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赣东北朱溪矿集区构造控岩-控矿特征

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摘要:赣东北朱溪为钦杭成矿带新发现的重要钨铜多金属矿集区, 位于钦杭结合带萍(乡)乐(平)拗陷带东部逆冲推覆抬升地段。朱溪矿集区燕山期受逆冲推覆构造影响, 形成由新元古界变质基底与上古生界-中生界沉积盖层组成的构造岩片堆叠构造; 具有逆冲推覆深断裂带控岩-控矿、碳酸盐岩构造岩片赋矿、燕山晚期浅层对冲构造破矿的构造背景; 发育有燕山早期 I 型花岗闪长(斑)岩钨铜和 S 型花岗岩钨铜两个岩浆岩成矿系列, 在空间上形成张家坞一月形与塔前一朱溪两个矿田、张家坞—毛家园和塔前一朱溪上下两个成岩台阶, 下成矿台阶朱溪巨大钨铜矿床的发现打开了钦杭成矿带拗陷区“深地”找矿的一扇窗户。

关键词:逆冲推覆构造带; I+S 型岩浆成矿系列; 多台阶成矿; 朱溪矿集区

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Rock- and ore-controlling structure in the Zhuxi ore concentration area in the northeastern Jiangxi Province

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Abstract: Zhuxi is a newly discovered W-Cu ore concentration area in the Qinzhou-Hangzhou metallogenic belt, northeastern Jiangxi Province. It is situated in the lifting area of the thrust nappe structure in eastern Pingxiang-Leping depression, Qinzhou-Hangzhou orogenic juncture belt. Due to the Yanshanian thrust nappe tectonic events, the Zhuxi ore concentration area is dominated by Neoproterozoic metamorphic basement and Upper Paleozoic-Mesozoic tectonic slices. The ore deposits in this region are

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controlled by thrust nappe deep faults, hosted in carbonatite tectonic slices, and transformed by superficial hedge structures. The early Yanshanian metallogenesis can be subdivided into two groups: (1) Mo-Cu metallogenic system related to the I-type (porphyritic) granodiorite; (2) W-Cu metallogenic system related to the S-type granite. The Zhuxi area is characterized by two orefields (Taqian-Zhuxi and Zhangjiawu-Yuexing) and two magmatic steps (Maojiayuan-Zhangjiawu and Zhuxi-Taqian). The discovery of the Zhuxi giant W-Cu deposit opens a window for the deep mineral exploration in the geotectogene of the Qinzhou-Hangzhou metallogenic belt.

Key words: thrust nappe structure; I- and S-type magma-mineralization series; multi-step mineralization; Zhuxi ore concentration area

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1 引言

赣东北朱溪钨铜多金属矿集区位于钦杭成矿带中段萍乐拗陷带东部,以景德镇浮梁县朱溪为中心,西起乐平市塔前,东至婺源县赋春。萍乐拗陷带斜贯江西省中部,上叠于晋宁期钦杭结合带和加里东期由新元古界万年群组成的万年逆冲推覆体之上,呈NEE向展布(杨明桂等,1997,2009,2015,2016a,b;刘一等,2016),由上古生界一中侏罗统海陆交互相沉积组成。该拗陷带长期以煤、非金属矿产较为重要,内生金属矿床规模相对较小。21世纪以来,通过深部找矿,朱溪深隐伏世界级钨矿床的发现,引起了广泛关注。早在20世纪90年代初项新葵(1992)就系统论述过朱溪矿集区中新生代构造演化与成矿作用,认为区内构造演化与成矿可分为三个阶段,一是海西-印支期海底火山喷发在石炭-二叠系中形成矿源层,二是燕山期岩浆侵位对矿源层叠加改造形成以铜为主的多金属矿床,三是喜马拉雅早期万年群浅变质岩推滑覆于晚古生代-中生代地层及矿床之上,最终形成隐伏-半隐伏矿床。近些年众多学者围绕朱溪钨铜多金属矿床对该区做了大量科研工作,主要集中在朱溪矿床地质特征(Chen et al., 2016)、成岩-成矿年代学(李岩等, 2014; 万浩章等, 2015; 刘战庆, 2016; Chen et al., 2016; Pan et al., 2017, 2018; Song et al., 2018a, b; 贺晓龙等, 2018)、成矿作用过程(欧阳永棚等, 2018)等方面,刘善宝等(2014)还对朱溪矿集区岩浆岩进行

了系统研究,表明区内至少有3期岩浆活动(850 Ma 花岗闪长斑岩和煌斑岩脉、160 Ma 花岗闪长斑岩和煌斑岩、146 Ma 花岗岩)和2期成矿作用(与160 Ma 岩浆岩有关的塔前钨钼矿床和月形铜多金属矿床、与146 Ma 花岗岩有关的朱溪矽卡岩型铜钨矿床)。此外,亦有部分研究涉及到构造控岩-控矿特征,其中谢涛等(2015)对朱溪矿区构造控矿特征进行了简要论述,认为朱溪矿床的形成受岩体接触带、滑脱构造、层间裂隙-破碎带等多重构造控制;霍海龙等(2018)通过对矿集区中生代推覆构造特征及年代学研究,认为区内与推覆构造同期或稍晚的岩浆作用沿推覆构造界面侵位并在适当的位置成矿。但目前为止,鲜有学者系统总结过矿集区找矿进展及构造控岩-控矿作用。本次研究是在前人已有研究成果的基础上,结合最新勘查成果,对矿集区进行了整体性时空四维结构研究和深层找矿预测。

2 成矿地质背景

研究区位于萍乐拗陷带东部(图1a),七宝山-朱溪逆冲推覆深断裂带东段,包括塔前-赋春、横路-大游山两支深断裂带(图1b)。受九岭逆冲推覆带前缘影响,形成构造岩片堆叠结构,使拗陷带遭到破坏,沉积盖层已碎片化,自北向南依次为毛公尖、景德镇、朱溪、涌山等构造岩片。朱溪矿集区可分为2个构造岩片。北部为朱溪岩片,处于七宝山-朱溪逆冲推覆深断裂带下盘,由塔前-赋春单斜与杨草尖变质岩块组成;南部为涌山岩片,处于

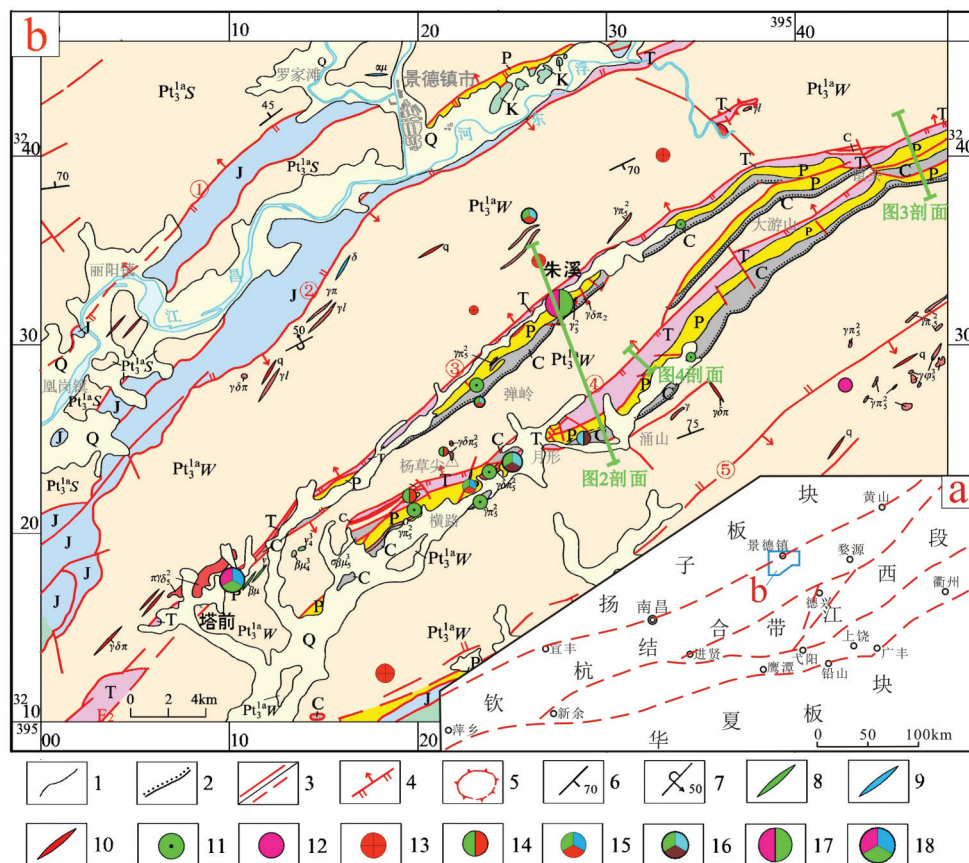


图1 朱溪矿区地质矿产简图(据饶建锋等,2017修改)

