分割而不连续,并有雅鲁藏布江缝合带南亚带的几个大型超基性岩的产出,包括当穷—休古嘎布岩体、普兰岩体和东波岩体等;喜马拉雅地质分区,沿山脉主脊大面积出露的是前震旦纪中高变质岩系,并有喜马拉雅期花岗岩体侵入。普兰县马甲藏布以西的山脉北坡地层发育较全,由南往北时代变新,震旦—寒武系至泥盆系为连续沉积,由浅海相的碳酸盐岩、碎屑岩组成;上石炭统和下二叠统的冈瓦纳型滨海相沉积;中生界广泛分布为海相碳酸盐岩和碎屑岩沉积^[21]。

3.1 东波岩体

阿里地区的东波超镁铁岩体位于札达县,属雅鲁藏布江缝合带西段的蛇绿岩。雅鲁藏布江缝合带西段中达巴—休古嘎布分区内,其北侧为仲巴古地块。该岩体地表出露近等轴状,总面积约400 km²,中部被新近系不整合覆盖而分成两部分(图6)^[21],呈北西西—南东东向展布。岩体边界与火山岩、三叠系浅变质的浅海—半深海相碳酸盐岩—碎屑岩呈断层接触,或被第四系覆盖(图6);西南缘也与火山岩呈断层接触,火山岩倾向岩体外侧,并在南侧见到岩体底部的炭质板岩倾向岩体内部,在岩体的边界硅质岩直接盖在火山岩或地幔橄榄岩之上。岩体未被构造肢解,出露连续,岩石较新鲜,但在断层通过之处蛇纹石化明显增强,局部成蛇纹岩。

岩体内岩相分带明显,以方辉橄榄岩(Harz)为主,另有纯橄岩(Dun)和二辉橄榄岩(Lhz)。在方辉橄榄岩中常见辉石定向性,主要为北西—南东走向

(290°~340°),少量为北东—南西向(近南北走向)。纯橄岩呈透镜状出露在方辉橄榄岩中(图6),从数十厘米至数十米的宽度皆有,以宽约1 m、长2~4 m为主,走向主要为北西—南东向。少量的二辉橄榄岩仍以透镜状出露于方辉橄榄岩中,宽约50 cm、长约2 m。纯橄岩中铬尖晶石的分布较多,而在方辉橄榄岩及二辉橄榄岩中偏少,在纯橄岩与方辉橄榄岩的接触部位偶见块状铬铁矿矿化。辉石岩和辉长岩在方辉橄榄岩中呈脉状出露,脉体宽度由数厘米至数米不等,定向性明显,为北西—南东走向,切割地幔橄榄岩,脉体总体走向为北西—南东向^[24]。地幔橄榄岩中各岩相组合均发生中等程度的变形,片理发育,大多数片理走向与东波蛇绿岩体展布方向一致,为北西走向,向南西倾斜,倾角45°~60°不等。但有个别产状倾向为北东方向,方辉橄榄岩中铬尖晶石的线理方向为北西—南东走向,同样与岩体的展布方向一致。

3.2 普兰蛇绿岩体

属雅鲁藏布江缝合带西段(图2),距离拉萨1500 km。岩体呈北西西—南东东向带状展布,东西长约60 km,南北宽约20 km,总面积达650 km²。岩体沿拉昂错南边向西北方向延长至姜叶玛,围岩为侏罗系—白垩系浅变质的碎屑岩、含放射虫硅质岩夹碳酸盐岩,除被第四系所覆盖的地段外,岩体四周与围岩均为断层接触(图7)。西部北侧为三叠系海相碎屑岩;中部及东部北侧为侏罗系—白垩系的砂、板岩,硅质岩及中基性火山岩;蛇绿岩西部南侧为三叠系硅质岩、玄武岩;东部南侧为前寒武结晶岩系及新近系陆相磨拉石组合,与蛇绿岩呈构造接触关系,局部可见新近系砾岩不整合覆盖在蛇绿岩之上(图7)^[25]。岩体边部的超镁铁岩岩石破碎,蛇纹石化强烈,在岩体内部岩石新鲜。

普兰蛇绿岩主要由地幔橄榄岩、辉长辉绿岩脉以及玄武岩夹硅质岩组成,岩石新鲜。辉长辉绿岩脉常侵入到地幔橄榄岩中,切割地幔橄榄岩,脉体总体走向为北西南东向,另有一个约10 km²辉长岩岩块出露于普兰蛇绿岩拉昂错南侧(图7)。玄武岩夹硅质岩分布于地幔橄榄岩的东西两侧,均倾向地幔橄榄岩,与地幔橄榄岩呈断层接触。岩体内的地幔橄榄岩以方辉橄榄岩为主,另有纯橄岩、二辉橄榄岩以及少量的橄榄单斜辉石岩组成。岩体内二

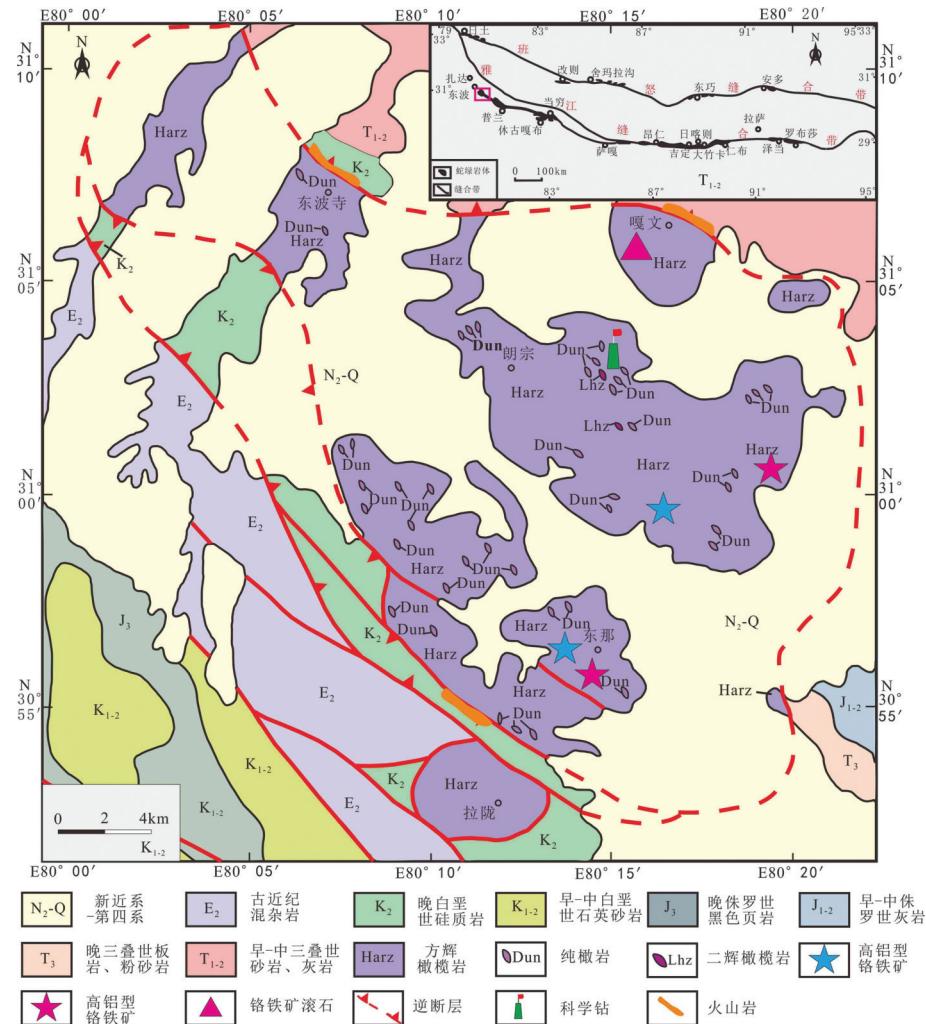


图6 雅鲁藏布江缝合带西段东波蛇绿岩的地质简图(据文献[16])

Fig.6 Detailed geological map of the Dongbo ophiolites in the YZSZ (modified after reference [16])

