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内蒙古西乌旗罕乌拉地区白音高老组火山岩特征及形成构造背景

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摘要:大兴安岭中生代火山岩的成因和构造背景一直存在争议。内蒙古西乌旗地区发育大面积的晚中生代火山岩, 是中国东部巨型火山岩带的重要组成部分。本文对西乌旗罕乌拉地区白音高老组火山岩开展了野外地质、岩石学、锆石 U-Pb 同位素年代学、地球化学研究, 以便对其岩石成因和构造背景给予制约。白音高老组火山岩主要由流纹岩及流纹质火山碎屑岩等一套中酸性火山岩组成。采集其中的球粒流纹岩和英安斑岩进行 LA-ICP-MS 锆石 U-Pb 测年, 测年锆石的 CL 图和 Th/U 值(0.34~1.25) 指示其为岩浆成因锆石, 测年结果分别为(140±0.8)Ma 和(133±0.7)Ma, 表明这套火山岩的形成时代为早白垩世早期。岩石地球化学研究表明, 白音高老组火山岩属高钾钙碱性系列, 具高硅、富碱、贫镁、钙, 高 FeO⁷/MgO 比值, 低 Mg[#] 值、Nb/Ta 比值的特征; 相对富集轻稀土元素, 亏损重稀土元素; 大部分样品富集 LILE, 而亏损 Ba、Sr 和 HFSE, 具 A 型花岗岩地球化学特征, 形成于伸展构造背景, 为地壳部分熔融的结果。结合区域中生代火山岩的空间展布特征, 认为该火山岩形成应与蒙古—鄂霍茨克洋闭合碰撞后伸展和古太平洋板块的俯冲作用有关。

关键词: 大兴安岭; 西乌旗; 白音高老组; 早白垩世; A 型花岗岩; 伸展构造环境; 地质调查工程
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Characteristics and tectonic setting of volcanic rocks of Baiyingaolao Formation in Hanwula of Xi Ujimqin Banner, Inner Mongolia

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Abstract: There exist different opinions concerning the petrogenesis and tectonic background of Mesozoic volcanic rocks developed in the Da Hinggan Mountains. The Late Mesozoic volcanic rocks in the Xi Ujimqin Banner of Inner Mongolia is a very important part of the huge volcanic rock belt in eastern China. The authors studied the volcanic rocks of Baiyingaolao Formation in Hanwula of Xi Ujimqin Banner in such aspects as field occurrence, petrology, zircon U–Pb isotopic geochronology and geochemistry in order to constrain their petrogenesis and tectonic background. The volcanic rocks of Baiyingaolao Formation are composed of rhyolite and volcanic clastic, which are a set of felsic volcanic rocks. The cathodoluminescence (CL) images of analyzed zircons of the pyromeride and dacite porphyry from Baiyingaolao Formation and their Th/U ratios (0.34–1.25) imply the igneous origin. LA–ICP–MS U–Pb dating shows that their ages are about (140±0.8)Ma and (133±0.7)Ma respectively, suggesting the early period of Early Cretaceous. Petrological and geochemical data reveal that the rocks belong to the high potassium calc–alkaline rock series characterized by rich Si and alkali, poor magnesium and calcium, high FeO⁷/MgO ratio and low Mg[#], Nb/Ta ratio. LREE are richer than HREE. The trace element geochemistry is characterized evidently by enrichment of LILE, depletion of Ba, Sr and HFSE. All these geochemical characteristics of rocks show an affinity with the A–type granites, which were most probably formed in an extensional setting and originated from the partial melting of the crust. Combined with spacial distribution of the Mesozoic volcanic rocks, the authors hold that they were probably related to the post–orogenic extension following the closure of the Mongol–Okhotsk orogen, and were also affected by the subduction of the Paleo–Pacific plate.

Key words: Da Hinggan Mountains; Xi Ujimqin Banner; Baiyingaolao Formation; Early Cretaceous; A–type granite; extensional environment; geological survey engineering

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1 引言

大兴安岭地区位于中亚造山带东段,古生代以来,该区经历了古亚洲洋闭合、蒙古—鄂霍茨克洋闭合及古太平洋板块俯冲等演化阶段,其构造和岩浆作用复杂多样(赵越等,1994;吴福元等,1999; Zorin, 1999; Kravchinsky et al., 2002; Wu et al., 2002; 李锦轶等, 2004; Sorokin et al., 2004; 孙德有等, 2004; 邵济安等, 2010, 2015; 余宏全等, 2012; Chen et al., 2016; Dash et al., 2016)。该地区中生代岩浆活动强烈,形成了大面积的火山岩和众多的花

岗岩(吕志成等,2004; Ying et al., 2010; 孟凡超等, 2013; Xu et al., 2013a; 许文良等, 2013b; 赵丕忠等, 2014),既是重要的多金属成矿带,也是研究中国乃至东北亚中生代构造岩浆演化的重要地区之一(葛文春等, 1999; Davis et al., 2001; 林强等, 2004; Wu et al., 2011; 马玉波等, 2016)。而大兴安岭地区中生代火山岩的成因及其产生的构造背景争议较大,该问题的研究对于揭示区内中生代构造—岩浆作用和成矿地质背景具有重要的意义。

大兴安岭地区中生代火山活动主要发生在晚侏罗世—早白垩世(张吉衡, 2009),在中南部地区

主要出露为满克头鄂博组、玛尼吐组、白音高老组
和梅勒图组。早期研究者对白音高老组火山岩进行
了大量研究工作,但在岩石的形成时代、岩石成因
和构造背景上还存在争议。形成时代上有晚侏
罗世(赵国龙等,1989;李文国等,1996;邵积东等,
2011)、早白垩世(葛文春等,2001;张吉衡,2009;苟
军等,2010;王建国等,2013;刘哲等,2017)。岩石
成因上一般认为是陆壳部分熔融的产物,但有与A₁
型花岗岩(Jahn et al., 2000, 2001;王兴安等,2012;
秦涛等,2014)、A₂型花岗岩(王雄等,2015;张学斌
等,2015;司秋亮等,2016)和I型花岗岩(Wu et al.,
2003; Dong et al., 2014;李研等,2017)相当的3种认
识。构造背景上一般认为该期火山-岩浆活动是蒙
古-鄂霍次克洋闭合后的后碰撞拉伸作用的响应
(Fan et al., 2003; 孟恩等, 2011; 徐美君等, 2011;
Ouyang et al., 2013, 2015; Xu et al., 2013a; Li et al.,

2016; 刘凯等, 2018); 另有学者认为该期岩浆活动
受控于古太平洋板块向古亚洲板块的俯冲作用(葛
文春等, 2007; Zhang et al., 2010); 还有部分学者认
为岩浆作用与地幔柱相关(林强等, 1998; 葛文春
等, 1999, 2005)。存在上述争议的原因主要是中生
代火山岩不同地区可能存在一定差异, 地层划分不
统一, 研究对象不一致; 早期研究, 火山岩时代的
确定主要依据岩石组合、古生物特征、区域地层对比
和Rb-Sr、K-Ar及少量U-Pb同位素数据(赵国龙
等, 1989; 李文国等, 1996; 邵积东等, 2011; 张学斌
等, 2015)。

笔者在西乌旗罕乌拉地区呼格吉勒图、巴彦
华、巴彦布拉格、彦吉嘎庙4个图幅1:5万区域地质
填图基础上, 以该区白音高老组火山岩为研究对
象, 进行岩石学、锆石U-Pb同位素年代学及岩石地
球化学等方面研究, 探讨该套岩石的成因、形成背

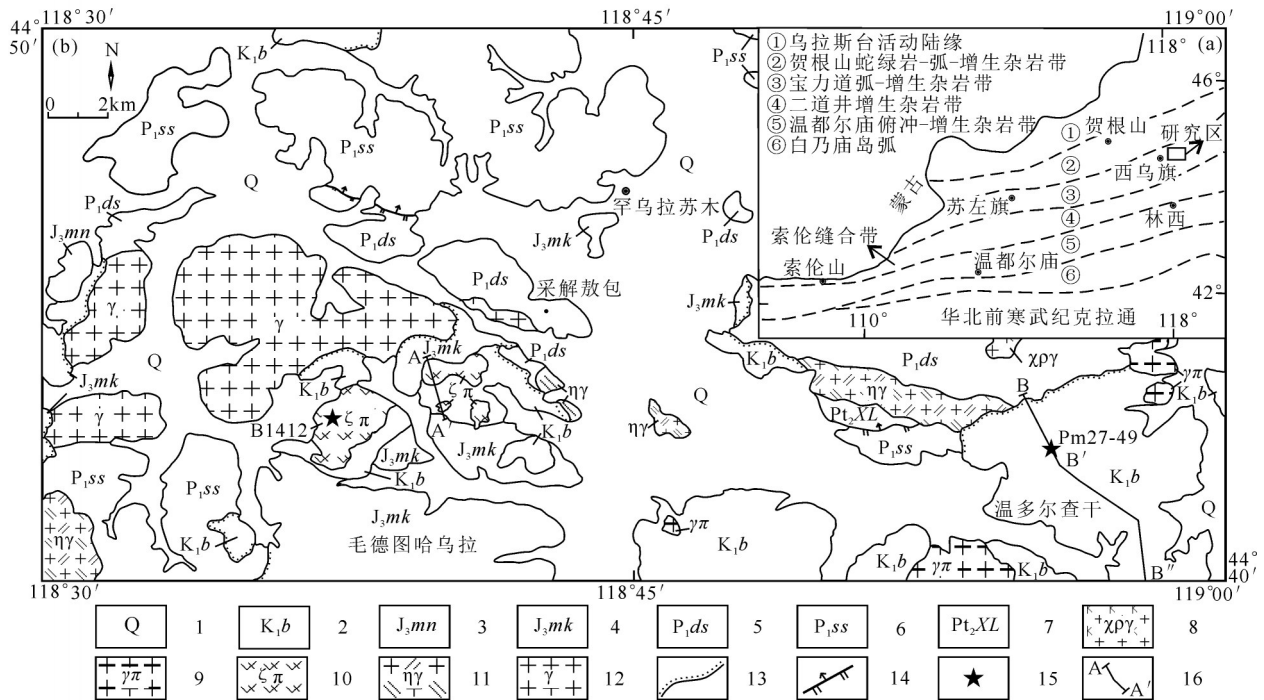


图1 研究区大地构造位置图(a, 据 Xiao et al., 2003)及区域地质简图(b)

