

doi: 10.12029/gc20200203

胡二红, 贺中银, 张善明, 吕新彪, 阮班晓, 赵鹏彬, 孙浩, 纪鹏飞, 何世明, 苏波, 刘婷. 2020. 内蒙古呼伦西白地区铌稀土多金属矿成矿特征及找矿标志[J]. 中国地质, 47(2): 300–314.

Hu Erhong, He Zhongyin, Zhang Shanming, Lü Xinbiao, Ruan Banxiao, Zhao Pengbin, Sun Hao, Ji Pengfei, He Shiming, Su Bo, Liu Ting. 2020. Metallogenic characteristics and prospecting indicators of Nb–REE polymetallic deposits in Hulunxibai area, Inner Mongolia[J]. *Geology in China*, 47(2): 300–314(in Chinese with English abstract).

内蒙古呼伦西白地区铌稀土多金属矿成矿特征及找矿标志

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摘要:北山成矿带东缘的内蒙古呼伦西白地区近年来发现了多处铌稀土矿床(点), 矿床(点)产于二叠纪二长花岗岩外接触带上, 矿体赋存于滨浅海相碳酸盐岩-变碎屑岩内, 其形态、产状、规模受构造-岩浆活动控制, 并与磁异常对应, 其中灰石山东北铌稀土矿区 Nb₂O₅ 品位 0.02%~0.56%、平均 0.11%, REO 品位 0.13%~2.63%, 平均 0.93%, 辉森乌拉西铌金矿区 Nb₂O₅ 品位 0.01%~0.1%, 与白云鄂博 Nb-REE-Fe 超大型矿床特征类似, 成矿潜力巨大。结合区域地质背景及铌稀土多金属矿成矿特征, 初步认为区内该类型矿床的找矿标志为: 二叠系滨浅海相碳酸盐岩-变碎屑岩地层、构造挤压形成的虚脱空间或韧-脆性转换部位、二叠纪二长花岗岩外接触带、放射性异常梯度带-磁异常位置、岩石蚀变(褐铁矿化、硅化、碳酸盐化等)和石英细网脉发育地段。对该地区铌稀土多金属矿成矿特征和找矿标志的研究有利于拓展北山地区三稀矿产的找矿思路, 也可为此类矿产的找矿方向提供借鉴。

关键词: 铌稀土多金属矿; 成矿特征; 找矿标志; 航磁航放异常; 矿产勘查工程; 呼伦西白; 内蒙古
中图分类号: P618.7 文献标志码: A 文章编号: 1000-3657(2020)02-0300-15

Metallogenic characteristics and prospecting indicators of Nb–REE polymetallic deposits in Hulunxibai area, Inner Mongolia

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Abstract: In recent years, many Nb–REE deposits (ore spots) have been discovered in Hulunxibai area on the eastern margin of

收稿日期: 2019-7-10; 改回日期: 2020-03-09

基金项目: 内蒙古自治区地质勘查基金项目(NMKD2016-06, NMKD2014-10)资助。

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Beishan metallogenic belt. The deposits (ore spot) occur in the external contact zone of Permian adamellite, and the orebodies occur in the littoral-shallow marine carbonate and metamorphic detrital rocks. The morphology, attitude and scale of the orebody are controlled by tectonic-magmatic activity, and orebody position is consistent with the magnetic anomaly. The Nb₂O₅ grade of northeast Huishi Mountain mining area is generally between 0.02% and 0.56%, with an average of 0.11%, and the REO grade is between 0.13% and 2.63%, with an average of 0.93%. The Nb₂O₅ grade of the west Huisenwula mining area is between 0.01% and 0.1%. The metallogenic type is similar to that of the superlarge Nb-REE-Fe deposit in Bayan Obo. The metallogenic potential of Nb-REE deposits in this area is enormous. Combined with regional geological setting and metallogenic characteristics of the Nb-REE polymetallic mineralization zone, the authors hold that prospecting indicators for this type of Nb-REE polymetallic deposits in this area are as follows: i. Permian littoral-shallow marine carbonate and metamorphic detrital rocks strata; ii. free space or transfer sites between ductile and brittle deformation from the tectonic compression; iii. Permian adamellite outer contact zone; iv. aeroradiometric anomaly gradient zone and magnetic anomaly; v. wall-rock alteration (limonitization, silicification, carbonatization, etc); vi. fine quartz network veins development area. The study of the metallogenic characteristics and prospecting indicators of the Nb-REE polymetallic deposits in this area is helpful to promoting the prospecting ideas of rare earth, rare metal and rare-scattered elements mineral resources in Beishan area, and can also provide a reference for the prospecting direction of such deposits.