辉橄榄岩与方辉橄榄岩之间往往呈渐变过渡关系,而纯橄岩、橄榄单斜辉石岩等多呈透镜状或团块状分布于二辉橄榄岩或方辉橄榄岩内,在岩体北部边界出露厚度约为500 m的二辉橄榄岩,直接与围岩砂板岩呈断层接触(图7)。致密块状和稠密浸染状铬铁矿呈透镜状出露于方辉橄榄岩中,呈突变关系,铬铁矿矿体宽为1~3 m,最大长轴方向为北西—南东向,长1~4 m^[25]。地幔橄榄岩中各岩相组合均发生中等程度的变形,片理发育,大多数片理走向与普兰蛇绿岩体展布方向一致,为北西走向,向南西倾斜,倾角45°~60°不等,但有个别产状倾向为北东方向,方辉橄榄岩中铬尖晶石的线理方向为北西—南东走向,同样与岩体的展布方向一致。

3.3 休古嘎布蛇绿岩

长15 km,宽13 km,面积约165 km²,西侧和南

侧被第四系覆盖,东侧和北侧与侏罗—白垩系石英砂岩、硅质岩等断层接触。岩体主要由方辉橄榄岩、纯橄岩及辉绿岩、辉石岩脉等几种岩相组成。方辉橄榄岩岩石新鲜蛇纹石中等,纯橄岩呈透镜体分布于方辉橄榄岩内,两者之间渐变过渡,在南部偶见辉石岩小透镜体,辉绿岩脉群相(30 km²),分布于岩体的南部,成群密集产出,局部在100 m范围内多达20条。岩体的边界主要为混杂岩带,由蛇纹岩和侏罗—白垩系硅质岩经构造作用混杂而成。

4 雅江蛇绿岩中基性岩脉特征对比

在雅鲁藏布江缝合带内基性岩脉大多以岩席或岩脉形式存在于橄榄岩或玄武岩之中。它们的成因解释被认为是洋壳慢速扩张的产物^[28~29],但也有学者认为代表了大洋的快速扩张^[30]。此外,这些

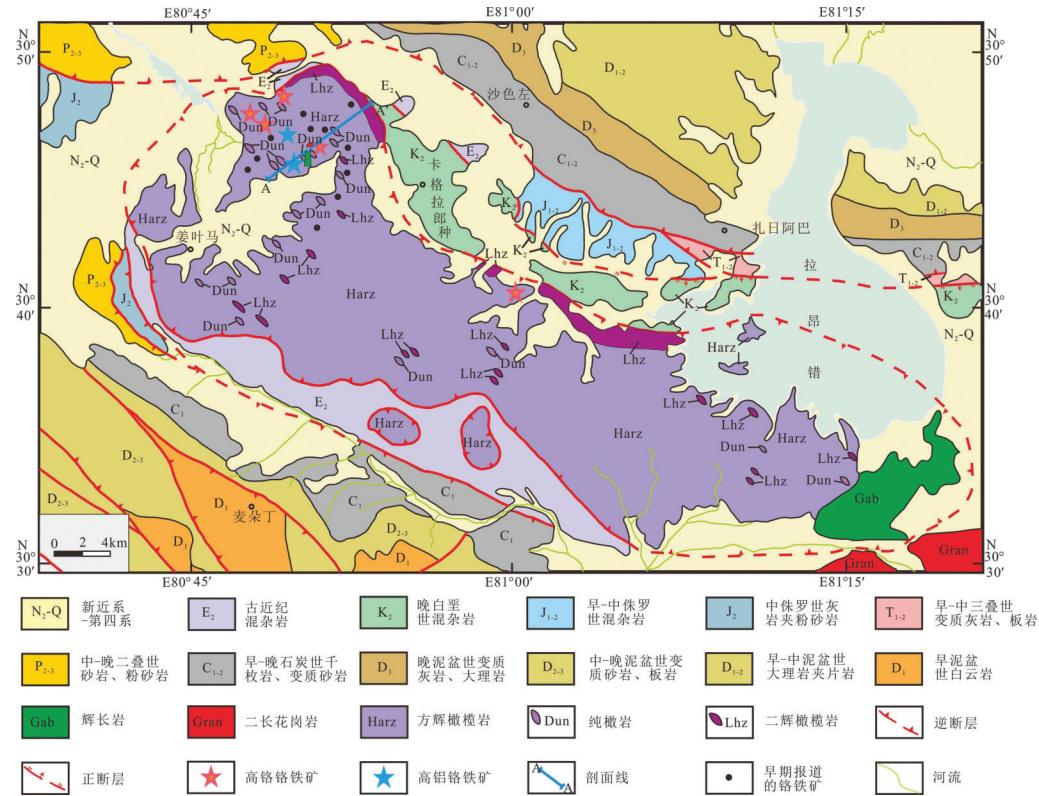


图7 雅鲁藏布江缝合带西段普兰蛇绿岩地质简图(据参考文献[16,26])

Fig.7 Geological map of Purang ophiolite

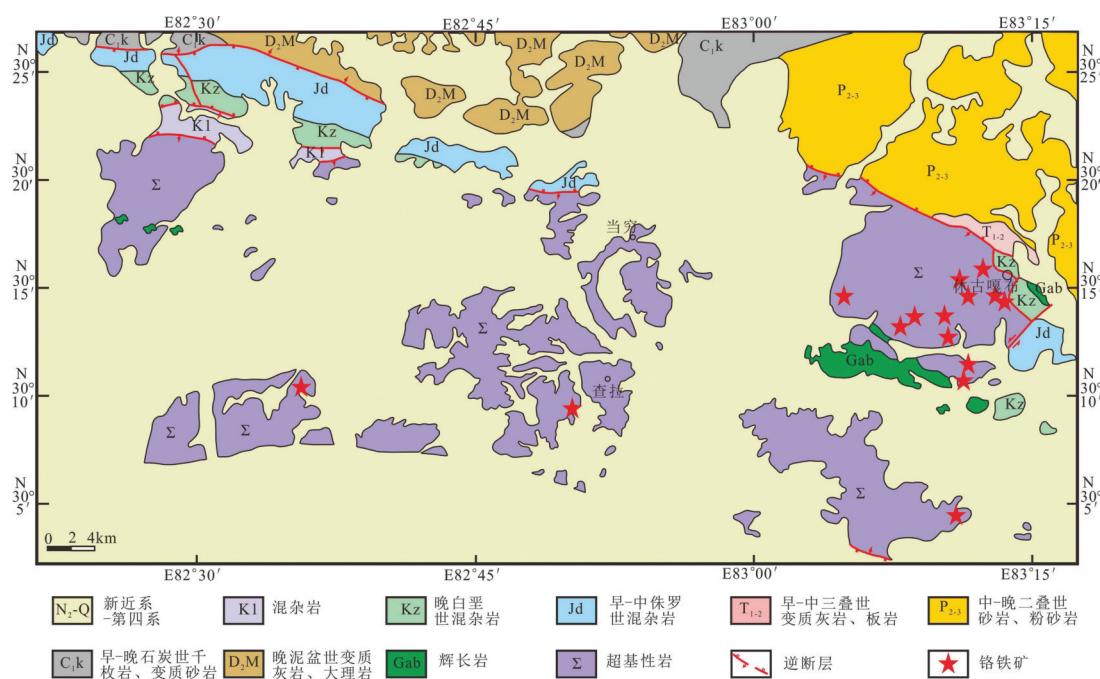


图8 雅鲁藏布江缝合带西段休古嘎布蛇绿岩地质简图(据文献[27])

Fig.8 Geological map of Xiugugabu ophiolite (after reference [27])