- 1—第四系; 2—下白垩统白音高老组; 3—上侏罗统玛尼吐组; 4—上侏罗统满克头鄂博组; 5—下二叠统大石寨组; 6—下二叠统寿山沟组;
- 7—中元古界锡林浩特岩群; 8—早白垩世碱长花岗岩; 9—早白垩世花岗斑岩; 10—早白垩世英安斑岩; 11—早中三叠世侵入岩;
- 12—早二叠世侵入岩; 13—不整合接触界线; 14—实测断层; 15—同位素年龄采样点及编号; 16—剖面位置

Fig. 1 Regional geological location of the study area (a, after Xiao et al., 2003) and geological sketch map of the study area (b)
1—Quaternary; 2—Lower Cretaceous Baiyingaolao Formation; 3—Upper Jurassic Manitu Formation; 4—Upper Jurassic Manketouebo Formation;
5—Lower Permian Dashizhai Formation; 6—Lower Permian Shoushangou Formation; 7—Middle Proterozoic Xilinhot Group; 8—Early Cretaceous
alkali - feldspar granite; 9—Early Cretaceous granite porphyry; 10—Early Cretaceous dacite porphyry; 11—Early - Middle Triassic intrusive rock; 12—
Early Permian intrusive rock; 13—Unconformity; 14—Measured fault; 15—Isotopic age sampling position and serial number; 16—Location of section

景,希望对该地区中生代火山岩形成的构造背景提供科学依据。

2 区域地质背景

2.1 区域地质

内蒙古西乌旗地区位于华北板块与西伯利亚板块所夹持的中亚造山带南缘东段,处于贺根山蛇绿岩带和索伦山—西拉木伦河缝合带之间(图1a),大兴安岭中生代火山岩带中南段。研究区位于西乌旗东北部罕乌拉地区,出露的地层单元由老至新依次为中元古界锡林浩特岩群,为一套二云母(石英)片岩、含石榴石英片岩夹斜长角闪岩,与下二叠统寿山沟组为逆断层接触。二叠系从下到上依次为:下二叠统寿山沟组,为一套浅变质海相细碎屑岩系组合,与大石寨组为逆断层接触;下二叠统大石寨组,为一套海相火山—沉积岩组合;中二叠统哲斯组,为一套滨浅海相碎屑岩夹碳酸盐岩组合,发育较为丰富的珊瑚、腕足、海百合茎、苔藓虫化石。

中生界见有上侏罗统满克头鄂博组,为一套酸性火山碎屑岩—沉积碎屑岩岩石组合,含植物化石;上侏罗统玛尼吐组,为一套中基性火山熔岩岩石组合;下白垩统白音高老组为一套中酸性火山熔岩、火山碎屑岩夹火山碎屑沉积岩岩石组合。区内岩浆活动发育,晚古生代—中生代以未变质的花岗岩—中酸性火山岩为主,总体呈北东—北东东向展布。早二叠世发育辉长岩、辉石闪长岩及碱长花岗岩侵入体、大石寨组火山岩;早、中三叠世发育花岗闪长岩、二长花岗岩侵入体;晚侏罗世—早白垩世发育陆相火山岩及早白垩世侵入体。

2.2 白音高老组火山岩特征

白音高老组火山岩主要分布于研究区南部毛德图哈乌拉—温多尔查干一带,呈北东向展布,出露面积约75.5 km²,不整合覆盖在早二叠世辉石闪长岩、中三叠世二长花岗岩岩体之上,角度不整合覆盖在下二叠统寿山沟组、大石寨组之上。其分布见图1b,实测剖面见图2。岩石类型主要见有喷溢

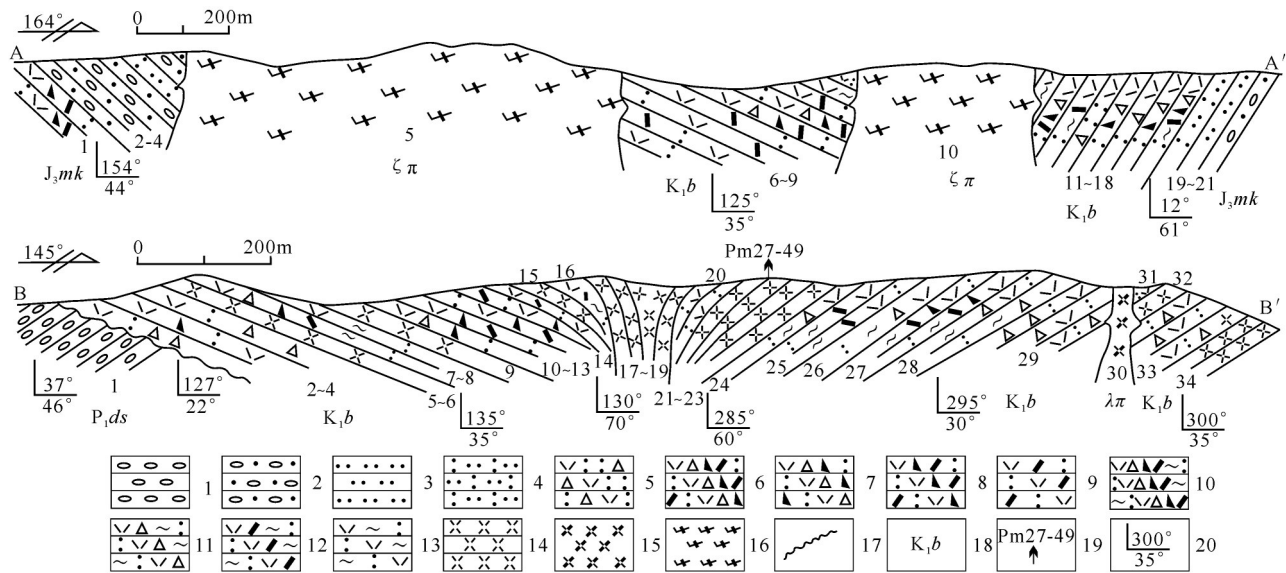


图2 罕乌拉地区白音高老组火山岩实测剖面

- 1—砾岩;2—砂砾岩;3—粉砂岩;4—凝灰质粉砂岩;5—流纹质凝灰角砾岩;6—流纹质含角砾岩屑晶屑凝灰岩;
- 7—流纹质含角砾岩屑凝灰岩;8—流纹质岩屑晶屑凝灰岩;9—流纹质晶屑凝灰岩;10—流纹质含角砾岩屑晶屑熔结凝灰岩;
- 11—流纹质含角砾熔结凝灰岩;12—流纹质晶屑熔结凝灰岩;13—流纹质熔结凝灰岩;14—流纹岩;15—流纹斑岩;16—英安斑岩;
- 17—不整合界线;18—地质代号;19—同位素年龄采样点及编号;20—产状

Fig. 2 Measured geological section of Baiyingaolao Formation in Hanwula

- 1—Conglomerates;2—Glutenites;3—Siltstones;4—Tuffaceous siltstones;5—Rhyolitic tuff breccias;6—Rhyolitic tuffs with breccia lithic and crystal clasts;7—Rhyolitic tuffs with breccia and lithic clasts;8—Rhyolitic tuffs with lithic and crystal clasts;9—Rhyolitic tuffs with crystal clasts;
- 10—Rhyolitic welded tuffs with breccia lithic and crystal clasts;11—Rhyolitic welded tuffs with breccia clasts;12—Rhyolitic welded tuffs with crystal clasts;13—Rhyolitic welded tuffs;14—Rhyolite;15—Rhyolite porphyry;16—Dacite porphyry;17—Unconformity;18—Geological code;19—Isotopic age sampling position and serial number;20—Attitude

相:灰白色、灰紫色流纹岩、球粒流纹岩、黑曜岩;爆发相:灰白色流纹质(含角砾)凝灰岩、流纹质凝灰岩、角砾岩、流纹质(含角砾)熔结凝灰岩、流纹质富晶屑熔结凝灰岩、英安质晶屑熔结凝灰岩等;次火山岩相:流纹斑岩、英安斑岩;喷发-沉积相:凝灰质砂砾岩等。火山岩相组合表现为爆发相-喷溢相系列组合,火山作用过程经历强烈爆发和宁静溢流的交互变化过程。火山活动晚期主要为次火山岩相产出,沿火山口或火山断裂(环形断裂及放射状断裂)侵入,形成似侵入岩的次火山岩。

流纹岩呈灰白色—青灰色—紫红色,具流纹构造,斑状结构(图3a、b)。斑晶占5%~10%,主要由斜长石、钾长石和石英组成,粒径0.2~1.5 mm。斜长石呈板状,隐约可见聚片双晶,主要为钠-更长石;钾长石呈板状,发育熔蚀结构,可见斑块状条纹;石英呈熔蚀不规则粒状。基质中发育流纹结构,沿流纹脱玻化程度不同,主要以长英质脱玻霏细粒状、隐晶质结构为主。

球粒流纹岩呈灰红色、浅紫色,野外露头上可见直径1~5 cm不等的石泡(图3c、d)。岩石具斑状结构、球粒结构。斑晶为斜长石、黑云母,自形程度高,粒径0.5~1 mm,含量小于5%。斜长石呈板状,聚片双晶宽窄不一,为钠更长石;黑云母呈鳞片状,褐色,表面浑浊;基质中圆形、椭圆形球粒结构发育,球粒细小密集,为纤维状长英质形成的放射状球体,直径0.15~1.0 mm,球粒之间被少量微粒长英质矿物和浑浊不规则的石英充填。副矿物主要为磷灰石、磁铁矿。

流纹质晶屑熔结凝灰岩,岩石风化面多呈灰红色、紫红色,新鲜面呈浅灰黑色、深紫色,假流动构造,熔结凝灰结构,主要由晶屑、浆屑及火山灰组成(图3e、f)。晶屑呈棱角状—次棱角状,少数自形晶,普遍有熔蚀,成分为石英、钾长石(条纹长石)、斜长石(黏土化),分布均匀,含量30%~40%。浆屑呈条带状、带状、透镜状,边缘轮廓不明显,脱玻化后边缘以纤维状长石为主,内部由嵌晶长英质或不规则粒状石英、长石组成,含量10%~20%。岩石中局部含角砾,且分布极其不均,含量1%~3%,其大小0.2~3 cm不等,大者可达5 cm,呈棱角状—次棱角状,岩性主要见有花岗岩、片岩、板岩等。

英安斑岩,呈次火山岩相产出,岩石多呈浅灰

红色、灰黄色,坚硬致密,具块状构造,斑状结构(图3g、h)。斑晶主要为斜长石、钾长石,含量10%~30%,粒径0.5~4 mm。斜长石呈板状,聚片双晶发育,为更长石。钾长石呈宽板状、聚斑状,条纹稀疏斑块状,为条纹长石。少量暗色矿物呈短柱状、粒状。基质粒状石英或长英质颗粒中嵌布柱状、粒状长石微晶,无规则排列。粒状石英、长英质颗粒界限不清,粒径小于0.2 mm。