Keyword: Nb-REE polymetallic deposits (ore spots); metallogenic characteristics; prospecting indicators; aeromagnetic and aeroradiometric anomalies; Hulunxibai; mineral exploration engineering; Inner Mongolia

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Fund Support: Supported by Geological Exploration Fund of Inner Mongolia Autonomous Region Program (No. NMKD2016-06, No. NMKD2014-10).

1 引言

三稀资源是稀土、稀有和稀散资源的统称,是新一代尖端武器、信息技术、节能环保、医药和医疗设备、高端装备制造、新材料、新能源汽车等所需要的功能材料、结构材料和关键性原料,2011年起将三稀资源作为国家支持的重点对象(王登红等,2012,2013,2016a)。目前已探明的三稀金属矿床数量主要集中在中南地区(湖北、湖南、广西、广东、海南)(王登红等,2016b;Wang et al., 2018)。内蒙古也发现了多处三稀元素矿产地,从东至西几乎均有发现,多数因规模小、品位低、综合利用价值低等特点而未进行报道,目前可达工业价值的主要有白云鄂博Fe-REE-Nb多金属矿(白鸽等,1996;刘健等,2009;Chen et al., 2011)、额济纳旗灰石山东北铌多金属矿(刘强等,2018)、锡林浩特石灰窑稀有金属矿(朱京占等,2013)、扎鲁特旗巴尔哲(八〇一)稀有多金属矿(于桂梅等,1980;杨武斌等,2009;陈金勇等,2019)、武川赵井沟铌钽稀有多金属矿(聂风军等,2013;李志丹等,2019)、阿右旗桃花拉山稀有稀土矿(蒋荣良,1989)、额济纳旗东七一山铷钨钼多金属矿(王勇等,2009;张善明等,2014)等等,其中

白云鄂博铁-稀土-铌超大型矿床是世界上已知最大的稀土矿床,也是中国最大的钽矿床(刘健等,2009;杨晓勇等,2015)。北山成矿带也发现了多处稀有稀土矿产,主要类型为碱性花岗岩型、花岗伟晶岩型(张善明等,2014;周会武等,2015a,2015b;贾志磊,2016;胡二红,2018)。

研究区位于北山成矿带东缘,许多学者对北山地区的构造格局、演化、各时代成矿背景、矿产空间分布特征等进行了详细深入的研究,但主要集中在额济纳旗以西的甘新蒙地区(左国朝等,1990a,1990b,2003;聂风军等,2002;何世平等,2002,2005;龚全胜等,2003;杨合群等,2008;杨建国等,2012;李俊健等,2015,2016;Wan et al., 2018),呼伦西白及附近的报道则相对较少,而关于区域三稀类型矿产的报道更是寥寥无几。呼伦西白地区地质工作起始于20世纪50年代,具有找矿突破是在2000年以后的近20年,主要发现了呼伦西白金矿床^①、珠斯楞海尔罕铜金矿床^②及多处金银铜铁多金属矿点^{③④}。关于三稀资源的发现及重视则开始于2007年,鄂尔多斯市三鑫矿业有限责任公司通过挂牌取得额济纳旗灰石山东北铌多金属矿项目^⑤,目前已完成勘探阶段初步找矿成果,发现了3条铌稀

土多金属矿带;内蒙古第一地勘院于2011—2016年间在辉森乌拉西在寻找金的同时,发现蚀变带内赋存铌矿(化)体^⑥,2018—2019年内蒙古地质矿产勘查院在该地区工作时亦发现了多处铌矿点,刘强等(2018)认为该地区铌稀土矿既有沉积变质特征,又有热液叠加的特征,此类型铌稀土矿的发现大大拓展了内蒙西部以及整个北山地区的找矿思路,意义重大。笔者通过对前人资料综合研究、野外分析