1—地质界线;2—角度不整合界线;3—实测/推测断裂;4—逆冲推覆断裂;5—构造窗;6—正常岩层产状;7—倒转岩层产状;8—基性岩脉;9—中性岩脉;10—酸性岩脉;11—铜矿(化)点;12—钨矿(化)点;13—金矿(化)点;14—铜金矿点;15—铜钨金矿点;16—铜铅锌矿点;17—钨铜矿床;18—钨钼铜矿床。Q—第四系;K—白垩系;J—侏罗系;T—三叠系;P—二叠系;C—石炭系;Pt^{1a}W—新元古界万年群;Pt^{1a}S—新元古界双桥山群; $\gamma\phi_3^1$ —燕山晚期钠长花岗岩; $\gamma\delta\pi_2^1$ —燕山早期花岗岩闪长斑岩; $\gamma\pi_2^1$ —燕山早期花岗岩斑岩; γ_2^1 —燕山早期花岗岩; $v_4^3/\beta\mu_4^3$ —华力西晚期辉长岩/辉绿岩; $\gamma\delta\pi_2$ —晋宁期花岗岩闪长斑岩;①—丽阳—罗家滩逆冲推覆断裂带;②—凰岗—湘湖逆冲推覆断裂带;③—塔前—赋春逆冲推覆深断裂带;④—横路—大游山逆冲推覆断裂带;⑤—临港—乐河逆冲推覆断裂带

Fig.1 Geological and mineral resources map of the Zhuxi ore concentration area (modified from Rao et al., 2017)

1—Geological boundary; 2—Angular unconformity boundary; 3—Measured/inferred fracture; 4—Thrust nappe fracture; 5—Tectonic window; 6—Normal strata attitude; 7—Reversed strata attitude; 8—Basic vein; 9—Neutral vein; 10—Acid vein; 11—Copper ore spot; 12—Tungsten ore spot; 13—Gold ore spot; 14—Copper-gold deposit; 15—Copper-molybdenum gold ore; 16—Copper-lead-zinc deposit; 17—Tungsten-copper deposit; 18—Tungsten-molybdenum-copper deposit. Q—Quaternary; K—Cretaceous; J—Jurassic; T—Triassic; P—Permian; C—Carboniferous; Pt^{1a}W—Neoproterozoic Wannian Group; Pt^{1a}S—Neoproterozoic Shuangqiaoshan Group; $\gamma\phi_3^1$ —Late Yanshanian albite granite; $\gamma\delta\pi_2^1$ —Early Yanshanian granodiorite porphyry; $\gamma\pi_2^1$ —Early Yanshanian granite porphyry; γ_2^1 —Early Yanshanian granite; $v_4^3/\beta\mu_4^3$ —Late Hualixi gabbro/diabase; $\gamma\delta\pi_2$ —Jinning period granodiorite porphyry; ①—Liyang—Luojiatan thrust nappe fault zone; ②—Huanggang—Xianghu thrust nappe fault zone; ③—Taqian—Fuchun thrust nappe deep fault zone; ④—Henglu—Dayoushan thrust nappe fault zone; ⑤—Lingang—Lehe thrust nappe fault zone

横路—大游山逆冲推覆断裂带下盘,由横路—涌山单斜与珍珠山变质岩块组成。朱溪、涌山岩片分别控制着塔前—朱溪与张家坞—月形两个矿田。燕山晚期又受到来自南北侧的对冲式推覆构造,对成矿起到破坏作用(黄雷等,2017;杨明桂等,2018),加上燕山造山后伸展断陷,使构造更加复杂化。

2.1 地层

区内变质基底由新元古界万年群组成,为一套浅

变质的含火山凝灰质砂泥质浊积岩,沉积盖层则由上石炭统一上三叠统组成。上石炭统黄龙组沉积不整合于万年群之上,黄龙组底部石英砾岩厚数十厘米至几米不等,下部为白云质灰岩、白云岩,上部为灰岩,为最有利的成矿围岩(Chen et al., 2016; 欧阳永棚等, 2018);二叠系中下统以灰岩为主夹硅泥质岩、镁质黏土,为较有利的成矿围岩;二叠系上统为乐平组煤系与长兴组灰岩;三叠系下统为碳酸盐岩,与上统安源

组煤系之间为角度不整合接触关系。

2.2 岩浆岩

区内自晋宁期以来发生有多次岩浆侵入活动,但以燕山早期中酸性岩浆成矿为主。

(1) 燕山期前侵入岩

晋宁期含铜花岗闪长斑岩:见于朱溪矿区南部上石炭统与万年群接触带附近,花岗闪长斑岩呈岩瘤状产出,直径约15 m,发生斑岩铜矿化,受断裂错动为“断根”的透镜体。锆石U-Pb年龄(847.2±9.4) Ma(万浩章等,2015)。

华力西期辉长辉绿岩脉群:产于汪家—杨草尖一带,呈脉状侵入于万年群中。刘战庆(2016)获得绿泥石化辉绿岩锆石U-Pb年龄(293.5±2.6) Ma,可能为华力西期地壳伸展期岩墙。

(2) 燕山期成矿岩浆岩

为区内最主要成矿岩浆岩。时代为燕山早期,包括晚侏罗世早、晚两个阶段。分为I型中酸性花岗闪长(斑)岩和S型花岗岩两个岩浆岩成矿系列,组成塔前一赋春、横路一大游山南北两条侵入岩带,并形成上下两个成岩台阶。

①塔前一赋春侵入岩带:岩体自西向东分布于塔前一赋春逆冲推覆深断裂带中,主要岩体集中出露于毛家园、弹岭、朱溪3处。

塔前岩体包括南东部的花岗闪长岩墙和北东部的毛家园斑状花岗闪长岩,后者遭受较强烈剥蚀,出露面积约4 km²。斑状花岗闪长岩年龄值为(159.7±1.8) Ma(刘善宝等,2014)、(160.9±2.5) Ma(胡正华等,2015),属晚侏罗世早期,侵入于塔前花岗闪长岩墙。岩体中分布有晚侏罗世晚期花岗斑岩墙,锆石U-Pb年龄(150.8±1.3) Ma(Chen et al., 2016)。

弹岭岩体呈小岩瘤状产出,遭受强烈剥蚀,仅北部有零星露头。以往认为该岩体为花岗斑岩,此次查明其岩性为花岗闪长斑岩。锆石U-Pb年龄(160.2±0.8) Ma(刘战庆,2016),时代为晚侏罗世早期。

朱溪矿区浅部主要为中酸性岩脉,以闪长岩、闪长玢岩、煌斑岩脉为主,还有少量花岗闪长斑岩、花岗斑岩等岩脉。矿区深部标高-400 m以下见隐伏花岗岩前锋带,主岩体隐伏于-1600 m以下,岩性主要为黑云母花岗岩、碱长二云母花岗岩、云英岩化花岗岩,其锆石U-Pb加权平均年龄分别为(146.9±0.97) Ma~(153.5±1) Ma(王先广等,2015;

Song et al., 2018a; Pan et al., 2018)、(149.0±1.0) Ma~(152.9±1.7) Ma (Song et al., 2018a)、(149.2±1.5) Ma~(152.9±1.7) Ma(Chen et al., 2016; Song et al., 2018a; Pan et al., 2018);此外,钻孔中还见有大量花岗斑岩脉,锆石U-Pb加权平均年龄为(148.3±1.4) Ma~(153.4±1) Ma(李岩等,2014; Chen et al., 2015; Song et al., 2018a; Pan et al., 2018),成岩时代均为晚侏罗世末。此外,在该带南侧万年群变质岩中见有杨草尖黑云母花岗闪长斑岩岩瘤,面积0.16 km²。

②横路一大游山侵入岩带:岩体主要分布于横路一大游山逆冲推覆深断裂带下盘,自西向东有横路花岗闪长斑岩墙带、张家坞花岗闪长斑岩株(面积0.63 km²)及月形花岗闪长斑岩墙群。张家坞、月形花岗闪长斑岩锆石U-Pb加权平均年龄分别为(161.6±3) Ma(霍海龙等,2018)、(159.7±2.0) Ma(刘战庆,2016)。

2.3 构造

朱溪矿集区由北部的塔前一赋春单斜、杨草尖变质岩块、涌山构造岩片和横路一大游山单斜组成,且区内金属矿床(点)主要赋存于塔前一赋春单斜和横路一大游山单斜构造内,两单斜构造主要由石炭—三叠纪碳酸盐岩和含煤碎屑岩组成,呈走向北东50°~55°、倾向北西展布,局部形成叠瓦状构造。两个由上石炭统一上三叠统组成的单斜北南两侧分别为塔前一赋春、横路一大游山逆冲推覆导岩导矿的深断裂带。单斜底部黄龙组与万年群之间的角度不整合面往往受后期挤压或滑动改造。上述“两单斜夹一块”、“两断裂”及“两不整合面”构成了矿集区的基本构造框架(图2)。