3 样品与分析测试

本次共采集锆石U-Pb定年样品2件,样品编号为Pm27-49(球粒流纹岩)和B1412(英安斑岩),地理坐标分别为44°42'08"N, 118°55'56"E和44°42'54"N, 118°37'03"E。采集岩石地球化学样品7件,其中酸性火山岩岩性为流纹岩(B1404、B1408)、球粒流纹岩(Pm27-49)、流纹质晶屑熔结凝灰岩(Pm13-21),次火山岩岩性为英安斑岩(B1412、Pm21-88、Pm21-137)。

3.1 锆石U-Pb定年

样品破碎和锆石挑选均由河北省区域地质调查研究所实验室完成。采用常规方法进行粉碎,用常规浮选方法分选出锆石后,再用双目镜挑选出晶形和透明度较好的锆石颗粒作为测定对象。将锆石颗粒黏在双面胶上,经环氧树脂固定—环氧树脂固化—表面抛光工序后,进行锆石显微照相和阴极发光照相。锆石的反射光和透射光显微照相及阴极发光(CL)显微照相在北京锆年领航科技有限公司完成。

锆石U-Pb同位素定年在天津地质矿产研究所利用LA-ICP-MS进行分析,ICP-MS为ThermoFisher公司制造的Neptune。本次实验采用的激光束斑直径为35 μm,以氦气作为剥蚀物质的载气。具体实验过程及试验方法见李怀坤等(2010)。数据采用Andersen软件对测试数据进行普通铅校正,年龄计算及谐和图绘制采用ISOPLOT(2.49版)软件完成。所有数据点年龄值的误差均为1σ,采用²⁰⁶Pb/²³⁸U年龄,其加权平均值具95%的置信度(Anderson, 2002; Ludwig, 2003)。

3.2 主微量元素分析

样品碎样和地球化学成分测试工作均在河北省廊坊区域地质矿产调查研究所实验室完成。岩

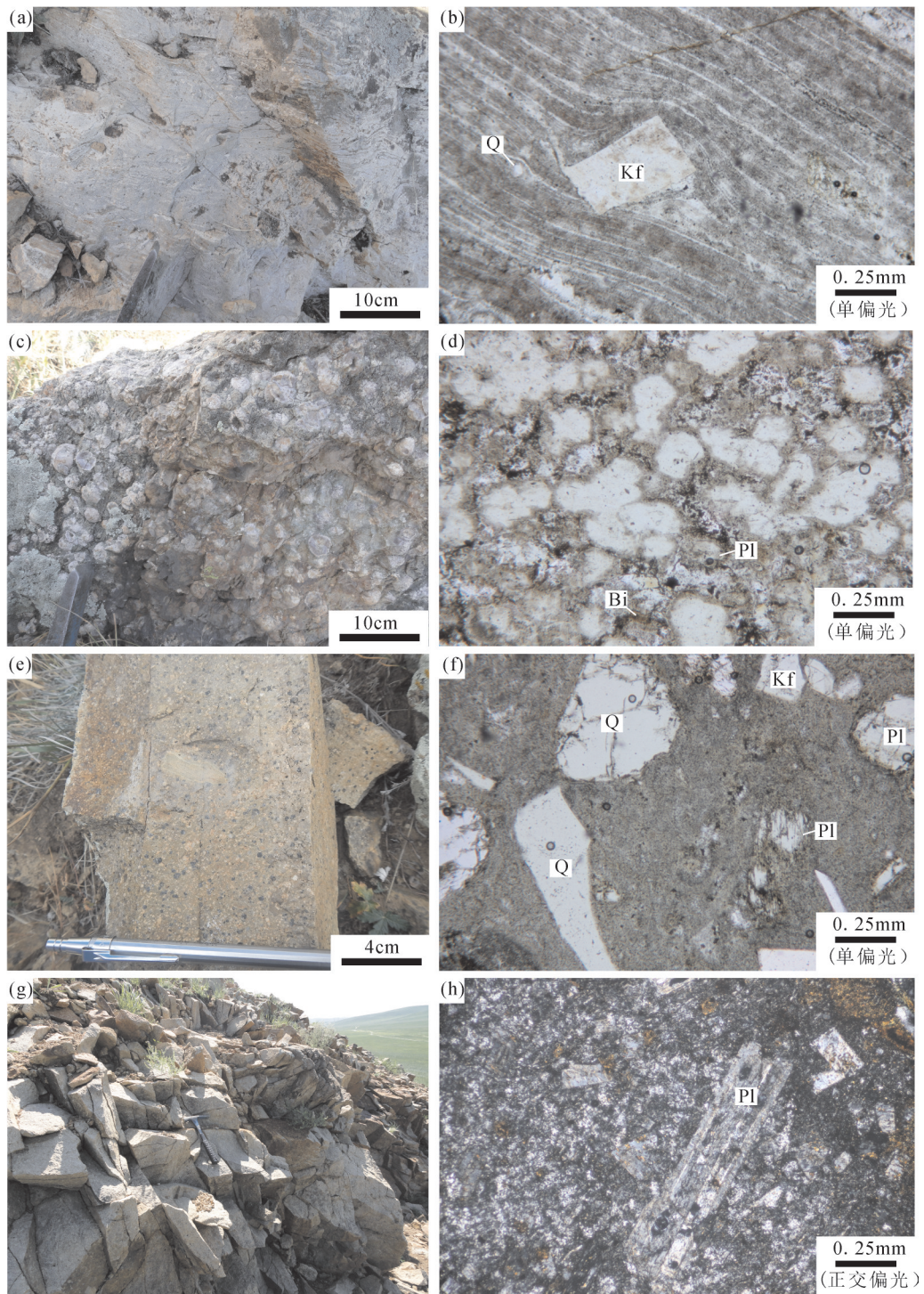


图3 白音高老组火山岩野外及显微照片

a,b—流纹岩; c,d—球粒流纹岩; e,f—流纹质晶屑熔结凝灰岩; g,h—英安斑岩; Q—石英; Pl—斜长石; Kf—钾长石; Bi—黑云母
 Fig. 3 Field outcrop and microscopic characteristics of Baiyingaolao Formation
 a,b—Rhyolite; c,d—Pyromeride; e,f—Rhyolitic welded tuffs with crystal clasts; g,h—Dacite porphyry; Q—Quartz; Pl—Plagioclase;
 Kf—K-feldspar; Bi—Biotite

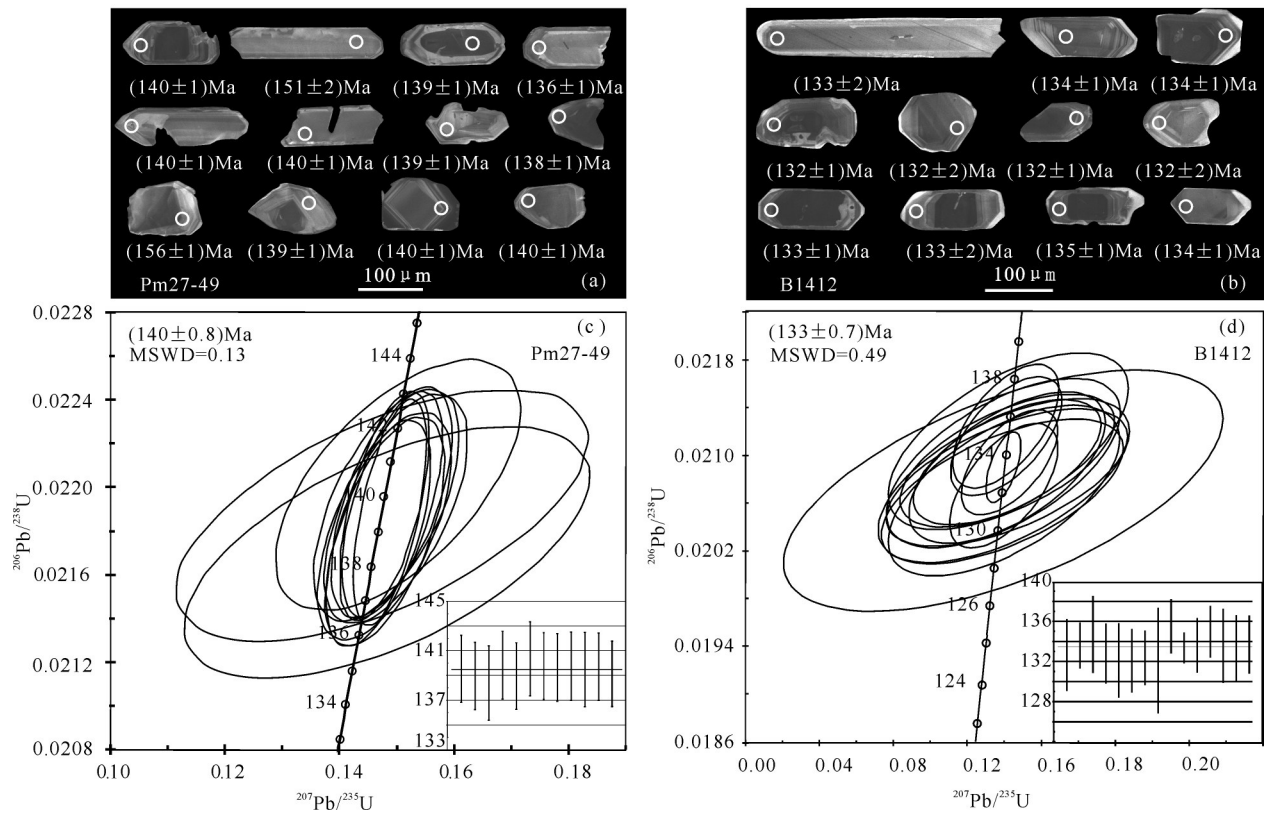


图4 罕乌拉地区白音高老组球粒流纹岩(Pm27-49)和英安斑岩(B1412)代表性单颗粒锆石阴极发光(CL)图像及其表面年龄(Ma)(a,b);罕乌拉地区白音高老组球粒流纹岩(Pm27-49)和英安斑岩(B1412)LA-ICP-MS 锆石U-Pb年龄谐和图(c,d)
Fig.4 Cathodoluminescence images of typical single-crystal zircons and their apparent ages (Ma) for the pyromeride (Pm27-49) and dacite porphyry (B1412) of Baiyingaolao Formation in Hanwula (a, b); LA-ICP-MS zircon U-Pb concordant age diagram for the pyromeride (Pm27-49) and dacite porphyry (B1412) of Baiyingaolao Formation in Hanwula (c, d)

石样品首先粗碎至2~4 cm,然后用3%~5%的稀盐酸经超声波清除表面杂质,再研磨至200目。主量元素采用X射线荧光光谱仪法,分析精度优于5%;微量元素采用X Serises 2 电感耦合等离子体质谱ICP-MS分析方法,精度和准确度优于5%。