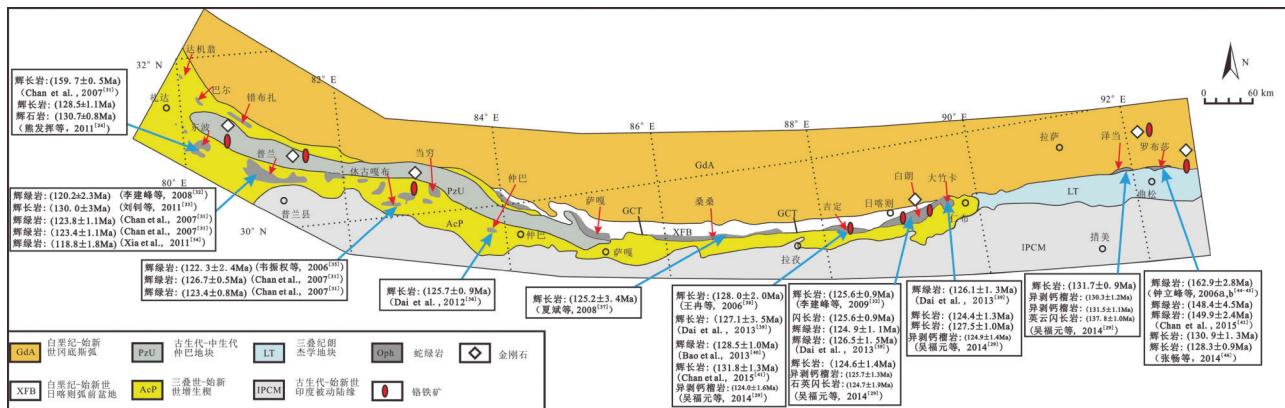


图9 雅鲁藏布江缝合带主要蛇绿岩岩体的地质简图

Fig.9 Age, chromitite and diamonds from the main ophiolite in the YZSZ

基性岩脉所代表雅鲁藏布江缝合带蛇绿岩的时代在各段岩体中存在差异性(图9),年龄在120~170 Ma^[24,29,32~42]。例如周肃等用Sm-Nd法测定东段罗布莎蛇绿岩中辉长辉绿岩脉斜长石、辉石及两个全岩的内部等时线年龄为(177±31) Ma^[43],钟立峰等^[44~45]报道了东段罗布莎蛇绿岩中辉绿岩锆石的SHRIMPU-Pb年龄为(162.9±2.8) Ma,张畅等^[46]报道罗布莎辉长岩中锆石和榍石的年龄为(128.3±0.9) Ma,其西侧相邻的泽当岩体的年龄为(131.5±1.1) Ma。西段前人报道休古嘎布蛇绿岩中辉绿岩的锆石SHRIMPU-Pb年龄(122.3±2.4) Ma^[35],辉长辉绿岩的全岩和矿物等时线年龄为(126±1.5) Ma^[47],获得普兰蛇绿岩中辉绿岩岩墙锆石SHRIMPU-Pb年龄(120.2±2.3) Ma^[32],Miller等报道的拉昂错南侧的普兰(文中称为“Yungbwa”)蛇绿岩中拉斑玄武岩的Sm-Nd等时线年龄为(147±25) Ma,⁴⁰Ar-³⁹Ar年龄为(152±33) Ma存在差异^[48],刘钊等获得普兰岩体东侧MOR型辉长岩的年龄为(130±5) Ma^[33]。桑桑蛇绿岩中辉绿岩的年龄为(125.2±3.4) Ma^[37]。东波蛇绿岩中辉石岩与辉长岩LA-ICP-MS的年龄分别为(130±0.5) Ma、(128±1.1) Ma(早白垩世中期)^[24]。因此,总体上雅鲁藏布江缝合带西段-中段年龄稍晚于东段的年龄,代表新特提斯洋形成的不均一性,可以类比于Izu-Bonin岛弧蛇绿岩的特征^[29],当然地幔橄榄岩体中辉长/绿岩所代表的构造意义还值得探讨。总结雅江蛇绿岩中各段岩体的基性岩脉的地球化学特征,可知罗布莎、普兰和东波岩体形成于洋中脊环境,具N-MORB的亲缘特

征,泽当、日喀则和休古嘎布岩体具有较复杂的洋中脊和岛弧叠加等特点(图10~图11),这对应于前人认为雅鲁藏布蛇绿岩是形成于扩张中心的大洋岩石圈的残留^[10,28,49~53],也可能形成于洋内俯冲带之上的岛弧-弧后盆地等环境^[27,54~59]。

5 雅江蛇绿岩中火山岩特征

该类岩石在区内分布较广泛,有块状、枕状、杏仁状等多种类型之分。它主要出现在雅鲁藏布蛇绿岩各岩体的南北两侧,这些玄武质岩石主要为形成于洋中脊环境下的MORB型,但也存在大洋内部的洋岛型(OIB)以及与俯冲作用有关的岛弧型和大陆边缘型等^[60~63],表明该岩石形成的复杂性。从空间分布来看,不同岩体其分段特征不同,例如日喀则南侧混杂岩带中较早期的玄武岩岩块较多显示OIB的特征,而北侧覆盖在超基性岩之上较年轻的玄武岩主要为MORB型^[64]。总结雅江蛇绿岩中各段岩体出露的火山岩地球化学特征,在Hf/3-Th-Ta图上表明罗布莎玄武岩和休古嘎布岩体的火山岩形成于洋中脊环境或弧后盆地的环境^[31,65~66],具N-MORB的亲缘特征,而东波和日喀则岩体具有较复杂的洋岛等特点,泽当和普兰具岛弧的背景(图12)^[27,60,67~69]。这些不同构造背景的岩石单元被认为是后期增生拼接在一起或者形成于复杂的岛弧-弧后盆地环境。但最近在对美国海岸山脉蛇绿岩的研究揭示了在俯冲带之上的环境中,由于俯冲板片的断离和洋脊的俯冲,会导致在弧前环境形成MORB型和OIB型的火山岩和脉岩^[70~72],该演化过

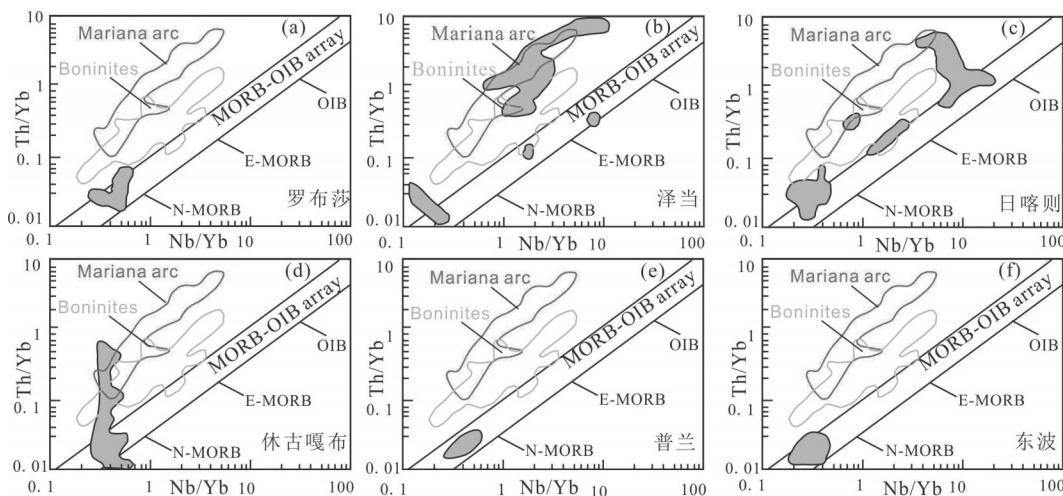


图10 雅鲁藏布江缝合带蛇绿岩中基性岩脉的Th/Yb–Nb/Yb构造背景图解
Fig.10 Th/Yb versus Nb/Yb diagram of tectonic setting from mafic rocks of the main ophiolite in the YZSZ