区内燕山期有多次断裂活动,燕山早期以向SSE逆冲推覆构造为主导,燕山晚期在塔前、月形矿区均有向NW逆冲破矿构造分布,与向SE的逆冲推覆构造形成“对冲”特征(图3),使矿床遭到严重破坏,而且相当强烈复杂。所以深层逆冲推覆深断裂导岩导矿,对冲式断裂破矿,是该矿集区构造的主要特征。

塔前一赋春逆冲推覆深断裂带:断裂带使由万年群组成的牛角岭变质岩片逆冲于塔前一赋春向斜之上,使向斜北翼遭到破坏叠覆,残留下南翼单斜。断裂带走向NE,西部塔前一朱溪段浅部产状陡立,朱溪矿区布设于断裂上盘的ZK4216钻孔至

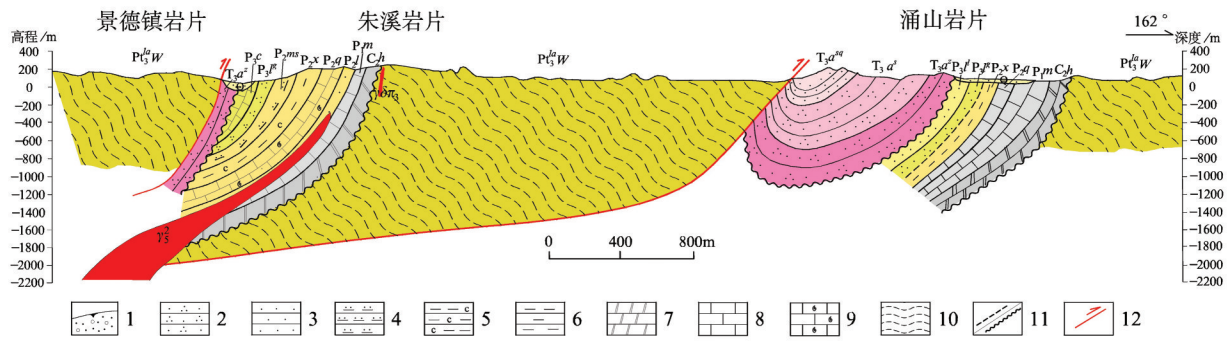


图2 朱溪矿集区构造剖面(据钟南昌, 1992 修改)

1—浮土; 2—石英砂岩; 3—砂岩; 4—硅质岩; 5—炭质泥岩; 6—泥岩; 7—白云岩; 8—灰岩; 9—生物碎屑灰岩; 10—千枚岩; 11—平行不整合/角度不整合界线; 12—逆冲推覆断裂; Q—第四系; T_3a^{st} —上三叠统安源组三丘田段; T_3a' —上三叠统安源组三家冲段; T_3a'' —上三叠统安源组紫家冲段; P_3c —上二叠统长兴组; P_3l —上二叠统乐平组老山段; P_3f —上二叠统乐平组官山段; P_2ms —中二叠统鸣山组; P_2x —中二叠统小江边组; P_2q —中二叠统栖霞组; P_1l —中二叠统梁山组; P_1m —下二叠统马平组; C_2h —上石炭统黄龙组; $Pt_5^{1a}W$ —新元古界万年群; γ_5^2 —燕山早期花岗岩; $\gamma\delta\pi_2$ —晋宁期花岗岩闪长斑岩

Fig. 2 Structural section of the Zhuxi ore concentration area (modified from Zhong, 1992)

1—Floating soil; 2—Quartz sandstone; 3—Sandstone; 4—Siliceous rock; 5—Carbonaceous mudstone; 6—Mudstone; 7—Dolomite; 8—Limestone; 9—Bioclastic limestone; 10—Phyllite; 11—Parallel/angular unconformity boundary; 12—Thrust nappe fracture; Q—Quaternary; T_3a^{st} —Upper Triassic Anyuan Formation Sanqitian member; T_3a' —Upper Triassic Anyuan Formation Sanjiachong member; T_3a'' —Upper Triassic Anyuan Formation Zijiaichong member; P_3c —Upper Permian Changxing Formation; P_3l —Upper Permian Leping Formation Laoshan member; P_3f —Upper Permian Leping Formation Guanshan member; P_2ms —Middle Permian Mingshan Formation; P_2x —Middle Permian Xiaojiangbian Formation; P_2q —Middle Permian Qixia Formation; P_1l —Middle Permian Liangshan Formation; P_1m —Lower Permian Maping Formation; C_2h —Upper Carboniferous Huanglong Formation; $Pt_5^{1a}W$ —Neoproterozoic Wannian Group; γ_5^2 —Early Yanshanian granite; $\gamma\delta\pi_2$ —Jinning period granodiorite porphyry

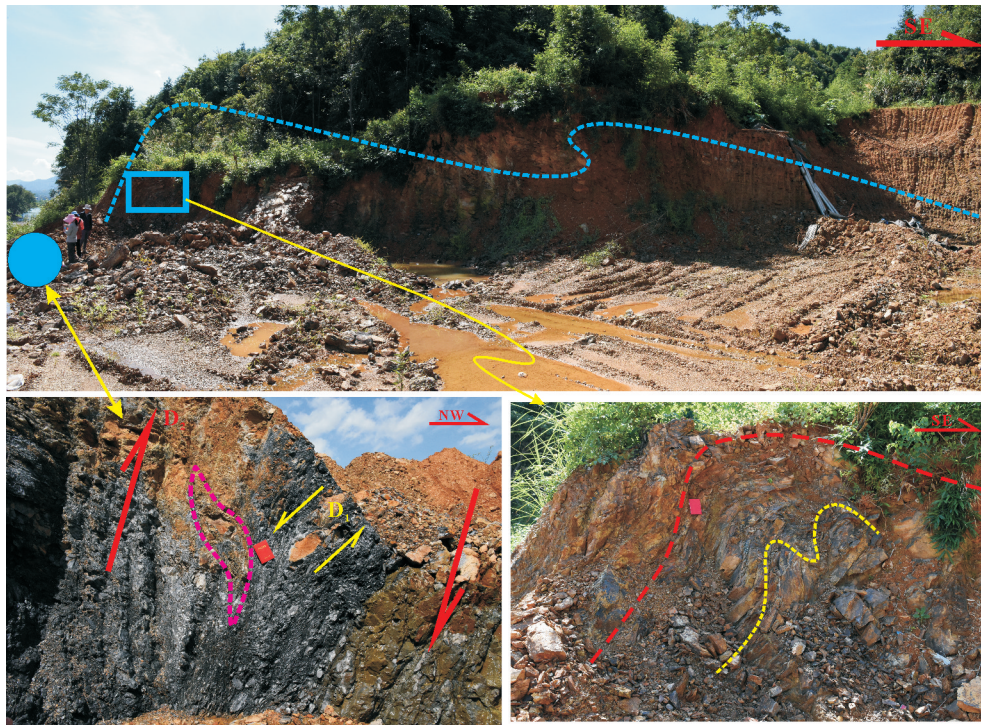


图3 塔前—赋春断裂带内“对冲”构造特征(张达提供)

Fig.3 Characteristics of hedging structures in the Taqian–Fuchun fault zone (after Zhang Da)

1217.5 m深处尚未见到断裂变缓,最近施工的钻孔发现断裂浅部呈铲式。东部镇头一赋春段产状趋缓,并出现飞来峰、构造窗构造(图4),总体具铲式特征。

横路一大游山逆冲推覆深断裂带:为朱溪岩片

与涌山岩片的分界断裂带,杨草尖万年群变质岩块逆冲推覆于横路一大游山向斜北侧,使向斜北翼遭到破坏叠覆,残留下向斜南翼成为单斜构造。在乐平市涌山万年群逆冲于上三叠统安源组煤系之上,

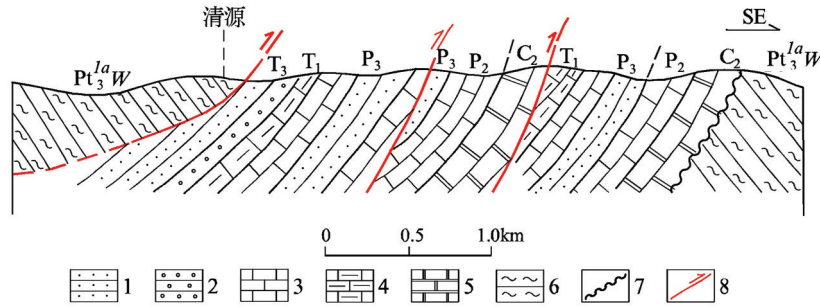


图4 塔前一赋春断裂带构造剖面素描图(钟南昌,1992)