4 分析结果

4.1 锆石U-Pb年代学

白音高老组火山岩样品 Pm27-49 (球粒流纹岩)的锆石为浅黄色-无色透明,在CL图像中,锆石自形程度较高,大部分呈长柱状,少量为短柱状(图4a),长度多为90~240 μm 。锆石具有明显的震荡环带,指示它们为岩浆成因锆石(Belousova et al., 2002; 吴元保等, 2004)。锆石LA-ICP-MS U-Pb分析结果(表1)显示,15个测点的 $^{206}\text{Pb}/^{238}\text{U}$ 表面年龄为 $(136 \pm 1) \text{ Ma}$ ~ $(140 \pm 1) \text{ Ma}$ (9、12号点分别为 $(156 \pm 2) \text{ Ma}$ 、 $(151 \pm 2) \text{ Ma}$,应为捕获锆石)。 $^{206}\text{Pb}/^{238}\text{U}$

和 $^{207}\text{Pb}/^{235}\text{U}$ 谐和性较好, $^{206}\text{Pb}/^{238}\text{U}$ 加权平均年龄为 $(140 \pm 0.8) \text{ Ma}$ ($\text{MSWD} = 0.13$)(图4c)。因此,球粒流纹岩的结晶年龄确定为 $(140 \pm 0.8) \text{ Ma}$ 。

白音高老组火山岩样品B1412(英安斑岩)的锆石为浅黄色-无色透明,在CL图像中,大部分锆石较自形,多呈长柱状,少量为短柱状(图4b),长度多为80~340 μm 。锆石LA-ICP-MS U-Pb分析结果(表1)显示,16个测点的 $^{206}\text{Pb}/^{238}\text{U}$ 年龄介于 $(132 \pm 1) \text{ Ma}$ ~ $(136 \pm 1) \text{ Ma}$, $^{206}\text{Pb}/^{238}\text{U}$ 和 $^{207}\text{Pb}/^{235}\text{U}$ 谐和性较好(图4d), $^{206}\text{Pb}/^{238}\text{U}$ 加权平均年龄为 $(133 \pm 0.7) \text{ Ma}$ ($\text{MSWD} = 0.49$)。因此,英安斑岩的结晶年龄确定为 $(133 \pm 0.7) \text{ Ma}$ 。

4.2 主量元素

白音高老组火山岩主量元素分析结果见表2。白音高老组酸性火山岩样品的 SiO_2 和碱含量较高,样品Pm13-21(流纹质晶屑熔结凝灰岩)的 K_2O 含量明显高于其他样品, Na_2O 和 $\text{Na}_2\text{O}/\text{K}_2\text{O}$ 比值低于

表1 罕乌拉地区白音高老组火山岩球粒流纹岩(Pm27-49)和英安斑岩(B1412)LA-ICP-MS 锆石 U-Pb 同位素分析结果
Table 1 LA-ICP-MS zircon U-Pb isotope analysis results for the pyromerite(Pm27-49)and dacite porphyry(B1412) of Baiyingaolao Formation in Hanwula

样品点号	含量/ 10^{-6}		同位素比值										表面年龄/Ma					
	Pb	U	$^{206}\text{Pb}/^{238}\text{U}$	1σ	$^{207}\text{Pb}/^{235}\text{U}$	1σ	$^{207}\text{Pb}/^{206}\text{Pb}$	1σ	$^{208}\text{Pb}/^{232}\text{Th}$	1σ	$^{232}\text{Th}/^{238}\text{U}$	1σ	$^{206}\text{Pb}/^{238}\text{U}$	1σ	$^{207}\text{Pb}/^{235}\text{U}$	1σ	$^{207}\text{Pb}/^{206}\text{Pb}$	1σ
球粒流纹岩(Pm27-49)																		
Pm27-49-01	61	2068	0.0219	0.0002	0.1488	0.0026	0.0493	0.0008	0.0086	0.0003	1.2500	0.0040	140	1	141	2	163	38
Pm27-49-02	21	931	0.0218	0.0002	0.1479	0.0043	0.0492	0.0014	0.0090	0.0003	0.3650	0.0000	139	1	140	4	158	65
Pm27-49-03	28	1157	0.0213	0.0002	0.1467	0.0059	0.0500	0.0019	0.0127	0.0004	0.4040	0.0020	136	1	139	6	194	90
Pm27-49-04	28	1199	0.0217	0.0002	0.1482	0.0144	0.0495	0.0047	0.0076	0.0002	0.4950	0.0000	138	1	140	14	173	224
Pm27-49-05	23	1022	0.0219	0.0002	0.1490	0.0043	0.0493	0.0014	0.0073	0.0002	0.4420	0.0010	140	1	141	4	161	65
Pm27-49-06	23	1003	0.0218	0.0002	0.1467	0.0037	0.0488	0.0012	0.0069	0.0002	0.4830	0.0030	139	1	139	4	140	57
Pm27-49-07	23	976	0.0220	0.0002	0.1502	0.0087	0.0495	0.0028	0.0075	0.0002	0.5020	0.0000	140	1	142	8	170	133
Pm27-49-08	30	1337	0.0219	0.0002	0.1491	0.0036	0.0493	0.0011	0.0067	0.0002	0.4680	0.0010	140	1	141	3	164	53
Pm27-49-09	29	990	0.0244	0.0003	0.3983	0.0157	0.1182	0.0037	0.0134	0.0006	0.4780	0.0020	156	2	340	13	1929	57
Pm27-49-10	29	1278	0.0219	0.0002	0.1498	0.0051	0.0496	0.0017	0.0071	0.0002	0.4990	0.0010	140	1	142	5	176	78
Pm27-49-11	35	1465	0.0219	0.0002	0.1482	0.0052	0.0490	0.0017	0.0078	0.0002	0.5620	0.0010	140	1	140	5	149	79
Pm27-49-12	58	1887	0.0237	0.0002	0.3408	0.0118	0.1043	0.0030	0.0100	0.0003	0.9070	0.0010	151	2	298	10	1703	54
Pm27-49-13	18	805	0.0219	0.0002	0.1473	0.0147	0.0488	0.0047	0.0070	0.0003	0.4160	0.0010	139	1	140	14	140	228
Pm27-49-14	35	1434	0.0219	0.0002	0.1478	0.0032	0.0489	0.0010	0.0080	0.0002	0.6440	0.0000	140	1	140	3	144	49
Pm27-49-15	25	1037	0.0218	0.0002	0.1478	0.0048	0.0491	0.0015	0.0081	0.0002	0.5420	0.0010	139	1	140	5	153	74
英安斑岩(B1412)																		
B1412-01	3	145	0.0208	0.0003	0.1397	0.0278	0.0488	0.0113	0.0043	0.0004	0.4363	0.0006	133	2	133	26	136	546
B1412-02	9	393	0.0209	0.0002	0.1427	0.0106	0.0494	0.0038	0.0059	0.0001	0.6605	0.0020	134	1	135	10	169	179
B1412-03	4	191	0.0211	0.0003	0.1422	0.0266	0.0488	0.0105	0.0053	0.0004	0.4335	0.0036	135	2	135	25	140	504
B1412-04	6	294	0.0208	0.0002	0.1413	0.0267	0.0492	0.0092	0.0072	0.0004	0.4469	0.0025	133	1	134	25	159	439
B1412-05	3	154	0.0207	0.0003	0.1398	0.0256	0.0490	0.0106	0.0065	0.0005	0.3389	0.0026	132	2	133	24	146	508
B1412-06	5	227	0.0207	0.0002	0.1388	0.0268	0.0486	0.0099	0.0065	0.0004	0.5014	0.0030	132	2	132	25	131	477
B1412-07	9	414	0.0207	0.0002	0.1398	0.0119	0.0489	0.0042	0.0072	0.0002	0.3660	0.0020	132	1	133	11	141	201
B1412-08	4	170	0.0207	0.0004	0.1394	0.0486	0.0488	0.0193	0.0083	0.0005	0.5733	0.0027	132	3	133	46	141	927
B1412-09	7	339	0.0212	0.0002	0.1408	0.0117	0.0481	0.0040	0.0078	0.0003	0.3765	0.0011	136	1	134	11	103	195
B1412-10	33	1376	0.0209	0.0001	0.1396	0.0038	0.0484	0.0013	0.0057	0.0000	0.9797	0.0007	133	1	133	4	120	66
B1412-11	1	65	0.0212	0.0009	0.1424	0.0910	0.0487	0.0000	0.0011	0.0014	0.3968	0.0023	135	6	135	86	135	-
B1412-12	6	294	0.0209	0.0002	0.1415	0.0208	0.0490	0.0072	0.0063	0.0002	0.5130	0.0004	134	1	134	20	148	345
B1412-13	7	317	0.0212	0.0002	0.1410	0.0123	0.0483	0.0042	0.0050	0.0001	0.7301	0.0004	135	1	134	12	115	206
B1412-14	4	191	0.0209	0.0003	0.1396	0.0258	0.0484	0.0102	0.0041	0.0003	0.3901	0.0017	134	2	133	25	116	497
B1412-15	6	267	0.0209	0.0003	0.1416	0.0217	0.0491	0.0077	0.0051	0.0003	0.4880	0.0040	133	2	134	21	155	365
B1412-16	6	296	0.0210	0.0002	0.1416	0.0237	0.0490	0.0085	0.0040	0.0001	0.8182	0.0015	134	1	134	23	147	405

其他样品,可能与分析样品里含有较多的钾长石晶屑有关。其余样品的SiO₂含量为73.39%~77.26%; Na₂O为3.29%~4.20%,平均为3.80%; K₂O为3.44%~5.05%,平均为4.14%; (Na₂O + K₂O)为7.65%~8.34%, Na₂O/K₂O为0.65~1.22,平均0.96。 Al₂O₃、FeO^T、CaO和MgO含量较低,其中Al₂O₃为11.58%~12.29%,平均为12.02%; FeO^T为1.03%~1.44%,平均为1.18%; CaO为0.27%~0.85%,平均为0.47%; MgO为0.04%~0.08%,平均为0.06%, Mg[#]为4.69~12.23; A/CNK介于1.00~1.07,平均为1.04,为弱过铝质高钾钙碱性系列(图5a)。分析结果(表2)显示,白音高老组火山岩部分样品烧失量较大(B1404),可能遭受蚀变作用,K、Na等活性较强的元素可能有一定程度的变化,本文采用Nb/Y-Zr/TiO₂微量元素