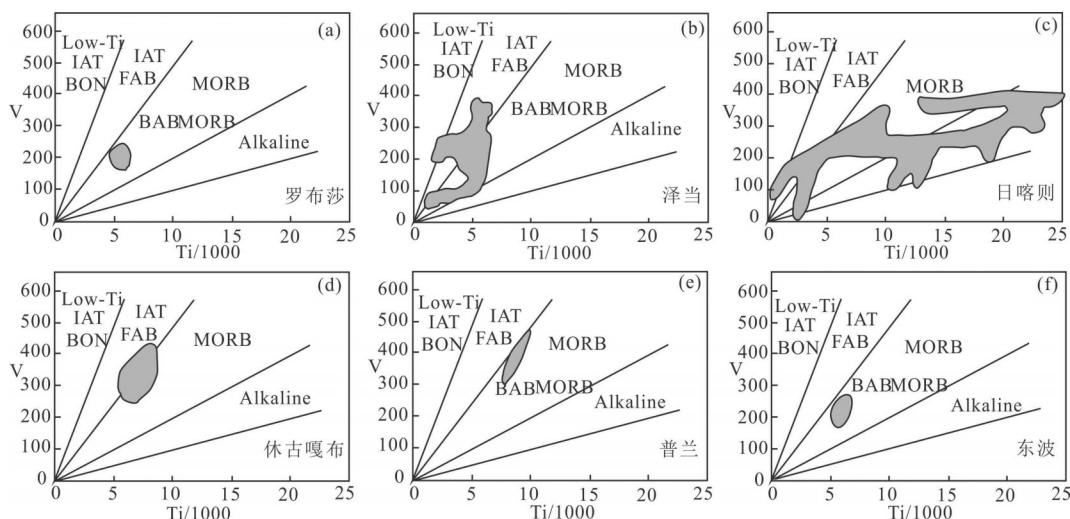


图11 雅鲁藏布江缝合带蛇绿岩中基性岩脉的V–Ti/1000构造背景图解
Fig.11 V–Ti/1000 diagram of tectonic setting from mafic rocks of the main ophiolite in the YZSZ

程在西特提斯洋也得到了很好的验证^[73–76]。

6 雅江蛇绿岩中铬铁矿特征

雅鲁藏布江蛇绿岩带是国内铬铁矿出露点最多,也是铬铁矿石储量、产量最大的岩带,占国产铬铁矿石的90%以上,其中罗布莎铬铁矿床是国内目前唯一的大型铬铁矿床。罗布莎铬铁矿体的分布十分广泛,截止目前共发现600个大小不等、形态各异的铬铁矿矿体,各岩相带中均有铬铁矿体产出,但分布特征不同。第一类是分布于矿区北部纯橄

岩岩相带中。除东西向的矿体外,还有北东、北西向分布的矿体,以各种浸染状矿石为主,局部有致密块状矿石。铬尖晶石颗粒较细,一般1~2 cm,自形晶-半自形晶,与围岩为逐渐过渡关系。除个别矿体规模较大,大多为条带状或不规则透镜状的小矿体,且由于品位较低,分布不集中,一般不具有工业价值。第二类是分布于中部和南部方辉橄榄岩夹纯橄岩带中的矿体,矿体主要以岩体走向相似的东西向、近东西向为主。另有少量的为北东向,矿体总体向南倾,与蛇绿岩构造岩片的倾向一

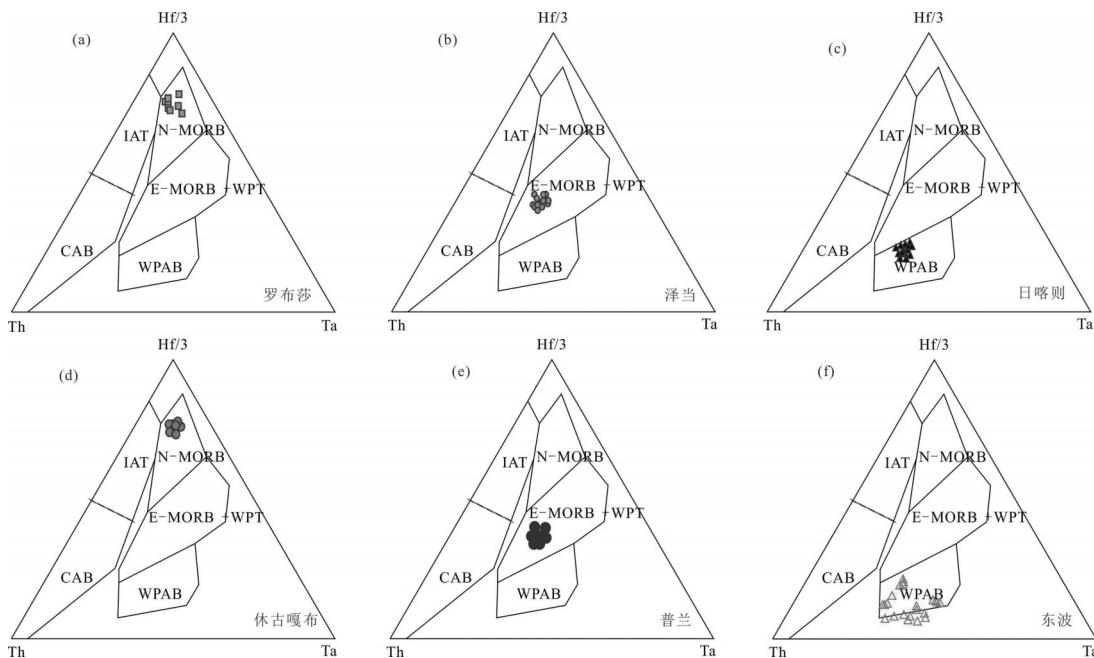


图12 雅鲁藏布江缝合带蛇绿岩中火山岩构造环境判别图解(据参考文献[76])

MORB/OFB—洋中脊玄武岩;WPB—板内玄武岩;LKT—岛弧低钾拉斑玄武岩;CAB—岛弧碱性玄武岩;SH—橄榄粗安岩;
TH—拉斑玄武岩;TR—过渡性;IAB/IAT—岛弧拉斑玄武岩

Fig. 12 Tectonic environment discrimination diagrams for volcanic rocks from the main ophiolite in the YZSZ (after reference [77])
WPB—Intraplate basalts; LKT—Low-K arc tholeiite; CAB—Calc-alkali basalt; SH—Mugearite; TH—Tholeiite; TR—Transition basalt;
IAB/IAT—Island-arc tholeiites

致,沿倾向方向有较大的延伸,多个矿体以雁行式(迭瓦式)排列为特征。以致密块状矿石为主,这类矿体也是罗布莎铬铁矿主要采集的工业矿体^[25]。

由于地质研究和勘查程度不同,雅鲁藏布江蛇绿岩带内已发现的铬铁矿规模和分布极不均衡。通过综合前人资料及野外考察按区段简要介绍带内铬铁矿床点的空间分布,雅鲁藏布江蛇绿岩带东段铬铁矿床数量最多、规模最大,除了罗布莎大型矿床外,还有泽当岩体鲁沟中的铬铁矿床等;而中段虽然岩体数量最多、规模最大,但发现的铬铁矿床数量较少,规模也不大。主要出露在日喀则市到昂仁县之间岩体内,以及仁布岩体内等;西段北亚带的蛇绿岩体规模小,数量少,在达机翁岩体、措布扎岩体和日康巴岩体内均发现铬铁矿(图13),它们空间分布上受北西—南东向断裂带控制,呈透镜状、似透镜状断续产出^[25]。西段南亚带的蛇绿岩体规模大,目前已在拉昂错、东波、当穷、休古嘎布岩体(群)地表发现了大量的铬铁矿化点,为铬铁矿的研究与勘探提供了背景。其中休古嘎布岩体群共发现原生铬铁矿体18个^[21](图8),多分布于岩体群

的南部,距边界围岩一般不超过1 km。有分段集中、成群出现的特点,矿体形态为透镜状、长透镜状、不规则状,赋存于纯橄榄岩异离体内,或直接产于方辉橄榄岩内,多数矿体与围岩接触界线截然清楚,边界呈波状弯曲,少数为渐变过渡。铬铁矿矿石主要为致密块状,仅在部分矿体的边部出现稠密浸染状。

东波蛇绿岩中存在较多矿化点,其中黄圭成等报道已经发现12个铬铁矿矿化点,原生矿点3个,转石矿点8个^[78]。经本团队近三年来的野外工作,新增原位矿点4个,转石矿点1个。矿体主要集中在东波岩体的南侧,一般长1~3 m,厚0.2~2 m,其中高铝型铬铁矿的规模较小为宽约20 cm,长约1 m。矿体的延伸方向为北西向,与岩体展布的方向一致。本文列出了其中5个矿体并分别编号。东波铬铁矿矿体具有类似于普兰铬铁矿的特征,已发现的铬铁矿矿体呈透镜状分布于方辉橄榄岩与纯橄榄壳中,并与方辉橄榄岩/纯橄榄岩呈截然的突变关系,其中东波高铝型铬铁矿的围岩多为方辉橄榄岩,高铬型铬铁矿的围岩为纯橄榄岩。其中DCr-1及DCr-2为致密铬铁矿,DCr-3、DCr-4和DCr-5矿体为