1—粉砂岩;2—砾岩;3—灰岩;4—泥质灰岩;5—大理岩;6—绿泥石片岩;7—角度不整合界线;8—逆冲推覆断裂;T₃—上三叠统;T₁—下三叠统;P₃—上二叠统;P₂—中二叠统;C₂—上石炭统;Pt₃^{Ia}W—新元古界万年群

Fig.4 Sketch map of the structural section in the Taqian-Fuchun fault zone (after Zhong, 1992)

1-Siltstone; 2-Conglomerate; 3-Limestone; 4-Muddy limestone; 5-Marble; 6-Chlorite schist; 7-Angular unconformity boundary; 8-Thrust nappe fracture; T₃-Upper Triassic; T₁-Lower Triassic; P₃-Upper Permian; P₂-Middle Permian; C₂-Upper Carboniferous; Pt₃^{Ia}W-Neoproterozoic Wannian Group

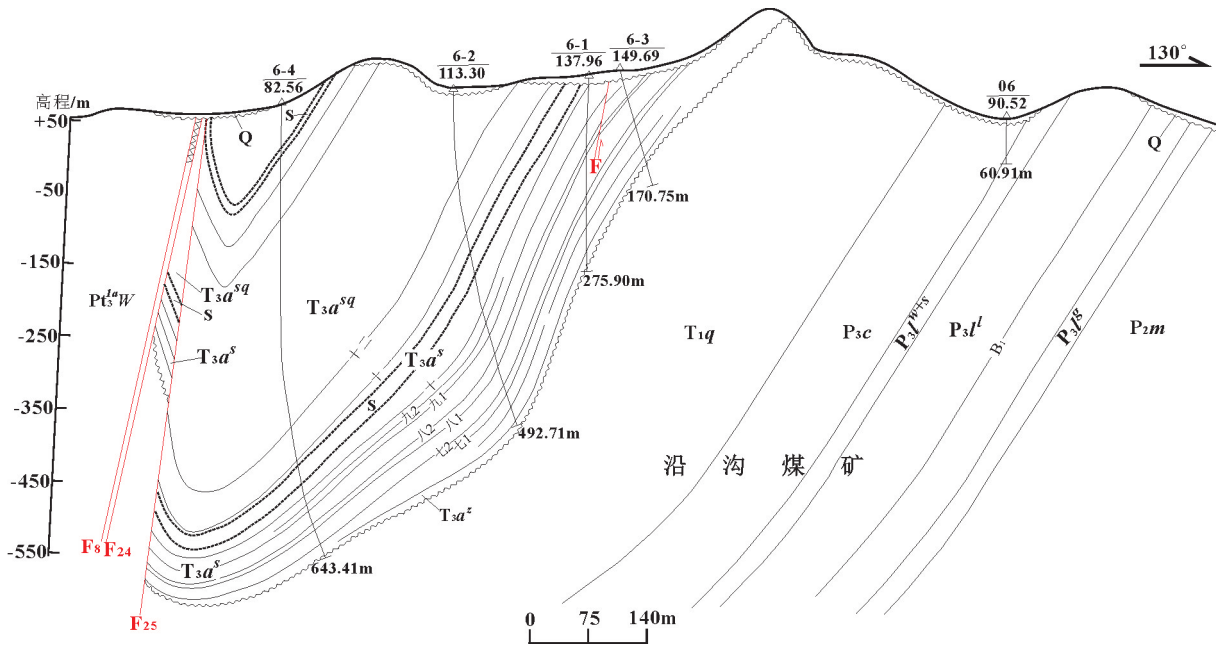


图5 涌山桥矿区6线剖面图(江西省地质矿产勘查开发局,2015)

Q—第四系;T₃a^{sq}—上三叠统安源组三丘田段;T₃a^s—上三叠统安源组三家冲段;T₃a^s—上三叠统安源组紫家冲段;T₁q—下三叠统青龙组;P₃c—上二叠统长兴组;P₃l^{W/S}—上二叠统乐平组王潘里段+狮子山段;P₃l^I—上二叠统乐平组老山段;P₃l^B—上二叠统乐平组官山段;P₂m—中二叠统茅口组;Pt₃^{Ia}W—新元古界万年群;七、八—煤层编号;S—砂(砾)岩层

Fig.5 Geological section along No. 6 exploration line of the Yongshanqiao mining area (after JBGMED, 2015)

Q-Quaternary; T₃a^{sq}-Upper Triassic Anyuan Formation Sanqituan member; T₃a^s-Upper Triassic Anyuan Formation Sanjiachong member; T₃a^s-Upper Triassic Anyuan Formation Zijiachong member; T₁q-Lower Triassic Qinglong Formation; P₃c-Upper Permian Changxing Formation; P₃l^{W/S}-Upper Permian Leping Formation Wangpanli and Shizishan member; P₃l^I-Upper Permian Leping Formation Laoshan member; P₃l^B-Upper Permian Leping Group Guanshan member; P₂m-Middle Permian Maokou Formation; Pt₃^{Ia}W-Neoproterozoic Wannian Group; 七、八-COal seam number; S-Sand (Gravel) rock formation

使得涌山向斜构造北翼(图5)遭到叠覆,但含煤向斜大部分得以保存(周宝直,2000),成为江西仅有的两个大型煤矿区之一。这里断面向北陡倾,总体上呈铲式特征。

3 构造控岩-控矿作用特征

矿集区受塔前一赋春、横路—大游山两条深断裂—侵入岩带控制,在空间上形成塔前一朱溪与张家坞一月形两个矿田。

3.1 塔前一朱溪钨铜钼矿田

该矿田受塔前一赋春深断裂控制,自SW向NE形成塔前一毛家园、弹岭、朱溪等矿床,成岩成矿具有I、S型两个系列和上下两个台阶,其中上台阶的毛家园、弹岭两个花岗闪长(斑)岩株,遭受较强剥蚀,成矿较弱;下台阶的朱溪与塔前矿段成矿条件优越。

(1)塔前一毛家园钨铜钼矿床

塔前一毛家园矿床包括塔前、毛家园两个矿段。塔前矿段与北侧的毛家园矿段相邻,故长期归于同一矿床。事实上,塔前矿段成矿在先,毛家园矿段成矿在后,成矿台阶前者在下,后者在上,矿床特征也有明显差异,塔前矿段以似层状矽卡岩型(白)钨铜钼矿体为主,而毛家园矿段主要为细脉型钨铜矿体,宜分别进行研究。

①毛家园矿段

毛家园矿段处于塔前一赋春断裂带上盘,成矿岩体为毛家园花岗闪长岩株,走向NE,呈椭圆状,长4 km,宽约1 km,侵入于万年群和石炭—三叠系中。沿接触带发生角岩化,在岩体西南部形成细脉型钨铜矿体(图6)。与铜厂、城门山等重要成铜岩体相比(Wang et al., 2012; Zhong et al., 2017), SiO_2 含量偏高为68.7%^①,故出现钨钼矿化,而铜成矿较弱。

②塔前矿段

塔前矿段形成于塔前一赋春导岩导矿深断裂带(F_1)下盘(图7)与万年隆起北坡逆冲构造前缘破坏矿床的 F_2 逆冲断裂带之间,形成强烈的对冲挤压结构。中部被毛家园花岗闪长岩体侵入,分为北东、南西两段,相伴有花岗闪长斑岩、闪长玢岩墙。主要有似层状矽卡岩型(白)钨铜钼矿体4条,产于茅口组或栖霞组及安源组中,分别为含铜钼白钨矿、白钨矿—辉钼矿、辉钼矿体,走向NE,长1200~1400 m,厚度6.11~15.83 m,矿体埋深控制至-350 m。受对

冲式构造挤压,岩层与矿体陡立倒转扭曲,矿床具中型规模,按照“以脉找体”的思路(杨明桂等,2008),深部可能存在成矿主侵入岩体,形成下台阶矿床,有待进一步勘查。

(2)弹岭铜铅锌矿床

成矿岩体侵入于黄龙组碳酸盐岩,为一小型花岗闪长斑岩瘤,黄龙组碳酸盐岩为区内最重要赋矿围岩,花岗闪长斑岩为区内重要成铜母岩,故其岩浆、围岩成矿条件均有利,经勘查在接触带有脉状、浸染状及矽卡岩型铜钨钼铅锌矿化,但成矿岩体与矿体规模均较小。

(3)朱溪钨铜矿床

朱溪矿床发现于20世纪60年代,中浅部经多次勘查为脉状、似层状小型铜矿床。21世纪以来,在全国第二空间找矿的推动下,江西省地矿局九一二大队通过“以层(脉)找体”矿区深部勘查,发现了巨厚的钨铜矿体,沿走向和倾向均控制超过2000 m(图8)。目前矿区已探获333类 WO_3 资源量285.5万t,居世界之首, WO_3 平均品位0.52%,其中富钨113万t, WO_3 平均品位1.793%;共伴生Cu 35.6万t、Zn 2.2万t、Ag 1165 t。白钨矿可选性良好,回收率达60.33%~80.24%,具有巨大的开发利用价值。且矿床沿走向和倾向均未控制边界,目前仍在勘查中, WO_3 资源量预计可超过400万t。