证等工作,总结了研究区铌稀土多金属矿成矿特征及找矿标志,旨在对北山地区寻找三稀资源矿产开拓思路、提供借鉴。

2 区域地质背景

据左国朝等(2003)划分方案,研究区位于塔里木板块之敦煌—玉门构造区之雅干南晚二叠世拉分盆地褶皱带(图1a)。呼伦西白地区已发现的铌

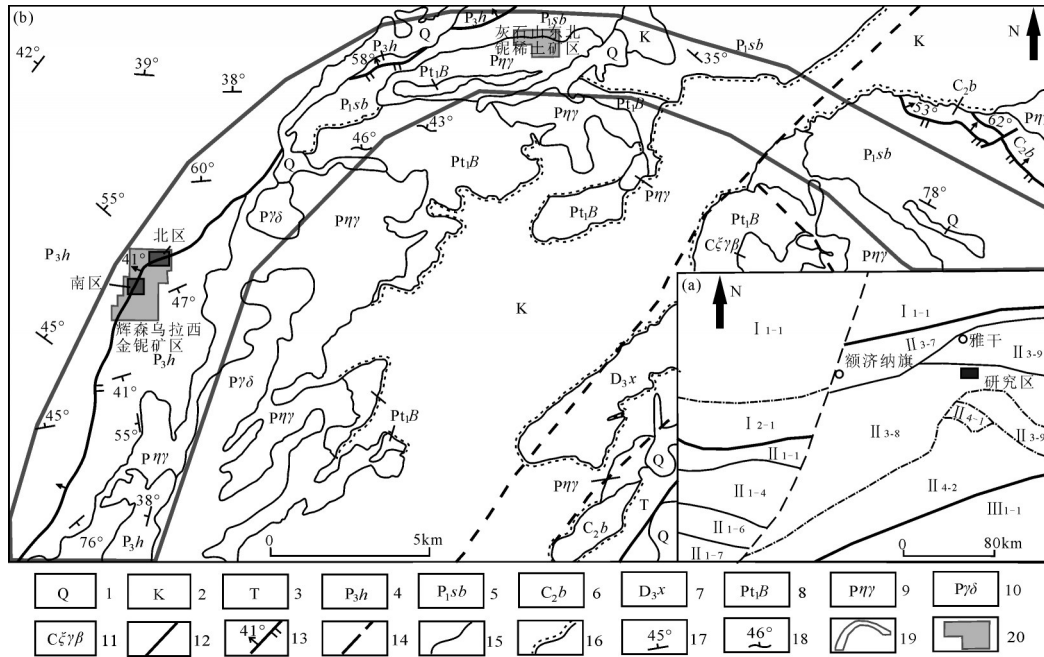


图1 呼伦西白地区大地构造位置(a)及地质略图(b)(据左国朝等,2003修改)

I—哈萨克斯坦板块; I₁—雅满苏—红石山构造区; I₂—星星峡—旱山构造区; I₁₋₁—雅满苏—红石山晚古生代裂陷海盆褶皱带; I₂₋₁—尾垭—星星峡—旱山地块; II—塔里木板块; II₁—公婆泉—洪果尔构造区; II₁₋₁—东七—山—洪果尔早古生代洋盆褶皱带; II₁₋₄—鹰嘴红山地块; II₁₋₆—沙红山二叠纪裂谷褶皱带; II₁₋₇—红柳河—红柳大泉志留纪弧后盆地褶皱带; II₃—敦煌—玉门构造区; II₃₋₇—雅干北中泥盆世裂陷海槽褶皱带; II₃₋₈—雅干南晚二叠世拉分盆地褶皱带; II₃₋₉—恩格尔乌苏北早二叠世裂谷褶皱带; II₄—珠斯楞构造区; II₄₋₁—珠斯楞寒武—奥陶纪陆棚海褶皱带; II₄₋₂—金塔—恩格尔乌苏晚古生代洋区褶皱带; III—华北板块; III₁—阿拉善构造区; III₁₋₁—宗乃山—沙拉扎山晚古生代岛弧褶皱带; 1—第四系; 2—白垩系; 3—三叠系; 4—二叠系上统哈尔苏海组; 5—二叠系下统双堡塘组; 6—石炭系上统白山组; 7—泥盆系上统西屏山组; 8—古元古界北山岩群; 9—二叠纪二长花岗岩; 10—二叠纪花岗闪长岩; 11—石炭纪黑云钾长花岗岩; 12—实测性质不明断层; 13—实测逆断层; 14—推测断层; 15—实测地质界线; 16—实测角度不整合界线; 17—地层产状; 18—面理产状; 19—铌稀土多金属矿化蚀变带; 20—典型矿床位置