图13 雅鲁藏布江缝合带西段几个岩体中铬铁矿野外照片

a—措布扎铬铁矿;b—巴尔铬铁矿;c—达机翁铬铁矿

Fig. 13 Outcrop of the different ophiolites in western YZSZ

稠密浸染状到中等稠密浸染状。DCr-3、DCr-4 和 DCr-5 的 Al_2O_3 (29%~56%) , Cr_2O_3 (< 40%), $\text{Cr}^{\#}$ =11~47(< 60) 为高铝型铬铁矿, Cr-1 和 Cr-2 矿体为 $\text{Cr}^{\#}$ > 60 属高铬型铬铁矿^[25,79~80]。

前人报道过普兰蛇绿岩中存在较多矿化点, 并认为普兰岩体内的豆荚状铬铁矿都为高铬型铬铁矿, 然而经本团队近五年的野外工作, 新发现原位矿点 7 个, 其中 4 个为致密块状矿体, 3 个稠密浸染状矿体^[25], 并且存在高铬和高铝型不同类型。矿体主要集中在普兰岩体的北侧, 姜叶马一带(图 7)。矿体一般长 2~6 m, 厚 0.5~2 m, 最大者长 10 m, 宽 6 m。矿体的延伸方向为北西向, 与岩体展布的方向一致。已发现的矿体均呈透镜状分布于方辉橄榄岩中, 并与方辉橄榄岩呈截然的突变关系, 无蚀变带及纯橄榄岩外壳。不同矿体的矿石中 $\text{Cr}^{\#}$ 不同, 变化在 52~88, 其中 PCr-1、PCr-2、PCr-3 矿体为致密块状矿体, PCr-4 中矿体的边部为稠密浸染状矿石 $\text{PCr-4}^{\#}=88$, 内部为致密块状矿石 $\text{PCr-4}^{\#}=68$, PCr-5、6 为稠密浸染状矿体。两类铬铁矿的矿石都有致密块状和浸染状, 在致密块状铬铁矿中由 > 95% 的自形至他形的铬尖晶石组成, 脉石矿物主要为蛇纹石和橄榄石, 而浸染状铬铁矿中有 > 30% 的脉石矿物。其中 PCr-4 中稀疏浸染状矿石中的铬尖晶石 $\text{Cr}^{\#}$ 要高于致密块状矿石, 而 PCr-3 及 PCr-5 致密块状又要高于 PCr-2 和 PCr-4 致密块状铬铁矿[。]

PCr-1 及 PCr-6 的 Al_2O_3 (22%~29%), Cr_2O_3 (< 47%), $\text{Cr}^{\#}$ =52~55(< 60) 为高铝型铬铁矿, 而其他矿体为 $\text{Cr}^{\#}> 60$ 属高铬型铬铁矿^[25,79~80]。

铬尖晶石的化学组成可以作为其寄主岩石——地幔橄榄岩的成因指示标志, 这是因为尖晶石的形成受地幔橄榄岩熔融程度、构造背景及形成压力的控制^[81]。总结和对比雅鲁藏布江缝合带主要蛇绿岩岩体中地幔橄榄岩和铬铁矿的铬尖晶石特征, 可知地幔橄榄岩中铬尖晶石的 $\text{Cr}^{\#}$ 值与 $\text{Mg}^{\#}$ 值呈负相关关系(图 15), 具有与世界上大多数阿尔卑斯型超镁铁岩共有的特征, 即 $\text{Cr}^{\#}$ 值随着 $\text{Mg}^{\#}$ 值的升高而降低^[82]。其中部分橄榄岩可以类比于深海橄榄岩, 而部分却与岛弧玄武岩环境($\text{Cr}^{\#}> 60$)较接近。雅鲁藏布江缝合带岩体中基本都存在铬铁矿, 例如罗布莎铬铁矿矿床和普兰、东波铬铁矿^[25], 其铬尖晶石的矿物化学组分差异性和相似性明显。类比于前人对不同构造背景下铬尖晶石的形成环境总结, 可知雅鲁藏布江缝合带地幔橄榄岩和铬铁矿既有深海橄榄岩的特征, 也存在岛弧型橄榄岩的过程^[25,27,57,64~65,85], 并且雅鲁藏布江缝合带西段的岩体更趋复杂。

7 讨 论

7.1 雅鲁藏布江缝合带蛇绿岩组合对比

在近几年野外调查和总结前人的资料中^[17,20,24~25,29,33,44~45,48,51,56,63~65,87~92]可以看出雅江蛇绿岩具



图14 普兰和东波岩体中铬铁矿野外照片(据文献[25])
Fig. 14 Field photos of the Purang and Dongbo chromitite (after reference [25])

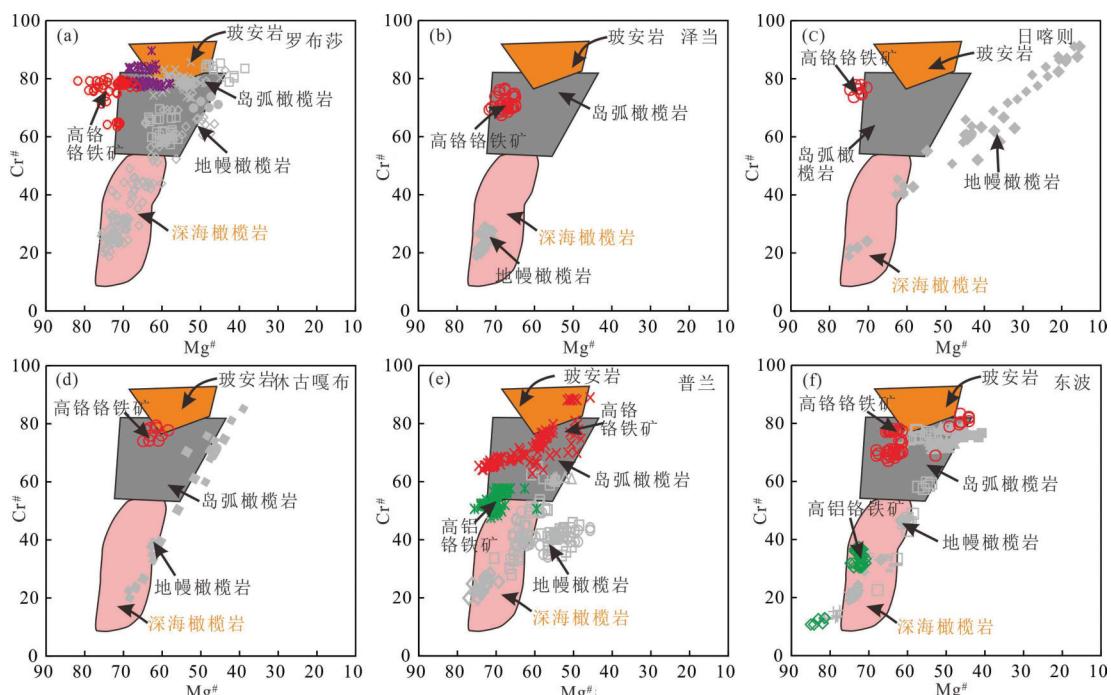


图15 雅鲁藏布江缝合带主要岩体中橄榄岩和铬铁矿的铬尖晶石Cr[#]-Mg[#]图解(据文献[83-84]);
其中数据来源于参考文献[25,27,57,64-65,85]
Fig. 15 Mg[#] versus Cr[#] diagram of chrome spinel in the peridotites from the main ophiolite in the YZSZ