矿床处于塔前一赋春深断裂带下盘,赋矿地层为上石炭统一中下二叠统,呈NW倾斜的单斜构造,隐伏成矿岩体(黑云母花岗岩和云英岩化花岗岩为主)由塔前一赋春逆冲推覆深断裂带下部呈多峰式由NW向SE方向上侵,前锋岩枝(碱长二云母花岗岩)已上侵至-400 m标高处,主岩体潜伏于-1600 m标高处,与上台阶的毛家园和弹岭岩体垂直距离约1800 m。隐伏成矿岩体为富硅($\text{SiO}_2=76.93\%\sim 78.34\%$)、富铝($\text{Al}_2\text{O}_3=12.96\%\sim 14.13\%$)、富碱($(\text{K}_2\text{O}+\text{Na}_2\text{O})=4.79\%\sim 5.86\%$)、过铝质($A/\text{CNK}=1.06\sim 1.37$)岩石(Chen et al., 2016),主要为S型花岗岩组合(Ouyang et al., 2014; Chen et al., 2016; Song et al., 2018b; Pan et al., 2018)。与典型的S型南岭组合花岗岩比较(Su et al., 2011; Song et al., 2016; Zhong et al., 2017),主量、微量元素特征基本相同,唯稀土配分曲线与海鸥式不同而呈平缓右倾斜式, Eu具弱亏损,显示向I型花岗岩过渡色彩。



图6 塔前一毛家园矿区地质简图(江西省地质矿产勘查开发局物化探大队[®],2013修改)
 1—晚古生代—中生代沉积盖层;2—新元古代万年群;3—似斑状花岗闪长岩;4—中酸性岩脉(墙);5—花岗岩岩脉;
 6—角岩化带;7—钨钼矿体;8—逆冲推覆构造;9—勘探线及钻孔

Fig.6 Geological sketch map of the Taqian-Maojiayuan mining area (modified from Geophysical & Geochemical Exploration Party of JBGMED[®], 2013)

1-Late Paleozoic-Mesozoic sedimentary caprock; 2-Neoproterozoic Wannian Group; 3-Porphyritic granite;
 4-Medium-acid vein (wall); 5-Granite porphyry vein; 6-Angular lithification zone; 7-Tungsten-Molybdenum orebody;
 8-Thrust nappe structure; 9-Exploration line and drill hole

主要赋矿地层为不整合面之上的上石炭统黄龙组白云质灰岩、灰岩、白云岩;其次为下二叠统马平组灰岩。

矿体具上小下大、上陡下缓、上铜下钨铜、钨铜共生等特征(欧阳永棚等,2018)(图9)。如ZK5406孔,揭露钨铜矿体累积视厚度576.40 m,以外带接触交代型矿体为主,兼有蚀变花岗岩型、似层状、脉状以钨为主的“四位一体”矿床。成矿具有垂向分带性,浅部为黄龙组似层状铜矿体及裂隙脉带中的脉状铜矿体,深部为厚大矽卡岩型钨铜矿体,蚀变花岗岩中形成细脉浸染型钨铜矿体。钨矿物以白钨矿为主,矽卡岩化发育,矿体与矽卡岩或矽卡岩化岩石空间位置基本一致(吴筱萍等,2015)。富矿主要为矽卡岩型矿体,矿床以42线为中心,两侧不完全对称,西南段矿体隐伏较深,北东段矿体略有抬

高之势,其浅部铅锌银矿化较强。

朱溪巨型钨(铜)矿床的发现,在成矿作用方面,有诸多启示。

一是进一步表明钦杭成矿带是我国最重要的钨铜贵稀多金属-非金属成矿带之一,其中江西段钨矿资源储量约600万t,铜矿资源储量约1600万t。朱溪矿集区是萍乐拗陷带的一个逆冲推覆抬升区段,拗陷遭破坏;第一台阶成矿岩体遭较强侵蚀,处于第二台阶的朱溪巨型钨铜矿床的找矿重大突破,进一步揭示了钦杭成矿带萍乐拗陷带强叠覆、深拗陷、深隐伏的成矿特征,第二空间找矿潜力巨大。

二是朱溪矿床产于拗陷带导岩导矿的逆冲推覆深断裂带下盘,是成矿岩体与矿床隐伏较深的基本因素,而且矿床未遭受后期浅层对冲断裂构造破坏,是保存巨大矿床的重要条件。

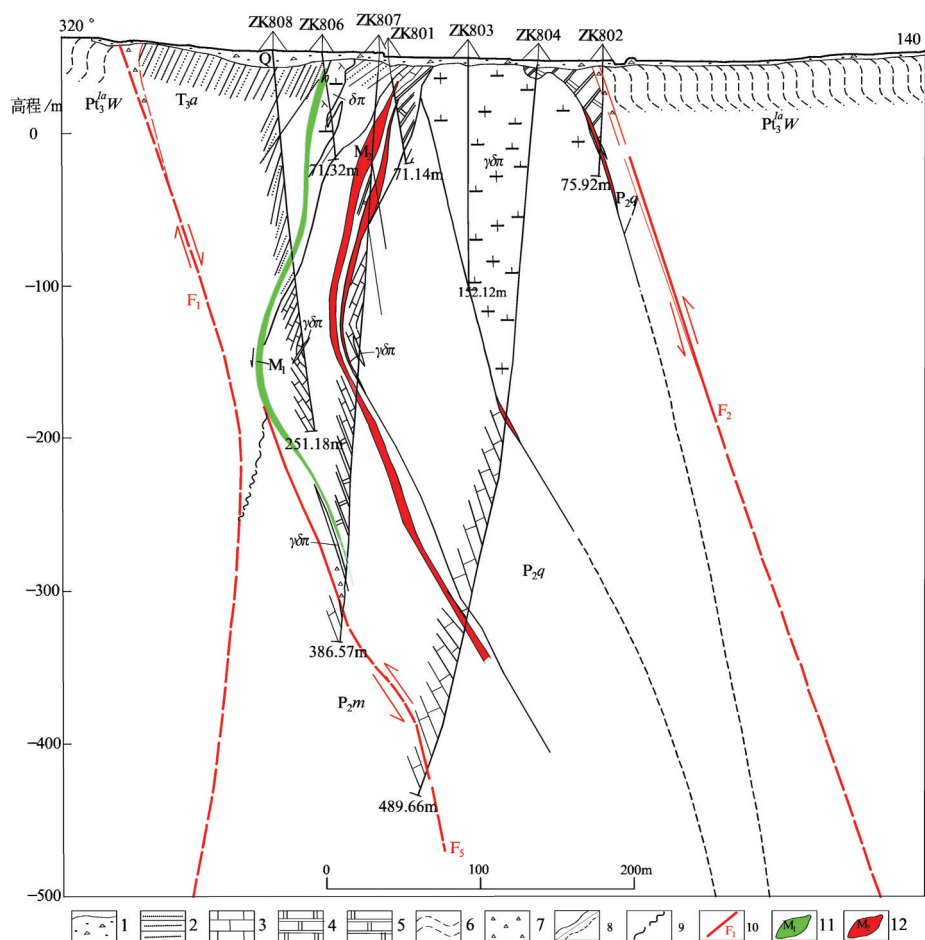


图7 塔前矿区8线剖面图(江西省地质矿产勘查开发局物化探大队,2013^⑥修改)

1—浮土;2—粉砂岩;3—灰岩;4—大理岩;5—砂卡岩;6—千枚岩;7—角砾岩;8—实测/推测地质界线;9—推测角度不整合界线;10—断裂;11—铜矿体及编号;12—钨钼矿体及编号;Q—第四系;T_{3a}—上三叠统安源组;P_{2m}—中二叠统茅口组;P_{2q}—中二叠统栖霞组;Pt_{3^{1a}}W—新元古界万年群;δπ—闪长斑岩;γδπ—花岗闪长斑岩

Fig.7 Geological section along No. 6 exploration line of the Taqian mining area (modified after Geophysical & Geochemical Exploration Party of JBGMED^⑥, 2013)

1—Floating soil; 2—Siltstone; 3—Limestone; 4—Marble; 5—Skarn; 6—Phyllite; 7—Breccia; 8—Measured/ inferred geological boundary; 9—Inferred angular unconformity boundary; 10—Fracture; 11—Copper orebody and its serial number; 12—Tungsten-molybdenum orebody and its serial number. Q—Quaternary; T_{3a}—Upper Triassic Anyuan Formation; P_{2m}—Middle Permian Maokou Formation; P_{2q}—Middle Permian Qixia Formation; Pt_{3^{1a}}W—Neoproterozoic Wannian Group; δπ—Diorite porphyry; γδπ—Granodiorite porphyry