分类图解,样品落入到流纹岩范围(图5b)。

次火山岩与酸性火山岩相比具有较低的SiO₂含量,为67.07%~67.96%;较高的Na₂O+K₂O、Al₂O₃、FeO^T、CaO和MgO含量,分别为8.36%~9.80%、15.98%~16.23%、2.88%~3.30%、0.34%~1.64%和0.18%~0.30%,以及较高的A/CNK和Mg[#]值,分别为1.15~1.20和8.89~15.47,为过铝质高钾钙碱性系列(图5a)。在Nb/Y-Zr/TiO₂图中,样品落入到流纹英安岩和英安岩范围(图5b)。在哈克图解中(图略),次火山岩和酸性火山岩具有一定演化性,随着SiO₂含量的增加,TiO₂、FeO^T、CaO、MgO和P₂O₅含量下降,Na₂O、Al₂O₃变化不明显,K₂O随SiO₂含量的增加而增加。

4.3 稀土元素

白音高老组火山岩稀土元素分析结果见表2。

表2 罕乌拉地区白音高老组火山岩主量元素(%)和微量元素(10⁻⁶)分析结果

Table2 Major(%) and trace element (10⁻⁶) analysis results of volcanic rocks of Baiyingaolao Formation in Hanwula

样品号	B1404	B1408	Pm13 - 21	Pm27 - 49	B1412	Pm21 - 88	Pm21 - 137	样品号	B1404	B1408	Pm13 - 21	Pm27 - 49	B1412	Pm21 - 88	Pm21 - 137
岩性	流纹岩	流纹质晶屑熔结凝灰岩	球粒流纹岩	球粒流纹岩	英安斑岩	英安斑岩	英安斑岩	岩性	流纹岩	流纹质晶屑熔结凝灰岩	球粒流纹岩	球粒流纹岩	英安斑岩	英安斑岩	英安斑岩
SiO ₂	73.39	77.26	80.37	77.18	67.96	67.07	67.72	Co	0.26	0.54	0.33	1.48	1.29	1.92	2.57
TiO ₂	0.08	0.08	0.04	0.09	0.37	0.44	0.47	Ni	1.00	1.60	1.80	21.90	2.00	2.10	2.90
Al ₂ O ₃	12.29	12.21	9.27	11.58	15.98	16.23	15.99	Li	33.15	13.36	88.70	20.03	15.75	39.72	32.94
FeO ^T	1.06	1.03	0.80	1.44	3.02	3.30	2.88	Sc	3.35	2.94	8.99	2.06	7.00	7.40	6.90
Fe ₂ O ₃	0.68	1.00	0.82	1.25	2.70	3.17	2.73	La	32.65	36.78	5.87	44.83	25.30	22.49	23.48
FeO	0.46	0.14	0.07	0.33	0.62	0.48	0.46	Ce	69.26	66.99	12.39	80.45	55.96	50.99	59.10
MnO	0.02	0.02	0.01	0.02	0.03	0.02	0.02	Pr	7.93	8.26	1.83	12.87	7.00	6.85	7.16
MgO	0.06	0.08	0.07	0.04	0.26	0.18	0.30	Nd	29.69	30.11	7.76	47.69	28.26	28.01	29.09
CaO	0.85	0.27	0.13	0.27	0.34	1.23	1.64	Sm	5.89	5.89	2.77	9.97	5.50	5.49	5.71
Na ₂ O	3.92	3.29	0.10	4.20	4.23	3.57	3.31	Eu	0.17	0.15	0.05	0.10	1.43	1.65	1.68
K ₂ O	3.93	5.05	7.69	3.44	5.57	5.04	5.05	Gd	5.00	5.07	2.47	8.05	4.63	4.62	4.62
P ₂ O ₅	0.01	0.01	0.02	0.03	0.09	0.12	0.13	Tb	0.89	0.90	0.54	1.36	0.74	0.77	0.77
LOI	4.26	0.53	1.29	1.47	1.62	2.21	1.92	Dy	4.79	5.02	3.29	7.35	3.82	3.99	4.10
Total	99.95	99.95	99.88	99.90	99.77	99.76	99.74	Ho	0.91	0.97	0.63	1.32	0.73	0.75	0.76
Cs	86.29	4.75	3.62	4.51	6.57	2.00	5.23	Er	2.72	2.93	1.93	3.79	2.18	2.18	2.19
Rb	228.80	234.90	193.50	131.50	109.40	94.20	105.40	Tm	0.45	0.49	0.36	0.60	0.35	0.34	0.35
Sr	64.10	24.50	24.80	16.10	77.30	171.40	226.10	Yb	2.99	3.29	2.68	4.19	2.40	2.31	2.31
Ba	99.20	115.60	71.50	134.10	1447.00	1373.00	1431.00	Lu	0.56	0.57	0.47	0.83	0.63	0.53	0.56
Ga	23.80	24.70	16.55	26.68	28.20	24.20	23.30	Y	28.36	31.38	16.47	34.73	20.96	21.12	20.19
Nb	11.82	8.65	16.77	14.19	9.42	8.18	7.86	Mg [#]	9.10	12.23	13.42	4.69	13.27	8.89	15.47
Ta	1.29	1.15	1.52	1.24	0.75	0.62	0.64	REE	163.93	167.43	59.50	223.38	138.92	130.96	141.89
Zr	145.70	108.40	69.50	244.60	341.40	304.90	305.40	δEu	0.09	0.08	0.05	0.03	0.86	1.00	1.00
Hf	6.14	4.77	3.92	7.80	8.24	8.75	8.42	Nb/Ta	9.14	7.51	11.04	11.46	12.49	13.18	12.36
Th	25.37	21.73	14.57	22.17	9.72	7.66	7.46	Rb/Sr	3.57	9.59	7.80	8.17	1.42	0.55	0.47
V	7.50	9.30	26.90	19.20	16.40	18.60	23.20	T ^{°C}	778.65	762.00	731.47	831.77	862.96	853.86	850.65
Cr	4.70	2.70	14.80	14.60	3.60	4.00	5.60	DI	94.03	96.40	97.72	96.41	91.57	86.54	85.23

注: Mg[#] = 100 × (Mg²⁺ / (Mg²⁺ + TFe²⁺)); δEu = Eu_n / (Sm_n × Gd_n)^{1/2}; T^{°C} = 12,900 / [2.95 + 0.85 × M + ln(496,000/Zr)] - 273.15, M = (Na + K + 2Ca) / (Al × Si); DI = Q + Or + Ab + Ne + Lc + Kp_o.

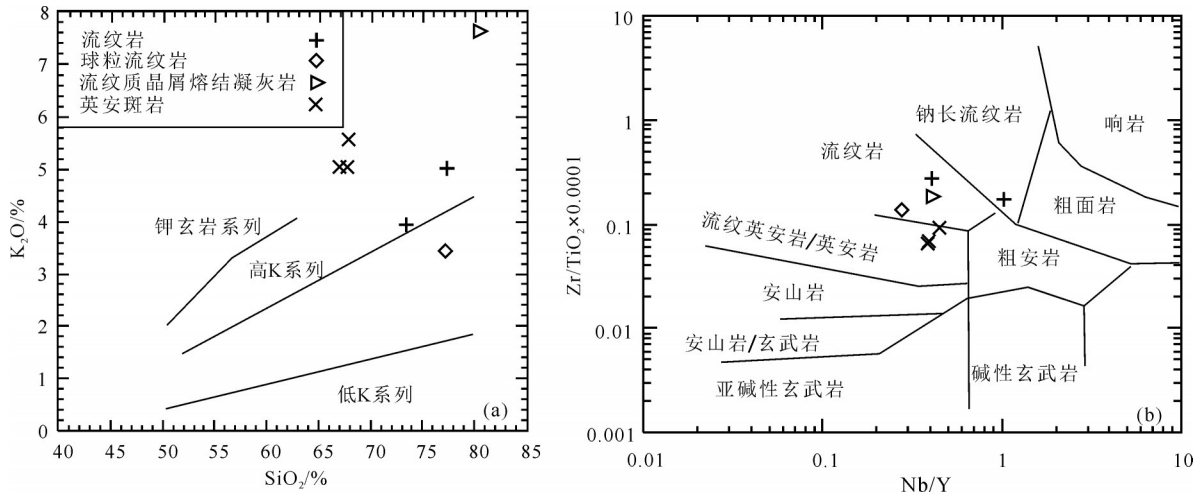


图5 罕乌拉地区白音高老组火山岩 SiO₂-K₂O图解(a)(转引自 Rollinson, 1993)和Nb/Y-Zr/TiO₂分类命名图解(b) (转引自 Wilson, 1989)

Fig.5 SiO₂-K₂O diagrams (a, after Rollinson, 1993) and Nb/Y-Zr/TiO₂ classifying-naming diagrams (b, after Wilson, 1989) for volcanic rocks of Baiyingaolao Formation in Hanwula

白音高老组酸性火山岩的稀土元素总量(ΣREE)为 43.03×10⁻⁶~233.38×10⁻⁶, 样品Pm13-21(流纹质晶屑熔结凝灰岩)含量明显低于其他样品, 为 43.03×10⁻⁶, 可能与样品含有较多石英、长石等晶屑有关。其余样品的稀土元素总量(ΣREE)为 163.93×10⁻⁶~233.38×10⁻⁶; LREE/HREE 比值为 7.13~7.94, 平均为

7.59; (La/Yb)_N 比值为 7.21~7.53, 平均为 7.37。在球粒陨石标准化稀土元素配分图解上(图6a), 稀土元素具有一定分馏, 富集轻稀土元素, 亏损重稀土元素。Yb 含量在 2.99×10⁻⁶~4.19×10⁻⁶, Lu 含量在 0.56×10⁻⁶~0.83×10⁻⁶, (Yb/Lu)_N 比值为 5.08~5.74, 平均为 5.37; (Gd/Yb)_N 比值为 1.24~1.55, 平均为 1.38,