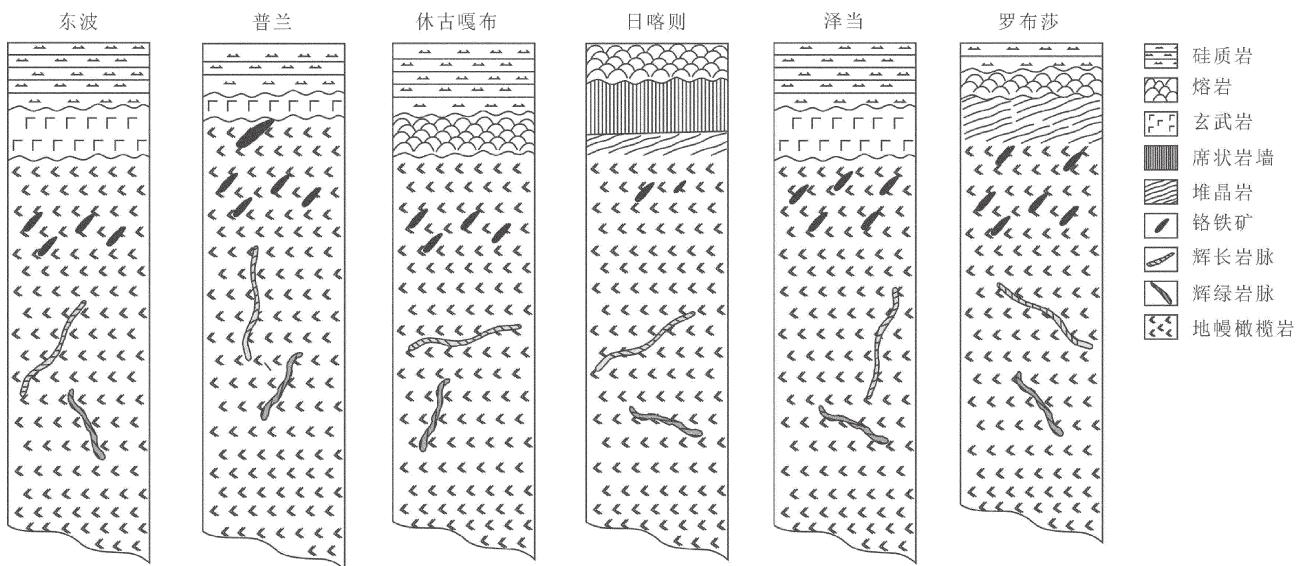


图 16 雅鲁藏布江缝合带蛇绿岩中主要岩体剖面对比图

Fig. 16 Comparison of sequence from the main ophiolite in the YZSZ

有以下几点显著的特征(图16):(1)雅江蛇绿岩中各岩体都以超基性岩为主,辉长辉绿岩分布较少;(2)橄榄岩体大多发生了强烈的蛇纹石化作用,如萨嘎和日喀则等岩体,特别是岩体的边界和断层部分,蛇纹石化作用明显,但罗布莎的康金拉、泽当岩体以及普兰和东波岩体存在大量新鲜的橄榄岩,以方辉橄榄岩和二辉橄榄岩为主;(3)雅江蛇绿岩中较少有席状岩墙群,除报道的日喀则蛇绿岩中顺层侵入的辉长岩辉绿岩岩席,在其他岩体中并未有此现象。但罗布莎岩体存在较厚的堆积辉长辉绿岩于地幔橄榄岩之上。此外,雅江蛇绿岩岩体中都可见辉长辉绿岩脉侵入于橄榄岩中;(4)在这些蛇绿岩的上部发育有一定厚度的玄武岩,如日喀则蛇绿岩。玄武岩与橄榄岩之间经常被辉长岩所占据,部分情况下直接与橄榄岩接触,如普兰蛇绿岩。此外,在部分剖面,可见玄武岩中经常出现硅质岩的夹层,如东波蛇绿岩;(5)豆荚状铬铁矿被发现于各个岩体内部,并常与纯橄榄岩伴生出现,目前罗布莎铬铁矿石储量达700万t,其余岩体也存在找矿的前景;(6)深部地幔矿物普遍存在于雅鲁藏布江缝合带蛇绿岩中,表明其存在深部地幔的演化阶段,记录了超高压强还原环境的痕迹。

雅江橄榄岩地球化学的原始地幔^[93]标准化稀土元素配分型式显示极为一致的“V”或“U”型^[27,55,59],

^{63–65,94]}(图17),LREE均为富集型,存在变化范围较大的REE丰度,可能是由于橄榄岩的熔融程度和亏损程度的差异所致^[95]。与阿尔卑斯橄榄岩的轻稀土亏损型模式明显不同^[96],地幔橄榄岩可能不只是早期的地幔交代作用所致,反映了岩石可能还经历了俯冲洋壳流体的改造。并且在泽当和普兰、东波岩体中二辉橄榄岩显示了阿尔卑斯型轻稀土亏损的特点,表明其与阿尔卑斯二辉橄榄岩以及现代洋底橄榄岩的REE丰度及其LREE亏损型较为接近^[97],反映原始地幔橄榄岩的部分熔融特征形成于MOR环境,富集的LREE指示其SSZ环境改造的特点。

7.2 雅鲁藏布江缝合带铬铁矿的前景讨论

世界上著名的大型豆荚状铬铁矿床的特征指示大多数铬铁矿床产在规模较大的蛇绿岩体内,如中生代古巴的Camaguey岩体为1200 km²,菲律宾的Acoje岩体大于4500 km²,阿尔巴尼亚的Bulqiza岩体面积为350 km²,土耳其的Guleman–Soridag岩体达200 km²。古生代俄罗斯的Ray–IZ岩体400 km²,Saranov岩体为1950 km²,哈萨克斯坦的Kempirsai岩体为1000 km²。而普兰地幔橄榄岩接近600~700 km²,类似的雅鲁藏布江缝合带西段几个大岩体,还包括东波岩体400 km²和休古嘎布岩体700 km²。另外岩体的部分熔融程度的高低也是成矿的一个重要的标注,豆荚状铬铁矿的产出多出现在方辉橄榄

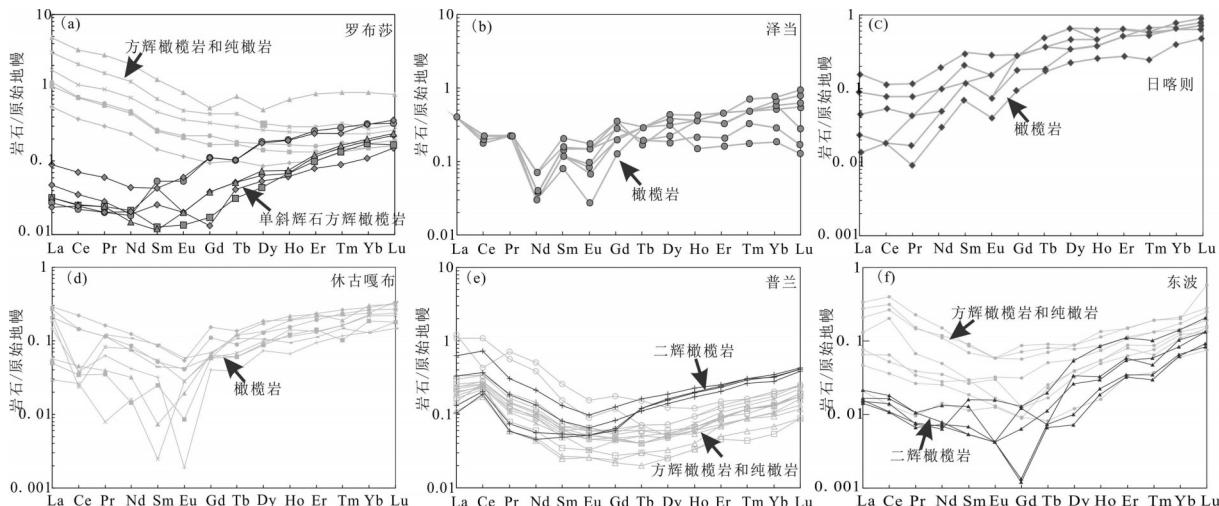


图 17 雅鲁藏布江缝合带主要岩体的全岩REE模式图(据文献[93])

Fig. 17 Primitive mantle normalized REE patterns from the main ophiolite in the YZSZ(after reference[93])