三是矿床浅部为产于上石炭统黄龙组中的似层状铜矿体、长期被部分专家学者视为晚石炭世同生(火山)沉积矿床或晚石炭世海底热水喷流沉积、燕山期岩浆热液叠改矿床。朱溪矿床勘查研究表明,矿床呈上铜下钨铜分带,上部铜矿体具似层状热液成因特征,中深部递变为砂卡岩钨铜矿体,均形成于晚侏罗世末((144±5) Ma~(150.6±1.5) Ma(李岩等,2014;刘善宝等,2017;Pan et al., 2017),且黄铜矿化稍晚于白钨矿化。再者上石炭统黄龙组以角度不整合沉积于万年群变质岩系之上,黄龙组为

浅海碳酸盐岩沉积,底部有数十厘米厚石英砾岩层,显然不具备海底热水喷流沉积的形成条件,而且钨铜矿化已进入不整合面之下,万年群浅变质岩中也有少量钨铜矿体。均表明花岗岩是成矿的必要条件,黄龙组及马平组碳酸盐岩是有利的围岩条件,构成了成矿的基本前提。

四是朱溪成矿S型花岗岩体,形成了巨型钨矿床和中型铜矿床,这是钦杭成矿带发现的又一钨铜共生矿床。笔者认为这种现象,很可能与钦杭板块结合带地壳中不均匀地混染有晋宁期华南的含铜洋壳物

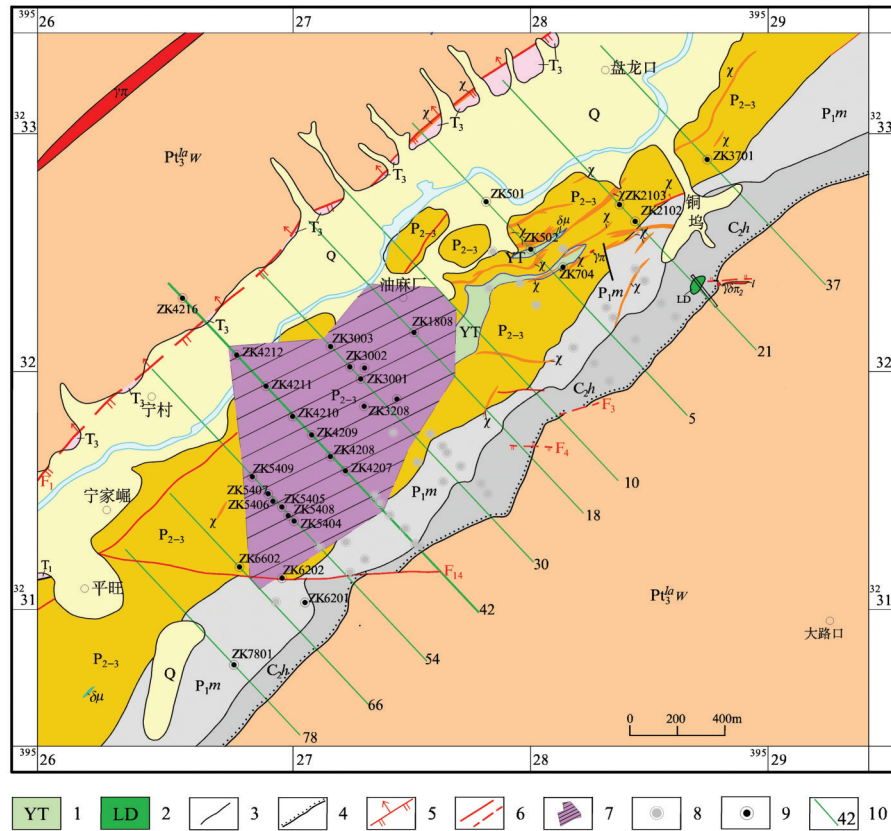


图8 朱溪矿区地质图(据江西省地质矿产勘查开发局九一二大队,2015^⑥修改)

1—透闪石-阳起石化带;2—绿色蚀变带;3—地质界线;4—角度不整合界线;5—逆冲推覆断裂;6—实测/推测断裂;7—矿体平面投影范围;8—以往施工钻孔;9—912队普查完工钻孔;10—勘探线及编号;Q—第四系;T₃—上三叠统;T₁—下三叠统;P₂₋₃—中-上二叠统;P_{1m}—下二叠统马平组;C_{2h}—上石炭统黄龙岗组;Pt₃¹⁰W—新元古界万年群;γπ—花岗岩;δμ—闪长玢岩;χ—煌斑岩;γδπ₂—晋宁期花岗闪长斑岩

Fig.8 Geological map of the Zhuxi deposit (modified from No. 912 Geological Surveying Party of JBGMED^⑥, 2015)

1—Tremolite-actinolite belt; 2—Green alteration belt; 3—Geological boundary; 4—Angular unconformity boundary; 5—Thrust nappe fracture; 6—Measured/inferred fracture; 7—Orebody plane projection range; 8—Previous drill hole; 9—Prospecting completed drill hole; 10—Exploration line and its serial number. Q—Quaternary; T₃—Upper Triassic; T₁—Lower Triassic; P₂₋₃—Middle-Upper Permian; P_{1m}—Lower Permian Mapping Formation; C_{2h}—Upper Carboniferous Huanglong Formation; Pt₃¹⁰W—Neoproterozoic Wannian Group; γπ—Granite porphyry; δμ—Diorite porphyry; χ—Lamprophyre; γδπ₂—Jinning period granodiorite porphyry

质有关(杨明桂等,2016b),值得进一步研究。

3.2 张家坞—月形铜矿田

该矿田位于横路—大游山深断裂带南东侧(下盘),总体呈NE向展布,以张家坞花岗闪长斑岩株及铜矿床为中心,南西有横路含铜岩脉群分布,北东为月形铜矿床,大致以2 km等间距分布(图10)。该矿田于1966年发现后,经多次勘查,进展不大,近期江西省地矿局物化探大队通过对月形矿区勘查,进一步查明了矿区构造与成矿特征,矿床远景不断扩大,初步估算铜金属量超过15万t,显示矿田具有较大铜资源潜力。

(1) 张家坞铜矿床

位于矿田中部的张家坞铜矿床,其成矿岩体张

家坞花岗闪长斑岩岩株为矿田内出露的最大规模侵入岩,走向NE,其两侧的岩墙群分别向南西侧的横路、北东侧的月形延展。岩体遭受较强剥蚀,在边缘接触带及岩体中部的围岩残留体有接触交代型、斑岩型铜矿体分布,经勘查规模不大。

(2) 月形铜矿床

于1966年由江西省地质局901队发现,经勘查为一小型铜矿床,2012年以来经江西省地矿局物化探大队在矿区东部隐伏区进一步工作,发现为一受断层错断的“断根”中型铜矿床,矿床处于张家坞花岗闪长斑岩岩株以东(图10),横路—大游山深断裂带下盘,由上石炭统—上三叠统组成NE向向斜构造,矿床产于向斜南翼,先后遭受F₃铲式逆冲断层