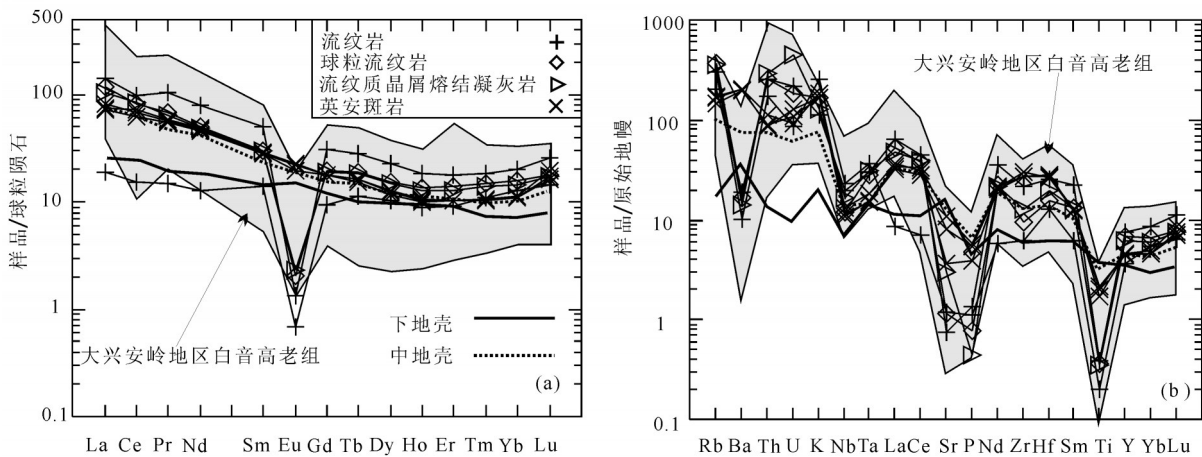


图6 罕乌拉地区白音高老组火山岩球粒陨石标准化稀土元素配分图(a)(球粒陨石标准化数据引自 Boynton, 1984)和原始地幔标准化蛛网图(b)(原始地幔标准化数据引自 Sun and Mc Donough, 1989)

下地壳、中地壳据 Rudnick et al., 2003; 大兴安岭地区白音高老组火山岩数据据 苟军等, 2010; Dong et al., 2014; Kong et al., 2014; 秦涛等, 2014; 聂立军等, 2015; 王雄等, 2015; Yang et al., 2015; 张乐彤等, 2015; 张学斌等, 2015

Fig.6 Chondrite-normalized REE patterns (a) (normalization values after Boynton, 1984) and primitive mantle-normalized trace element spider diagrams (b) (normalization values after Sun and Mc Donough, 1989) for volcanic rocks of Baiyingaolao Formation in Hanwula lower crust, middle Crust (after Rudnick et al., 2003); data of Baiyingaolao Formation in the Da Hinggan Mountains (after Gou Jun et al., 2010; Dong Yu et al., 2014; Kong Yuanming et al., 2014; Qin Tao et al., 2014; Nie Lijun et al., 2015; Wang Xiong et al., 2015; Yang Wubin et al., 2015; Zhang Letong et al., 2015; Zhang Xuebin et al., 2015)

表现出平坦的重稀土元素配分模式。样品 δEu 介于 0.03~0.09, 显示明显 Eu 负异常。

次火山岩的稀土元素总量(REE)低于酸性火山岩, 为 $130.96 \times 10^{-6} \sim 141.39 \times 10^{-6}$; $(\text{La}/\text{Yb})_N$ 比值低于酸性火山岩, 为 6.57~7.11; $(\text{Gd}/\text{Yb})_N$ 比值高于酸性火山岩, 为 1.56~1.62。整体上与酸性火山岩相似, 稀土元素具有一定分馏, 富集轻稀土元素, 亏损重稀土元素(图 6a)。样品 δEu 介于 0.86~1.00, 具有轻微的 Eu 负异常。

4.4 微量元素

白音高老组火山岩微量元素分析结果见表 2。在原始地幔标准化微量元素蛛网图中(图 6b), 白音高老组火山岩整体表现为富集 Rb、Th、Cs 等大离子亲石元素(LILE), 亏损 Sr 和 Nb、Ta、P、Ti 等高场强元素(HFSE)。但酸性火山岩与次火山岩具有一定区别, 酸性火山岩明显亏损 Ba、Sr、P、Ti 元素, 而次火山岩具有 Ba、Zr、Hf 的正异常。酸性火山岩的 Rb 含量为 $131.50 \times 10^{-6} \sim 234.90 \times 10^{-6}$, Sr 含量为 $16.13 \times 10^{-6} \sim 64.12 \times 10^{-6}$, Y 含量介于 $28.36 \times 10^{-6} \sim 34.73 \times 10^{-6}$, Rb/Sr 和 Sr/Y 比值分别为 3.57~9.59 和 0.46~2.26; Nb 含量为 $8.65 \times 10^{-6} \sim 14.19 \times 10^{-6}$, Ta 含量为 $1.15 \times 10^{-6} \sim 1.29 \times 10^{-6}$, Nb/Ta 比值为 7.51~11.46, 平均为 9.37; Cr 含量变化范围较大, 介于 $2.68 \times 10^{-6} \sim 14.60 \times 10^{-6}$, 平均为 7.33×10^{-6} ; Ni 含量介于 $1.02 \times 10^{-6} \sim 21.94 \times 10^{-6}$, 平均为 8.18×10^{-6} 。

次火山岩 Rb/Sr 比值小于酸性火山岩, 为 0.47~1.42; Sr/Y 和 Nb/Ta 比值大于酸性火山岩, 分别为 3.69~11.20 和 12.36~13.18; Cr 和 Ni 含量较为稳定, 分别为 $3.58 \times 10^{-6} \sim 5.55 \times 10^{-6}$ 和 $1.96 \times 10^{-6} \sim 2.87 \times 10^{-6}$ 。

5 讨 论

5.1 形成时代

内蒙古自治区地质局区域地质测量队在 1976 年 1:20 万罕乌拉幅区调中将该套火山岩划归为晚侏罗世^①, 缺少精确的同位素年代学及岩石地球化学资料。本文研究的西乌旗罕乌拉地区白音高老组喷溢相球粒流纹岩和次火山岩相英安斑岩中的锆石多呈自形, 内部结构清晰, 生长震荡环带发育, Th/U 值较高(0.34~1.25), 表明锆石为岩浆成因, 因此所测得的年龄代表岩浆结晶年龄。球粒流纹岩的形成时代为 $(140 \pm 0.8) \text{Ma}$ (MSWD=0.13), 英安斑

岩的形成时代为 $(133 \pm 0.7) \text{Ma}$ (MSWD=0.49), 该区白音高老组火山岩的形成时代介于 140~133 Ma, 应为早白垩世早期。对大兴安岭地区白音高老组测年数据进行归纳整理(表 3), 60 个年龄数据表明其形成年龄介于 144~121 Ma, 集中于 141~124 Ma, 整体上没有明显的间断(图 7), 与本文结果相符。在球粒流纹岩样品中见有捕获锆石(9、12 号), 锆石特征显示为岩浆成因, 其年龄值分别为 $(156 \pm 2) \text{Ma}$ 、 $(151 \pm 2) \text{Ma}$, 区域研究表明大兴安岭地区满克头鄂博组火山岩年龄峰值集中在 160~150 Ma (张吉衡, 2009; 王建国等, 2013; 李鹏川等, 2016), 这 2 颗锆石应代表该区满克头鄂博组火山岩的年龄信息。

5.2 岩石成因与源区

研究区白音高老组火山岩具有较高的 SiO_2 和 $(\text{Na}_2\text{O} + \text{K}_2\text{O})$ 含量, 较低的 CaO 和 MgO 含量; FeO^T/MgO 比值(平均为 17.74)远高于一般 I 型花岗岩(平均为 2.27, Whalen et al., 1987), 接近世界 A 型花岗岩的平均值(22.84, Whalen et al., 1987); 具较高的 Ga、Zr、Nb 和 Y 含量, 较低的 Sr 和 Ba 含量, $10000 \times \text{Ga}/\text{Al}$ 比值介于 2.75~4.35(平均 3.46), 明显高于 I 型和 S 型花岗岩平均值(分别为 2.1 和 2.28), 接近 A 型花岗岩(3.75, Whalen et al., 1987)。酸性火山岩 $\text{Zr} + \text{Nb} + \text{Ce} + \text{Y}$ 平均值为 281.50×10^{-6} , 稍低于 A 型花岗岩, 次火山岩 $\text{Zr} + \text{Nb} + \text{Ce} + \text{Y}$ 平均为 341.68×10^{-6} , 接近 A 型花岗岩的下限值(350×10^{-6})。白音高老组火山岩 Nb/Ta 平均值为 11.02, 高于高分异 I 型花岗岩(2.3~9.9)。相对于 A 型花岗岩, 高分异 I 型花岗岩具有较低的 FeO^T 含量($< 1.00\%$)和较高的 Rb 含量($> 270 \times 10^{-6}$), 而白音高老组火山岩具有较高的 FeO^T 含量(平均为 2.12%)和较低的 Rb 含量(平均为 150.70×10^{-6}), 区别于高分异 I 型花岗岩(Whalen et al., 1987)。据 Watson et al.(1983)的公式, 用岩石主要元素和 Zr 含量计算样品的锆石饱和温度, 白音高老组次火山岩的锆石饱和温度为 851~863°C, 较高的成岩温度同样不支持它们为 I 型(764°C; King et al., 1997)。在 $10000 \times \text{Ga}/\text{Al} - \text{Nb}$ 和 $10000 \times \text{Ga}/\text{Al} - \text{Zr}$ 判别图解中(图 8), 所有样品点都投影于 A 型花岗岩区。因此, 综上所述, 白音高老组火山岩应为 A 型。

地球化学特征显示, 白音高老组次火山岩和酸性火山岩具有一定差别。次火山岩的分异指数 DI (85.23~91.57) 低于酸性火山岩(94.03~96.41), 但结

表3 大兴安岭地区白音高老组火山岩年龄测定结果
Table3 The ages of volcanic rocks of Baiyingaolao Formation in Da Hinggan Mountains