岩型蛇绿岩中,而基本上不存在于二辉橄榄岩内^[51],这是由于铬主要赋存在辉石内,部分熔融的增加,铬才能被大量释放而聚集成矿,地幔橄榄岩经历不同构造背景的形成过程,其辉石大量的被熔融,铬铁矿也可能经历多期次叠加,以至于富集大型的铬铁矿矿床;依据世界上豆英状铬铁矿的产出规律可知其主要是位于堆积杂岩底部的橄榄岩中和MOHO面以下的橄榄岩中。例如阿曼铬铁矿、菲律宾的阿科杰铬铁矿都产于堆积杂岩之下的1~2 km的纯橄榄岩-方辉橄榄岩的过渡带,两类铬铁矿同时存在的土耳其Mugla岩体中高铝型铬铁矿位于含单斜辉石的方辉橄榄岩岩相内,而高铬型类位于MOHO面以下的亏损地幔橄榄岩中。古巴Sagua铬铁矿中两类铬铁矿都位于方辉橄榄岩-纯橄榄岩相带内,是不同期次不同性质的熔体与岩石反应的结果。阿尔巴尼亚Bulqiza铬铁矿呈现最为明显的特征,其高铝型铬铁矿处于堆晶纯橄榄岩相内,而高铬型铬铁矿位于岩体下层部位的方辉橄榄岩-纯橄榄岩相中;除上述规律外,豆英状铬铁矿中致密块状铬铁矿矿体多与方辉橄榄岩之间呈截然的接触关系,偶见宽度在0.5~1 m的蛇纹岩带,而浸染状和豆英状铬铁矿矿体多与纯橄榄岩壳呈渐变过渡的关系,这是由于铬铁矿的多期形成过程中,较完整的痕迹都是最晚期的特征。豆英状铬铁矿的最大延伸方向常与岩体展布方向,以及单斜辉石的叶理方向一致,这是由于洋中脊之下的地幔软流圈向上运移的

过程^[52],造成大量的透镜状纯橄榄岩、方辉橄榄岩和辉石岩的变形所致。罗布莎岩体是个倒转岩体,整个岩体的辉石含量较低,部分熔融程度较高,少见二辉橄榄岩。而阿里大量分布的超基性岩体,其岩相学,地球化学,矿体特征都可类似比于世界典型铬铁矿床的岩体特征,并已发现大量的铬铁矿野外露头和转石矿点,存在较大的找矿空间。

世界上目前仅在少数岩体中报道了同一岩体内出现两种类型的豆英状铬铁矿,其中有古巴的Mayari-Cristal超基性岩和阿尔巴尼亚的Bulqiza岩体。中国的几个主要铬铁矿床均为单一型,其中罗布莎、东巧、玉石沟等为高铬型铬铁矿;而贺根山、萨尔托海等为高铝型铬铁矿。普兰和东波地幔橄榄岩中均发现了高铬型和高铝型两类铬铁矿在国内属首次。经国内外对比研究,将两类铬铁矿的特征归纳为以下几点:①两类铬铁矿的铬尖晶石矿物化学组分特征不同。高铝型铬尖晶石的Cr#=11~55、Mg#=60~85,而高铬型铬铁矿的铬尖晶石为Cr#=65~80、Mg#=40~65。两者位于不同的演化区间,例如,东波高铝型铬铁矿落在橄榄岩演化趋势上,其Cr#-Mg#呈负相关性(图18),而高铬型铬铁矿落在Mg#和Cr#的正相关趋势线上;②无论是高铬型还是高铝型铬铁矿,其围岩不一定存在纯橄榄岩壳^[25]。罗布莎高铬型铬铁矿的野外特征显示,规模较大的致密块状铬铁矿围岩更多的为方辉橄榄岩,表明部分铬铁矿的形成可能并不是按岩石-熔体反应的观点所描

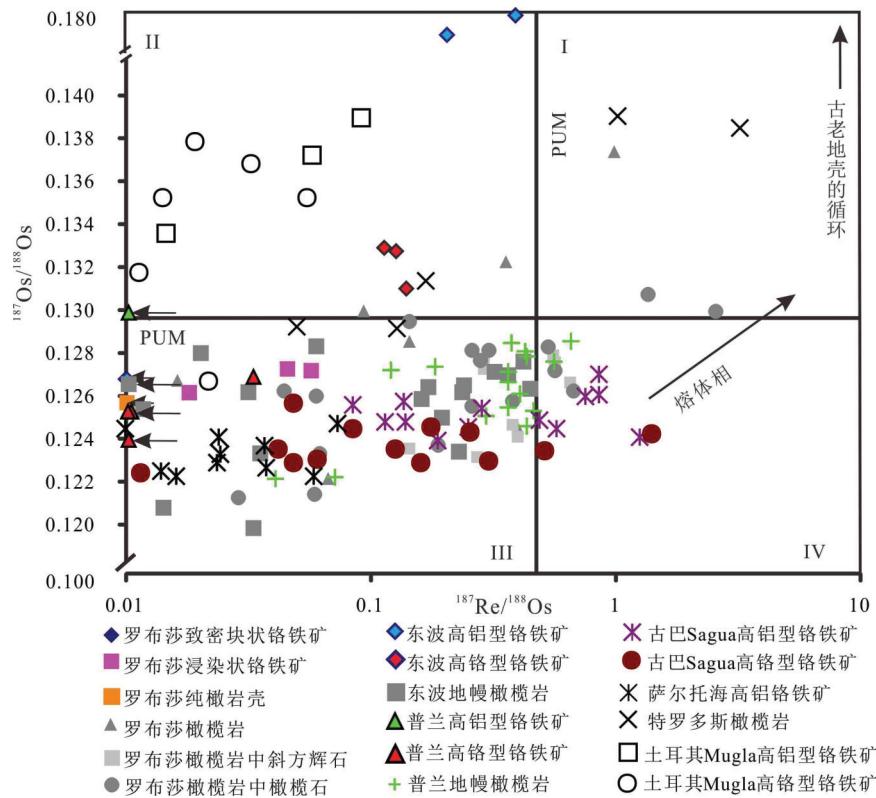


图18部分铬铁矿矿床中两类铬铁矿及其地幔橄榄岩的 $^{187}\text{Re}/^{188}\text{Os}$ - $^{187}\text{Os}/^{188}\text{Os}$ 图解

其中罗布莎橄榄岩、斜方辉石和橄榄石的Re-Os数据来自[98],普兰地幔橄榄岩来源于Liu et al^[89],东波地幔橄榄岩来源于Niu 等^[99],古巴铬铁矿来源于González-Jiménez 等^[100],萨尔托海铬铁矿来源于Shi 等^[101],特罗多斯橄榄岩来源于Büchl 等^[102],土耳其铬铁矿来源于Uysal 等^[103]

Fig. 18 $^{187}\text{Re}/^{188}\text{Os}$ versus $^{187}\text{Os}/^{188}\text{Os}$ diagram of two types of chromitite and their mantle peridotite

述的那样形成矿体与纯橄岩壳^[85]。罗布莎铬铁矿和围岩的矿物学特征显示致密块状铬铁矿矿体与其围岩特征并不相同,也不具渐变的过程,普兰岩体中高铬型致密块状铬铁矿进一步佐证了这一特征;③浸染状铬铁矿常与纯橄岩壳伴生,并呈现两者之间渐变连续的过程。该特征在罗布莎高铬型浸染状铬铁矿尤为明显,纯橄岩壳-浸染状铬铁矿之间表现为渐变过渡关系,橄榄石与铬尖晶石的矿物组分变化指示岩石-熔体反应的特征^[85],而东波高铝型浸染状铬铁矿及其围岩纯橄岩也具此特点;④不同岩体的高铬型铬铁矿的围岩通常显示轻稀土元素富集,认为是后期流体改造特征;东波高铝型铬铁矿的围岩显示轻稀土元素亏损特征,与岩体中呈透镜状的二辉橄榄岩特征相同,属亏损地幔特征(例如罗布莎、普兰及东波、萨尔托海等)。