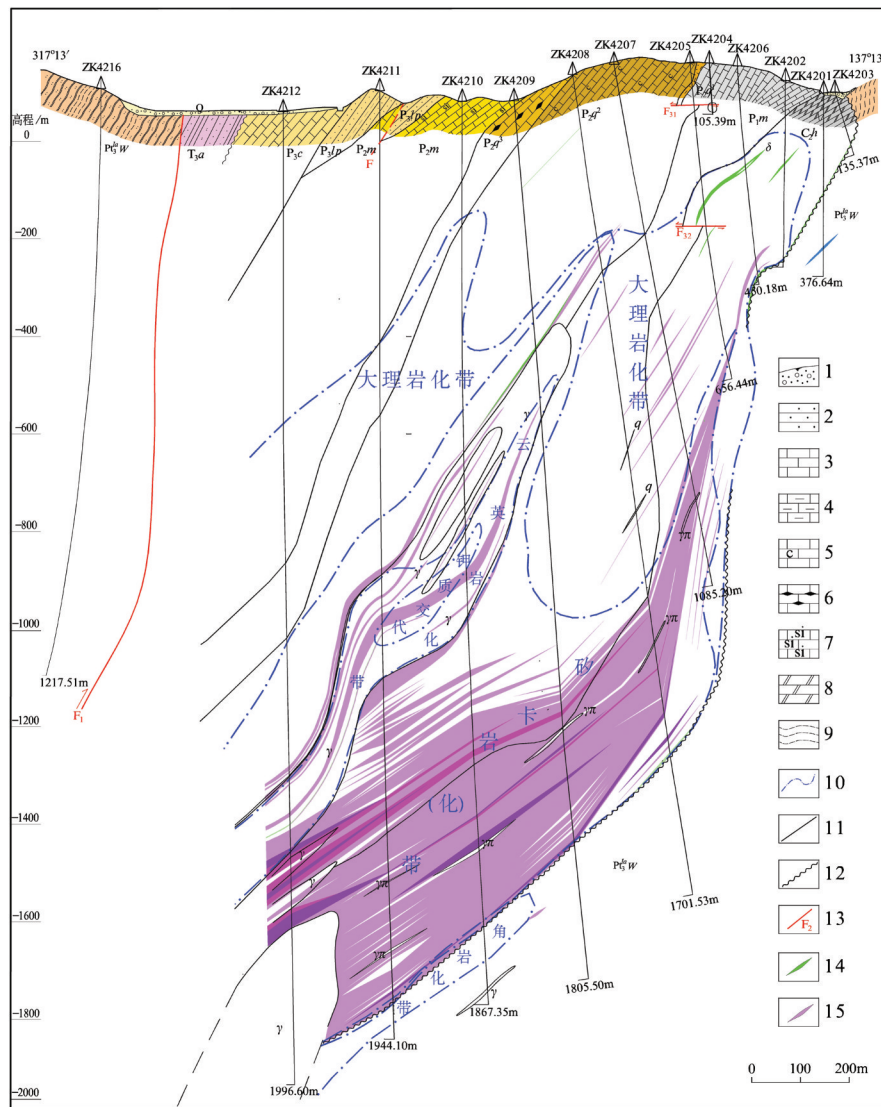


图9 朱溪矿床42线剖面图(据江西省地质矿产勘查开发局九一二大队[®], 2015 修改)

1—浮土;2—砂岩;3—灰岩;4—泥质灰岩;5—含炭灰岩;6—燧石灰岩;7—硅质灰岩;8—白云岩;9—千枚岩;10—蚀变带界线;11—地质界线;12—角度不整合界线;13—断裂;14—铜矿体;15—钨矿体;Q—第四系;T_{3a}—上三叠统安源组;P_{3c}—上二叠统长兴组;P_{3lp}—上二叠统乐平组;P_{2m}—中二叠统茅口组;P_{2q³}—中二叠统栖霞组上段;P_{2q²}—中二叠统栖霞组中段;P_{2q¹}—中二叠统栖霞组下段;P_{1m}—下二叠统马平组;C_{2h}—上石炭统黄龙组;Pt_{3^{1a}W}—新元古界万年群;γ—花岗岩;γπ—花岗斑岩;δ—闪长岩

Fig. 9 Geological section along No. 42 exploration line of the Zhuxi deposit (modified after No. 912 Geological Surveying Party of JBGMED[®], 2015)

1—Floating soil; 2—Sandstone; 3—Limestone; 4—Muddy limestone; 5—Carbonaceous limestone; 6—Chert limestone; 7—Siliceous limestone; 8—Dolomite; 9—Phyllite; 10—Alteration boundary; 11—Geological boundary; 12—Angular unconformity boundary; 13—Fracture; 14—Copper orebody; 15—Tungsten orebody; Q—Quaternary; T_{3a}—Upper Triassic Anyuan Formation; P_{3c}—Upper Permian Changxing Formation; P_{3lp}—Upper Permian Leping Formation; P_{2m}—Middle Permian Maokou Formation; P_{2q³}—Upper member of Middle Permian Qixia Formation; P_{2q²}—Middle member of Middle Permian Qixia Formation; P_{2q¹}—Lower member of Middle Permian Qixia Formation; P_{1m}—Lower Permian Maping Formation; C_{2h}—Upper Carboniferous Huanglong Formation; Pt_{3^{1a}W}—Neoproterozoic Wannian Group; γ—Granite; γπ—Granite porphyry; δ—Diorite

和向斜中部的F₂正断层破坏,矿床成了一个单斜状残体(图11)。

矿区出露有花岗闪长斑岩、花岗斑岩、花岗岩岩墙群。矿区已圈定铜多金属矿体39条,全部呈隐伏

状,走向上已控制2100 m,矿体主要呈似层状,其次为透镜状、脉状。赋矿地层主要为上石炭统黄龙组,其次为下二叠统马平组。主要为砂卡岩型铜铁矿体,已控制33号似层状矿体规模最大,长约850 m,斜深

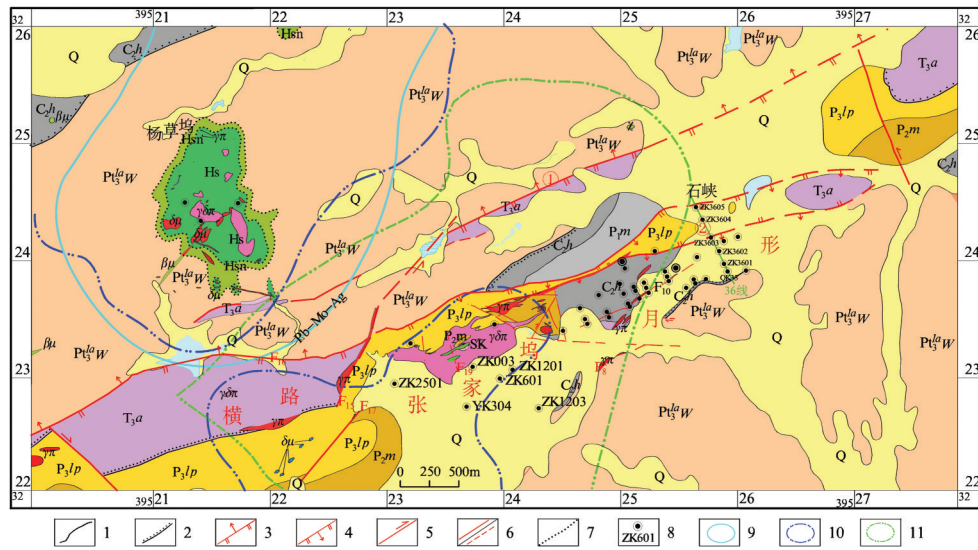


图10 横路—张家坞—月形矿带综合地质图

1—地质界线;2—角度不整合界线;3—逆冲推覆断裂;4—正断层;5—走滑断层;6—实测/推测性质不明断裂;7—蚀变界线;8—钻孔;9—水系沉积物异常;10—地磁异常;11—土壤异常;Q—第四系;T_{3a}—上三叠统安源组;P_{3lp}—上二叠统乐平组;P_{2m}—中二叠统茅口组;P_{1m}—下二叠统马平组;C_{3h}—上石炭统黄龙组;Pt_{3¹W}—新元古界万年群;γπ—花岗岩;γδπ—花岗闪长岩;δμ—闪长玢岩;βμ—辉绿岩;χ—煌斑岩;①—横路—大游山深断裂带;②—月形地堑式正断层

Fig.10 Geological map of the Henglu-Zhangjiawu-Yuexing ore belt

1-Geological boundary; 2-Angular unconformity boundary; 3-Thrust nappe fracture; 4-Normal fault; 5-Strike slip fault; 6-Measured/inferred unidentified fault; 7-Alteration boundary; 8-Drill hole; 9-Stream sediment geochemical anomaly; 10-Ground high accuracy magnetic survey anomaly; 11-Soil geochemical anomaly; Q-Quaternary; T_{3a}-Upper Triassic Anyuan Formation; P_{3lp}-Upper Permian Leping Formation; P_{2m}-Middle Permian Maokou Formation; P_{1m}-Lower Permian Maping Formation; C_{3h}-Upper Carboniferous Huanglong Formation; Pt_{3¹W}-Neoproterozoic Wannian Group; γπ-granite porphyry; γδπ-Granodiorite porphyry; δμ-Diorite porphyry; βμ-diabase; χ-Lamprophyre; ①-Henglu-Dayoushan deep fault zone; ②-Yuexing mantle type normal fault

386 m, 平均厚度 6.53 m, 主要矿体 Cu 平均品位为 1.7%~3.5%。矿床最大延深为 32 号矿体, 达 458 m。

为追索断失的下部矿床, 施工的 ZK3604 孔于 1266.33 m 处钻遇 F₂ 下盘黄龙组中的似层状铜矿体, 厚度 2.96 m, Cu 品位 0.42%~0.57%。这说明矿区向斜北翼可能存在下台阶隐伏成矿岩株和相关的接触交代型、斑岩型、似层状矿体, 有待验证。

(3) 横路铜矿点

位于张家坞矿床南西约 2 km, 有含铜花岗闪长岩墙群分布, 水系沉积物钼铜异常范围较大, 未进行详细勘查工作。

月形、横路含铜岩墙群, 长期以来大都认为属张家坞花岗闪长岩岩株的卫星岩脉。根据月形矿区最新勘查资料, 岩墙群与铜矿、张家坞岩株的最大距离已达 2200 m, 且未发现自岩株由近至远的矿化分带特征, 矿化较岩株接触带明显较强, 不排除月形矿床 F₂ 正断层下盘(北西侧)存在被错断的成矿岩体。若如此, 矿田可能出现横路、张家坞、月