序号	地区	岩性	测试方法	纬度	经度	年龄/ Ma	资料来源
1	西乌旗罕乌拉	球粒流纹岩	LA-ICP-MS 锆石 U-Pb	44°42'08"	118°55'56"	140±0.8	本文
2	西乌旗罕乌拉	英安斑岩	LA-ICP-MS 锆石 U-Pb	44°42'54"	118°37'03"	133±0.7	本文
3	赤峰	流纹岩	LA-ICP-MS 锆石 U-Pb	42°51'43"	118°06'04"	132±1.0	郝彬等,2016
4	赤峰	流纹岩	LA-ICP-MS 锆石 U-Pb	42°57'08"	118°04'42"	130±1.0	郝彬等,2016
5	兴安地块呼伦湖东岸	流纹岩	LA-ICP-MS 锆石 U-Pb	49°11'36"	117°52'14"	144±0.6	黄明达等,2016
6	塔尔气	流纹岩	LA-ICP-MS 锆石 U-Pb			132±2.0	李杰等,2016
7	二十四道沟盆地	流纹岩	LA-ICP-MS 锆石 U-Pb	47°32'35"	121°42'11"	131±2.0	司秋亮等,2016
8	敖尼尔林场东	流纹岩	LA-ICP-MS 锆石 U-Pb	47°45'50"	121°34'25"	136±1.0	司秋亮等,2016
9	伊勒呼里山	流纹质凝灰岩	LA-ICP-MS 锆石 U-Pb	51°17'01"	123°49'57"	124±1.1	尹志刚等,2016
10	伊勒呼里山	流纹质玻屑凝灰岩	LA-ICP-MS 锆石 U-Pb	51°22'46"	123°51'10"	125±0.7	尹志刚等,2016
11	扎兰屯韩家地窝棚	流纹岩	LA-ICP-MS 锆石 U-Pb	47°54'46"	122°22'07"	130±5.0	杜玉春,2015
12	柴河源林场天然砂子	流纹岩	LA-ICP-MS 锆石 U-Pb	47°34'22"	120°43'53"	127±0.5	聂立军等,2015
13	吐列毛杜	流纹岩	LA-ICP-MS 锆石 U-Pb	45°31'10"	120°44'11"	127±2.1	Yang et al.,2015
14	柴河	流纹岩	SHIRMP 锆石 U-Pb			131±1.0	张乐彤等,2015
15	锡林浩特	流纹岩	SHIRMP 锆石 U-Pb			134±0.8	张乐彤等,2015
16	锡林浩特	石泡流纹岩	SHIRMP 锆石 U-Pb			137±1.2	张乐彤等,2015
17	锡林浩特	碎斑熔岩	SHIRMP 锆石 U-Pb			139±0.9	张乐彤等,2015
18	五岔沟	流纹岩	LA-ICP-MS 锆石 U-Pb	46°57'40"	120°22'04"	134±1.0	Dong et al.,2014
19	五岔沟	流纹质凝灰岩	LA-ICP-MS 锆石 U-Pb	46°48'11"	120°22'11"	133±2.0	Dong et al.,2014
20	五岔沟	流纹岩	LA-ICP-MS 锆石 U-Pb	48°35'57"	120°22'02"	133±1.0	Dong et al.,2014
21	五岔沟	流纹质凝灰岩	LA-ICP-MS 锆石 U-Pb	46°46'50"	120°17'07"	132±2.0	Dong et al.,2014
22	五岔沟	流纹岩	LA-ICP-MS 锆石 U-Pb	46°46'38"	120°17'49"	130±3.0	Dong et al.,2014
23	乌兰哈拉盖	流纹质晶屑凝灰岩	LA-ICP-MS 锆石 U-Pb			131±0.5	黄猛,2014
24	西乌旗	流纹质熔结凝灰岩	LA-ICP-MS 锆石 U-Pb			135±0.8	黄猛,2014
25	西乌旗	流纹质角砾熔结凝灰岩	LA-ICP-MS 锆石 U-Pb			134±0.9	黄猛,2014
26	科右中旗	流纹质玻屑岩屑晶屑凝灰熔岩	LA-ICP-MS 锆石 U-Pb	45°34'51"	120°35'52"	125±1.0	孔元明,2014
27	科右中旗	流纹岩	LA-ICP-MS 锆石 U-Pb			122±1.0	Kong et al.,2014
28	嫩江县	流纹岩	LA-ICP-MS 锆石 U-Pb			128 ± 0.8	刘阁等,2014
29	苏尼特左旗	石英粗安岩	LA-ICP-MS 锆石 U-Pb	44°37'07.8"	112°20'33.1"	135±5.6	秦旭亮,2014
30	苏尼特左旗	粗安岩	LA-ICP-MS 锆石 U-Pb	44°37'29.8"	112°24'37.7"	135±3.8	秦旭亮,2014
31	大石门林场	流纹岩	LA-ICP-MS 锆石 U-Pb			126±0.94	秦涛等,2014
32	大旗村	流纹岩	LA-ICP-MS 锆石 U-Pb			130±2.0	秦涛等,2014
33	扎兰屯地区	流纹岩	LA-ICP-MS 锆石 U-Pb			130 ± 5.0	张亚明等,2014
34	西哲里木苏木	钠闪石流纹岩	LA-ICP-MS 锆石 U-Pb			141±1.0	王建国等,2013
35	东乌旗辉音敖包	流纹质熔结凝灰岩	SHIRMP 锆石 U-Pb			135±2.0	陈英富等,2012
36	东乌旗辉音敖包	流纹岩	SHIRMP 锆石 U-Pb			129±3.0	陈英富等,2012
37	东乌旗辉音敖包	石泡流纹岩	SHIRMP 锆石 U-Pb			124±0.5	陈英富等,2012
38	下营子	流纹岩	SHIRMP 锆石 U-Pb	44°03'53"	117°46'48"	144±1.4	李可等,2012
39	赤峰	粗安质熔结凝灰岩	LA-ICP-MS 锆石 U-Pb	42°56'09"	118°03'24"	132±1.0	杨杨等,2012
40	赤峰	流纹岩	LA-ICP-MS 锆石 U-Pb	43°00'13"	118°29'34"	138±3.0	杨杨等,2012
41	阿尔山	流纹岩	LA-ICP-MS 锆石 U-Pb			125±2.0	陈良,2010
42	阿尔山	流纹质凝灰熔岩	LA-ICP-MS 锆石 U-Pb			125±1.0	陈良,2010
43	五岔沟	熔结凝灰岩	LA-ICP-MS 锆石 U-Pb	46°45'38"	120°04'16"	126±1.0	方红薇,2010

续表3

序号	地区	岩性	测试方法	纬度	经度	年龄/Ma	资料来源
44	五岔沟	流纹岩	LA-ICP-MS 锆石 U-Pb	46°45'38"	120°04'16"	125±2.0	方红薇,2010
45	五岔沟	流纹质凝灰熔岩	LA-ICP-MS 锆石 U-Pb	46°45'38"	120°04'16"	125±1.0	方红薇,2010
46	满洲里南部	流纹岩	LA-ICP-MS 锆石 U-Pb	48°37'13"	116°49'02"	141±1.0	苟军等,2010
47	满洲里南部	流纹岩	LA-ICP-MS 锆石 U-Pb	49°19'48"	117°29'46"	139±1.0	苟军等,2010
48	满洲里南部	流纹岩	LA-ICP-MS 锆石 U-Pb	49°32'34"	116°44'41"	141±1.0	苟军等,2010
49	满洲里南部	流纹岩	LA-ICP-MS 锆石 U-Pb	49°30'07"	117°11'28"	141±1.0	苟军等,2010
50	林东	流纹岩	LA-ICP-MS 锆石 U-Pb	44°10'57"	119°15'37"	125±4.0	Ying et al., 2010
51	阿尔山	流纹岩	LA-ICP-MS 锆石 U-Pb			126±1.2	张吉衡,2009
52	灰通河	流纹岩	LA-ICP-MS 锆石 U-Pb	44°16'44"	118°35'29"	131±1.0	张吉衡,2009
53	阿勒坦大阪	流纹岩	LA-ICP-MS 锆石 U-Pb	48°19'48"	120°01'08"	129±1.0	张吉衡,2009
54	巴雅尔图胡硕	流纹岩	LA-ICP-MS 锆石 U-Pb	45°05'32"	120°22'39"	132±1.0	张吉衡,2009
55	沙仁台	流纹岩	LA-ICP-MS 锆石 U-Pb	46°21'18"	120°34'21"	134±1.0	张吉衡,2009
56	阿尔山	流纹岩	LA-ICP-MS 锆石 U-Pb	47°09'07"	119°56'33"	124±1.0	张吉衡,2009
57	乌兰哈德	凝灰熔岩	LA-ICP-MS 锆石 U-Pb	45°39'20"	120°47'39"	124±1.0	张吉衡,2009
58	中心屯	熔结凝灰熔岩	LA-ICP-MS 锆石 U-Pb	46°07'56"	121°10'23"	136±1.0	张吉衡,2009
59	白音高老	熔结凝灰熔岩	LA-ICP-MS 锆石 U-Pb	43°55'59"	119°17'41"	139±1.0	张吉衡,2009
60	上伙房	凝灰熔岩	LA-ICP-MS 锆石 U-Pb	43°07'54"	117°41'37"	141±1.0	张吉衡,2009

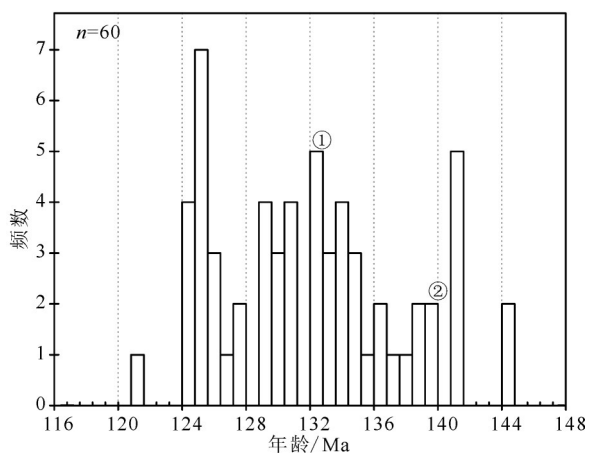


图7 大兴安岭地区白音高老组火山岩年龄直方图(图中白音高老组年龄数据见表3;①、②分别代表本文133、140 Ma火山岩所处年龄段)

Fig.7 Age probability diagram for volcanic rocks of Baiyingaolao Formation in the Da Hinggan Mountains(ages of Baiyingaolao Formation after Table 3; ① and ② representing 133, 140Ma of this paper, respectively)

晶温度高于酸性火山岩(分别为851~863℃和731~832℃)。相对于次火山岩,酸性火山岩具有更高的REE和Rb含量以及Rb/Sr比值。在Harker图解中(图略),次火山岩和酸性火山岩具一定变化趋势,显示酸性火山岩可能为次火山岩分异结晶的结果。但是,酸性火山岩和次火山岩成分上存在一定

的间断,并且次火山岩具有较高的钾含量,靠近或属于碱性系列,而酸性火山岩为高钾钙碱性系列,酸性火山岩具有比次火山岩明显的Sr、Ba和Ti异常,不符合结晶分异的结果。并且,如果两者为结晶分异的关系,次火山岩(英安斑岩)的年龄应早于酸性火山岩(流纹岩),这与研究结果不符。以上说明,次火山岩和酸性火山岩应为同一岩浆房,不同期次作用的结果。由于酸性火山岩分异程度较高,经过斜长石、钾长石等矿物分离结晶,具有较高的SiO₂和Rb含量以及Rb/Sr比值,较低的TiO₂、FeO^T、MgO和CaO含量,明显的Sr、Ba、Eu和Ti负异常以及较低的结晶温度。

白音高老组火山岩具有较低的Nb/Ta比值和Mg[#]值,相对富集轻稀土元素,亏损重稀土元素,大部分样品LILE富集,而Ba、Sr和HFSE亏损等特征,显示具壳源特征(Rudnick et al., 2014)。其Rb/Sr(原始地幔0.3)和Ti/Zr(地壳<20)比值与壳源岩浆近似,与幔源岩浆明显不同(Sun et al., 1989; Pearce, 1983),进一步说明白音高老组火山岩具有一定壳源特征。