地幔橄榄岩和铬铁矿的Re-Os特征表明除特罗多斯橄榄岩具有典型壳源物质混染外,其余岩体

的地幔橄榄岩较多的位于第三象限小于PUM值(0.1296)(图18),与实验证实的深海地幔橄榄岩的比值范围0.120~0.129一致。雅江西段两类铬铁矿的Re-Os同位素也具有此特征,它们的 $^{187}\text{Re}/^{188}\text{Os}$ 、 $^{187}\text{Os}/^{188}\text{Os}$ 的值区间基本相似,其中罗布莎高铬型铬铁矿与普兰高铬型铬铁矿、古巴Sagua高铬型铬铁矿比值接近,都位于第三象限的残余相内;东波高铬型铬铁矿与土耳其Mugla高铬型铬铁矿都大于0.1296的值。此外,高铝型铬铁矿也可分为两组,其中古巴Sagua和萨尔托海高铝型铬铁矿 $^{187}\text{Os}/^{188}\text{Os}$ 值较低,与其地幔橄榄岩相同;另外一组是东波、普兰以及土耳其Mugla高铝型铬铁矿明显大于0.1296,东波个别高铝型铬铁矿甚至达到0.180,表明地壳物质混染的现象明显,指示此类铬铁矿的形成过程有地壳流体的加入^[25]。因此,两类铬铁矿Re-Os同位素特征上有较大的差别,显示来源于不同岩浆的演化过程。尤其是东波岩体和普兰岩体中的两类铬铁矿,不仅仅有

含富集IPGE的铂族矿物(PGM),还存在富集PPGE的PGM矿物,进一步佐证了这一特征。

豆英状铬铁矿的形成过程与蛇绿岩的形成过程紧密相连。近些年的研究认为高铬型铬铁矿形成于SSZ^[104-111],并得到了一些实验岩石学的支持,例如原始橄榄石和铬尖晶石饱和的熔体在水富集的条件下,能形成蛇绿岩型铬铁矿^[11],而这正是SSZ环境所具有的特征。在研究西藏罗布莎高铬型铬铁矿和新疆萨尔托海高铝型铬铁矿后,Zhou等提出了高铬型铬铁矿形成于俯冲带(SSZ)构造背景下^[109],而高铝型铬铁矿形成在弧后盆地或者洋中脊的环境,并提出两类铬铁矿是由两种不同的熔体与岩石(方辉橄榄岩)反应形成,其中高铬型铬铁矿是俯冲带富水的玻安质熔体,而高铝型铬铁矿是在洋中脊/弧后盆地由拉斑玄武质熔体与岩石(方辉橄榄岩)反应而成。Malpas等在研究Sagua高铝型和高铬型铬铁矿后,认为它们形成于弧前的环境^[113]。Gerville等研究了同一缝合带上相邻的Mayari岩体、Sagua岩体和Moa-Baracoa岩体^[114],指出两岩体形成在一个“过渡性”的大地构造环境内,从“早期岛弧”演化成弧后盆地,分多阶段形成。俯冲的第一阶段为,Mayari高铬型铬铁矿形成于“早期岛弧”玻安质岩浆的反应;第二阶段,Sagua高铬和高铝型铬铁矿的岩体形成于“过渡性”大地构造背景下,即从“早期的岛弧”到新生的弧后盆地的阶段。Moa-Baracoa富铝型的铬铁矿形成于最后,即弧后盆地阶段。土耳其的Mugla两类铬铁矿中被解释为两期,其中第一期高铝型铬铁矿形成初期洋中脊环境,第二期高铬型的铬铁矿在俯冲带形成^[115],与岩体伴生的火山岩提供了这类构造背景的佐证^[116]。

雅鲁藏布江缝合带中不同蛇绿岩显示出多成因、多阶段演化的特征。综合前人的资料,雅鲁藏布江缝合带存在两种类型的蛇绿岩:①形成于典型的SSZ环境的弧前盆地,这种类型的蛇绿岩主要发育在雅鲁藏布江缝合带中段,包括日喀则蛇绿岩体^[60]、萨嘎和桑桑蛇绿岩^[55]。②地幔橄榄岩形成于MOR环境,但是后期普遍都受到了不同程度SSZ环境的改造,这种类型的地幔橄榄岩主要位于雅鲁藏布江缝合带的东段和西段,包括东段罗布莎蛇绿岩体的地幔橄榄岩^[117]。西段的东波和普兰两岩体的地幔橄榄岩矿物学、岩石学等特征都显示了MOR→

SSZ的过程,两岩体中的辉长岩/辉绿岩-玄武岩也反映类似结果^[24,48,90-91]。

在总结罗布莎、普兰和东波铬铁矿和地幔橄榄岩的特征,及对比其他地幔橄榄岩和铬铁矿的基础上,初步认为豆英状铬铁矿与地幔橄榄岩的形成过程是经历了不同期次的结果,提出多阶段的成因模式^[16,25,65,85]:早期俯冲的陆壳和洋壳物质到达地幔过渡带后成分被分解,岩石脱水,产生大量流体,地壳中放射性物质生热,导致地幔熔融,Cr等成矿元素从矿物中释放并聚集,在地幔柱作用下熔融物质上涌,在过渡带顶部,熔体开始结晶,铬铁矿聚集成矿,独立的携带微粒金刚石等超高压矿物的强还原性流体进入开始固化的铬铁矿和围岩,呈半固化的铬铁矿和熔融体继续向上运移,而这一阶段的形成时代可能大于300 Ma,数据基于McGowan等获得的290~630 Ma的年龄和铬铁矿中选出的大量锆石的U-Pb年龄为(376±7) Ma^[118],以及Robinson等获得大于500 Ma的锆石U-Pb年龄^[119];第二阶段在向上运移至洋中脊的过程中,压力不断降低,超高压相矿物转变成低压相矿物,溶解在铬铁矿矿浆中物质出溶成柯石英和单斜辉石等,铬铁矿矿浆进一步聚集形成致密块状铬铁矿,此阶段的时代主要限定来源于洋中脊环境下形成的各种基性岩脉,其年龄变化于120~170 Ma^[24,29,32-42];在俯冲环境下,含水熔体与方辉橄榄岩的反应,并普遍改造了地幔橄榄岩的地球化学特征,提高了橄榄岩的部分熔融程度,而其发生蛇绿岩仰冲在88 Ma^[113]。以罗布莎岩体为例,其具有早阶段MORB型地幔和辉长岩墙Sm-Nd等时线年龄为(177±33) Ma^[88],晚期为俯冲带(SSZ)性质蛇绿岩的锆石SHRIMP年龄为(126±2) Ma^[113]。并且罗布莎地幔橄榄岩的矿物学,岩石地球化学性质指示,岩体开始形成于MOR阶段,后受不同程度的SSZ环境改造^[120],同样的现象出现于雅江西段的岩体内^[121-122]。然而在该岩体的地幔橄榄岩和铬铁矿中发现的地幔超高压矿物,表明可能还有更早一期阶段来源于地幔深部过程,深度可能达到>400 km^[123-129]。因此,雅鲁藏布缝合带蛇绿岩及铬铁矿可能普遍经历了多阶段的演化过程,每一阶段对铬铁矿富集程度及贡献还有待进一步的研究,每个岩体其经历的部分熔融程度及俯冲流体改造的规模尚需明确。西段的几个大型岩体与罗布

莎存在较多相似之处,均经历了多阶段的构造背景演化和成矿过程,存在较大的找矿潜力。

8 结语

经典的板块构造理论认为蛇绿岩和其中的豆状铬铁矿,无论是产在MOR或SSZ环境,均是形成在软流圈之上的上地幔顶部。地幔中异常矿物和超高压矿物的发现是对上述传统理念提出挑战,需重新审视蛇绿岩和铬铁矿的成因。多年来人们通过金伯利岩金刚石中的包裹体矿物探讨深部地幔的物质组成、地幔的物质循环和动力学机制。尤其是近些年地幔柱和深部地幔对流的理念越来越引起人们的重视,地幔物质组成和地幔动力学成为固体地球科学研究的前沿。蛇绿岩铬铁矿和地幔橄榄岩中金刚石等超高压矿物的发现引起国际同行的高度关注,已经成为探讨深部地幔作用一个新的窗口和研究方向,属国际上研究热点和前沿问题。雅鲁藏布江缝合带蛇绿岩剖面完整连续,各类地质体发育齐全,是研究蛇绿岩及铬铁矿的理想场所,也是中国铬铁矿勘探的潜力所在,其科学意义和可探索的空间,必将会成为更多科技工作者的热土。

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