形上下两个台阶三星连珠式成矿岩体与矿床。

此外, 在横路以北的杨草尖地区有黑云母花岗闪长岩岩瘤(图 10)侵入于万年群中, 有水系沉积物铜钨地球化学异常, 经勘查有铜钨金矿化, 布格重力异常显示岩体深部有可能增大, 值得进一步调查研究。

4 讨论

华南板块自古元古代以来经历了多期次构造演化, 其中中生代以来构造-岩浆作用尤为显著, 自中晚侏罗世开始受古太平洋、特提斯及西伯利亚多向板块汇聚作用的影响, 华南地区出现大规模的岩石圈挤压增厚, 并形成大规模 NE-SW 走向推覆构造, 导致华南板块岩石圈加厚, 并诱发下地壳持续重熔, 为中生代大规模岩浆作用及成矿奠定了重要的构造背景(张岳桥等, 2009; 徐先兵等, 2009; Li et al., 2014; 霍海龙等, 2018)。

具体到朱溪矿集区, 燕山运动早期中晚侏罗世在区内发生了自 NNW 向 SSE 的深层逆冲推覆作用, 形

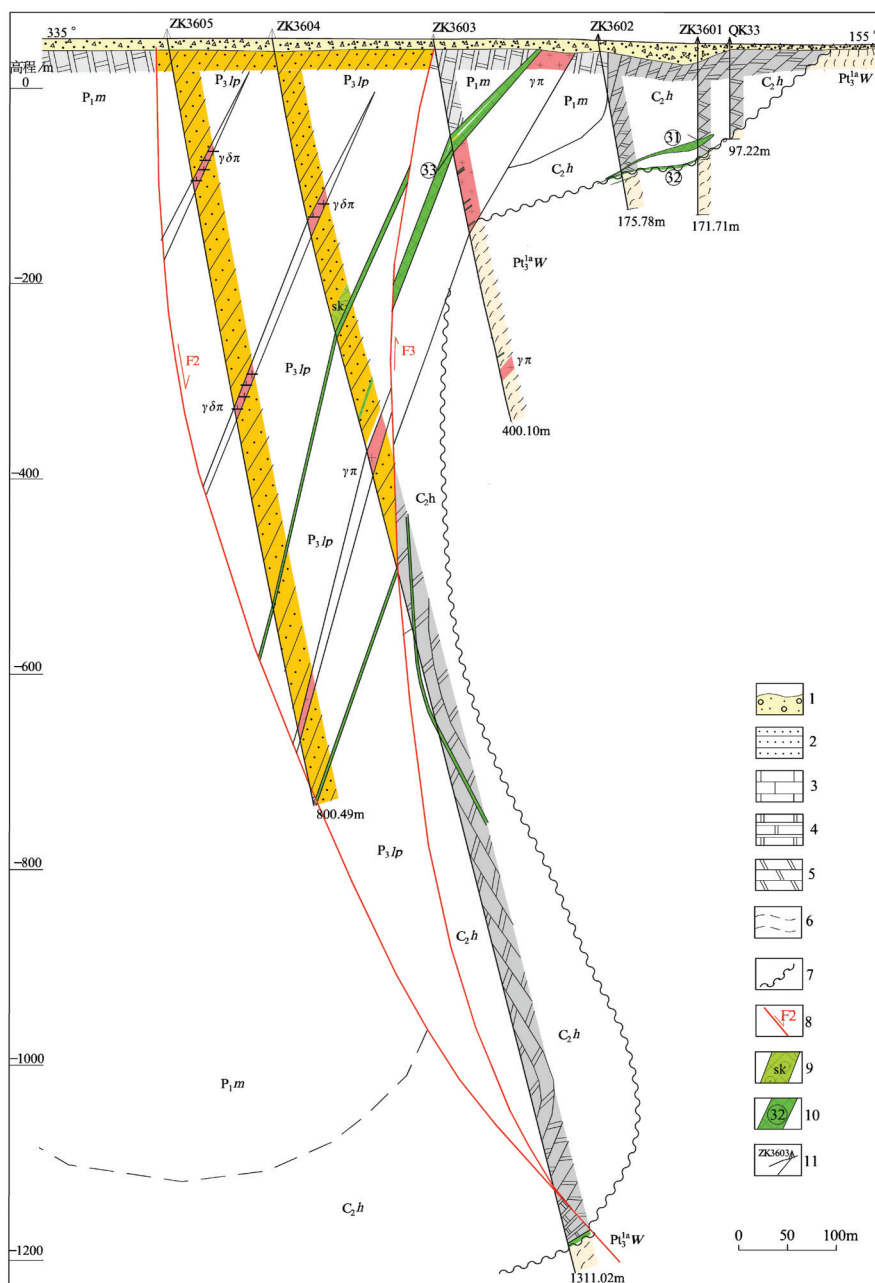


图 11 月形矿区 36 线剖面图

1—浮土;2—砂岩;3—灰岩;4—大理岩;5—白云岩;6—千枚岩;7—角度不整合界线;8—正断层及编号;9—矽卡岩;
10—铜矿体及编号;11—钻孔及编号。Q—第四系;P₃lp—上二叠统乐平组;P₁m—下二叠统马平组;C₂h—上石炭统黄龙组;
Pt₃^W—新元古界万年群;γδπ—花岗闪长斑岩;γπ—花岗斑岩

Fig.11 Geological section along No. 36 exploration line of the Yuexing deposit

1—Floating soil; 2—Sandstone; 3—Limestone; 4—Marble; 5—Dolomite; 6—Phyllite; 7—Angular unconformity boundary; 8—Normal fault and its serial number; 9—Skarn; 10—Copper orebody and its serial number; 11—Drill hole and its serial number. Q—Quaternary; P₃lp—Upper Permian Leping Formation; P₁m—Lower Permian Maping Formation; C₂h—Upper Carboniferous Huanglong Formation; Pt₃^W—Neoproterozoic Wannian Group; γδπ—Granite; γπ—Granite porphyry

成由变质基底与沉积盖层组成的构造岩片堆叠结构, 致使基底与盖层出现较复杂的空间叠置关系, 且活化了基底与盖层不整合面从而形成滑脱面; 除主推覆构造

造界面及不整合面之外, 挤压过程中在沉积盖层中亦会产生一系列垂直主压应力的次级破裂面, 推覆构造界面及伴生的次级断裂成为区内主要储矿构造, 当构

造属性由挤压转变为伸展状态时,岩浆岩沿早期破裂面侵位,含矿热液与碳酸盐岩发生交代成矿(霍海龙等,2018),导致沿七宝山—朱溪逆冲推覆深断裂带形成了I型与S型花岗岩带及朱溪多金属矿集区(杨明桂等,2015,2016a,2016b,2018)。大约于燕山晚期北侧的九岭逆冲推覆隆起再次向SSE方向逆冲推覆,南侧的万年隆起北坡向NNW方向滑脱,二者前缘于七宝山—朱溪一带发生对冲,使该带构造进一步复杂化(黄雷等,2017;杨明桂等,2018),同时使得燕山早期形成的矿床遭到严重破坏,由于推覆抬升多数遭到剥蚀成为残体,从而形成了一批被错断的“断根”矿床(诸如月形铜矿床和弹岭铜矿床),另一部分矿床(诸如朱溪钨铜多金属矿床)遭到推覆而埋深加大,增加了找矿难度。

5 结 论

(1)通过整体性时空四维结构研究,揭示了朱溪矿集区具有逆冲推覆深断裂带控岩—控矿、碳酸盐岩构造岩片赋矿、燕山晚期浅层对冲构造破矿的特征;发育有燕山早期I型花岗闪长(斑)岩钨铜和S型花岗岩钨铜两个岩浆岩成矿系列,在空间上形成张家坞—月形与塔前—朱溪两个矿田、张家坞—毛家园和塔前—朱溪上下两个成岩台阶。运用“以层(脉)找体”、“研断寻根”的思路,结合“多位一体”矿床模式,提出了塔前矿段、月形矿床以及横路地区深部找矿的预测意见,展示了区内第二台阶具有很大的钨铜多金属资源潜力及良好的找矿前景。

(2)朱溪矿集区为钦杭成矿带拗陷区中一个构造岩片堆叠抬升区段,处于矿集区第二成矿台阶的朱溪巨型钨铜矿床的发现是“深地”找矿的重大突破,揭开了萍乐拗陷带和钦杭成矿带广阔拗陷区的找矿序幕。有待于进一步跟踪研究和勘查,以取得更多更大找矿成果。

注释

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