研究区白音高老组火山岩较低的Sr/Y和(La/Yb)_N比值和高的重稀土含量,表明源区熔融时残留相中缺乏石榴石(Defant et al., 1990);另外白音高老组火山岩具有较低的Sr含量,并在微量元素蛛

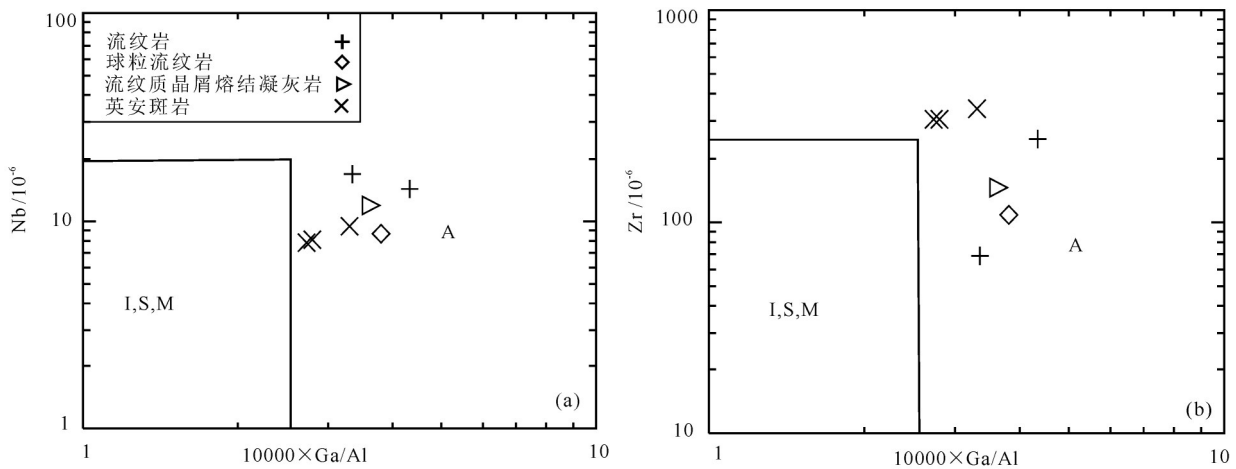


图8 罕乌拉地区白音高老组火山岩 10000×Ga/Al-Nb 图解(a)和 10000×Ga/Al-Zr 图解(b)(a,b,据 Whalen et al., 1987)
 Fig.8 Diagrams of 10000×Ga/Al-Nb(a)and 10000×Ga/Al-Zr(b)(a, b, after Whalen et al., 1987) for volcanic rocks of Baiyingaolao Formation in Hanwula

网图上具明显的 Ba 和 Sr 的负异常,这暗示源区残留相中可能存在斜长石。石榴石具有较高的稳定压力(大于 1.0~1.2 GPa),而当压力大于 1.5 GPa,斜长石会变得不稳定(Sen et al., 1994),因此白音高老组火山岩形成压力较低,应为正常厚度地壳部分熔融结果。大兴安岭地区白音高老组火山岩同位素研究显示,其具有较高的火山岩 $\epsilon_{\text{Hf}}(t)$ 和 $\epsilon_{\text{Nd}}(t)$ 值(张玉涛等, 2007; Zhang et al., 2008; 苟军等, 2010; Kong et al., 2014)。大兴安岭地区经过早期大洋的俯冲及其幔源岩浆的底侵作用,形成新生地壳,由于后碰撞阶段的拉伸作用使这些新生地壳部分熔融,形成了白音高老组火山岩,白音高老组火山岩的 Sr 和 Hf 同位素特征正是源区特征的体现。所以白音高老组火山岩可能为新生镁铁质地壳部分熔融的结果。

5.3 火山岩形成的构造环境探讨

Eby(1990, 1992)将 A 型花岗岩划分为 A_1 和 A_2 两种类型,并指出 A_1 型花岗岩产于与上地幔热柱、裂谷作用有关的非造山环境, A_2 型花岗岩主要产于与大陆边缘地壳伸展作用或与陆内剪切作用产生的拉张环境有关的后造山环境。利用 Nb-Y-Ce、Nb-Y-3×Ga 及 Y/Nb-Rb/Nb 判别图(图 9),研究区白音高老组火山岩样品基本都位于 A_2 型花岗岩区。目前对于 A 型花岗岩物质来源和成因的解释有着不同的认识,比较统一的观点是其形成于伸展构造环境(Eby, 1990, 1992; Frost et al., 2001; Bonin, 2007)。结合同时期大兴安岭地区广泛发育的双峰式岩浆作用(葛文春等, 1999; 郭锋等, 2001; Wang et al., 2006; 裴福萍等, 2008; Zhang et al., 2008,

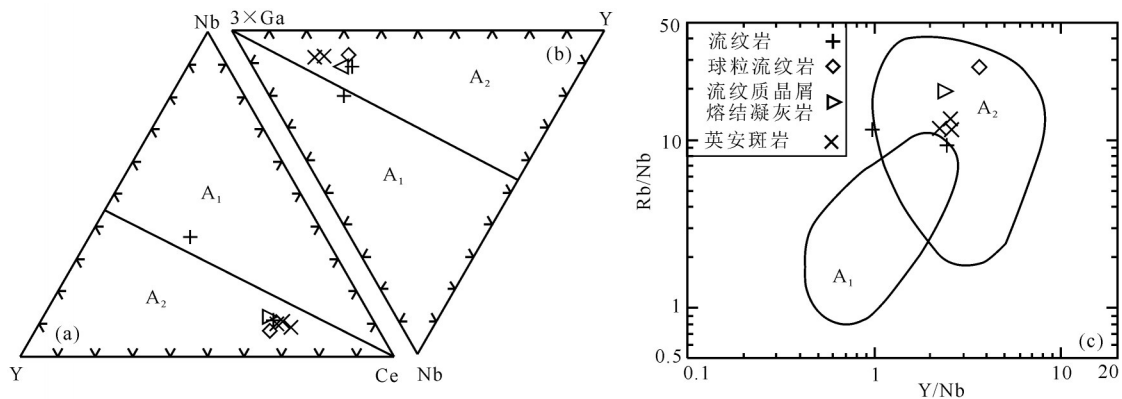


图9 罕乌拉地区白音高老组火山岩构造环境图解(据 Eby, 1992)
 Fig.9 Tectonic discrimination diagrams for volcanic rocks of Baiyingaolao Formation in Hanwula(after Eby, 1992)

2010;孟恩等,2011)、A型花岗岩(孙德有等,2004,2005;Wang et al.,2006;隋振民等,2007;施璐等,2013;李竞妍等,2014;吴涛涛等,2016)和变质核杂岩(张履桥等,1998;Liu et al.,2006,刘俊来等,2008,2011;林伟等,2013),进一步证实大兴安岭地区在早白垩世处于伸展构造背景。但对于这种伸展构造体制是后碰撞阶段(蒙古—鄂霍茨克构造体系),还是俯冲阶段弧后环境(环太平洋构造体系)或者是板内体系还存在较大的争论。

研究显示,大兴安岭地区发育S型花岗岩(170 Ma左右)(李宇等,2015)、具埃达克质特征的岩浆岩(164~140 Ma)(高晓峰等,2005;武广等,2008;何付兵等,2013;黄凡等,2014),结合辽西地区晚侏罗世—早白垩世自北向南推覆事件的存在(张长厚等,2002;赵越等,2004;张宏等,2008;孟恩等,2011),证明在中侏罗世—早白垩世大兴安岭地区存在地壳加厚事件(赵越等,1994),由于加厚陆壳的不稳定性,导致造山后的伸展环境或加厚陆壳的拆沉作用。所以西乌旗罕乌拉地区白音高老组火山岩可能与蒙古—鄂霍茨克洋闭合碰撞后的伸展环境相关。

从整个东北地区中生代火山的时空分布来看,目前在松辽盆地以东地区尚未发现160~138 Ma的火山岩,而早白垩世晚期(133~106 Ma)火山岩在大兴安岭、吉黑东部及松辽盆地底部断陷层广泛分布(Wang et al.,2006;Zhang et al.,2008,2010;孟恩等,2011;徐美君等,2011;许文良等,2013b),且显示出由东部陆缘向西部陆内,火山岩中的碱性组分具有增高的成分极性变化,在靠近陆内一侧的松辽盆地和大兴安岭地区则显示双峰式火山岩组合的特征,表明了早白垩世晚期受到了东部板块俯冲作用的影响(郭锋等,2001;Wang et al.,2006;Zhang et al.,2008,2010;孟恩等,2011;许文良等,2013b)。中国东部在早白垩世广泛出露变质核杂岩、断陷盆地和拆离构造,例如:呼和浩特、达子营和万福等变质核杂岩(关会梅等,2008;申亮等,2011);松辽、大杨树和孙吴等断陷盆地(杨建国等,2006),这些伸展构造具有近似的伸展方向(NW-SE),变质核杂岩的形成时代主体介于130~110 Ma(刘俊来等,2011;林伟等,2013),也印证了位于太平洋西岸的欧亚大陆东缘在早白垩世具伸展性质的构造背景(林伟等,

2013)。以上研究表明,古太平洋板块俯冲作用对中国东部早白垩世晚期火山岩具有一定影响。结合本文白音高老组火山岩球粒流纹岩和英安斑岩的形成时代分别为(140±0.8)Ma和(133±0.7)Ma,所以流纹岩的形成可能主要受控于蒙古—鄂霍茨克洋俯冲/碰撞后的后造山伸展垮塌作用,英安斑岩为后期古太平洋俯冲和蒙古—鄂霍茨克后碰撞阶段共同作用的结果。

6 结 论

(1)内蒙古西乌旗罕乌拉地区白音高老组球粒流纹岩的形成时代为(140±0.8)Ma,英安斑岩的形成时代为(133±0.7)Ma,整体形成于早白垩世早期。

(2)白音高老组火山岩具A型花岗岩地球化学特征,为地壳部分熔融的结果。

(3)西乌旗罕乌拉地区白音高老组火山岩形成于伸展构造环境,为蒙古—鄂霍茨克构造体系和环太平洋构造体系共同作用的结果,流纹岩主要受控于蒙古—鄂霍茨克洋俯冲/碰撞后的后造山伸展垮塌作用,英安斑岩为后期古太平洋俯冲和蒙古—鄂霍茨克后碰撞阶段共同叠加作用的结果。

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注释

①内蒙古自治区地质局区域地质测量队. 1976. 罕乌拉幅(L-30-29)1:20万区域地质调查报告.

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