Fig.1 Geotectonic location (a) and geological sketch map (b) of Hulunxibai area (modified from Zuo Guochao et al., 2003)

I—Kazakhstan plate; I₁—Yamansu—Hongshishan tectonic region; I₂—Xingxingxia—Hanshan tectonic area; I₁₋₁—Late Paleozoic chasmic basin fold belt in Yamansu—Hongshishan; I₂₋₁—Weiya—Xingxingxia—Hanshan block; II—Tarim plate; II₁—Gongpoquan—Honggouer tectonic area; II₁₋₁—Early Paleozoic ocean basin fold belt in Dongqiyishan—Honggouer; II₁₋₄—Yingzuihongshan block; II₁₋₆—Permian rift fold belt in Shahongshan; II₁₋₇—Silurian back-arc basin fold belt in Hongliuhe—Hongliudaquan; II₃—Dunhuang—Yumen tectonic area; II₃₋₇—Middle Devonian chasmic trough fold belt in northern Yagan; II₃₋₈—Late Permian pull-apart basin fold belt in south Yagan; II₃₋₉—Early Permian rift fold belt in northern Engeerwusu; II₄—Zhusileng tectonic area; II₄₋₁—Cambrian—Ordovician continental shelf sea fold belt in Zhusileng; II₄₋₂—Late Paleozoic ocean area fold belt of Jinta—Engeerwusu; III—North China plate; III₁—Alax tectonic area; III₁₋₁—Late Paleozoic island arc old belt in Zongnaishan—Shalazhashan; 1—Quaternary; 2—Cretaceous; 3—Triassic; 4—Upper Permian Haersuhai Formation; 5—Lower Permian Shuangbaotang Formation; 6—Upper Carboniferous Baishan Formation; 7—Upper Devonian Xipingshan Formation; 8—Paleoproterozoic Beishan Group; 9—Permian adamellite; 10—Permian granodiorite; 11—Carboniferous biotite K—feldspar granite; 12—Unknown fault; 13—Reverse fault; 14—Inferred fault; 15—Geological boundary; 16—Unconformity; 17—Attitude of strata; 18—Attitude of foliation; 19—Niobium and rare earth polymetallic mineralization zone; 20—Typical deposit location

稀土多金属矿床(点)整体呈弧形分布于二叠系与海西晚期岩体的外接触带。出露地层主要有古元古界北山群(P_{tB})、古生界上泥盆统西屏山组(D_{3x})、上石炭统白山组(C_2b)、下二叠统双堡塘组(P_{1sb})、上二叠统哈尔苏海组(P_3h);中生界三叠系(T)、白垩系(K)及新生界第四系(Q)。区域侵入岩较为发育,主要为二叠纪二长花岗岩($P_{\eta\gamma}$)、花岗闪长岩($P_{\gamma\delta}$)、石炭纪黑云钾长花岗岩($C_{\zeta\gamma\beta}$)。产出形态以岩基和岩株为主,主要呈NW和NE展布。区内断裂构造发育,按照断裂的展布方向可划分为NW向、NE两组,其中NW断裂控制了区内地层和岩浆岩的分布格局,NE断裂则对NW断裂、地层、岩体进行了改造,这些断裂均具多期活动性,它们共同构成了区内的构造格架(图1b)。

区内与稀有稀土多金属成矿元素有关的地层主要与下二叠统双堡塘组(P_{1sb})、上二叠统哈尔苏海(P_3h)组滨浅海相碎屑岩内。如灰石山东北区内发现的铌-稀土矿均产自二叠系下统双堡塘组内,含矿岩性段为变质粉砂岩、硅泥质板岩、钙质砂泥质板岩等等;辉森乌拉西区内发现铌矿体主要产于二叠系上统哈尔苏海组变质岩屑石英砂岩中。区域一带内生矿产的成矿作用与岩浆侵入和期后热液活动关系密切,岩浆不仅是成矿元素的源泉,而且由于岩浆热液的活动,给成矿物质的搬运、富集成矿创造了有利条件(翟裕生等,1993,2011),区内目前已发现的稀有稀土矿床、矿(化)体、矿点均分布于二叠纪二长花岗岩($P_{\eta\gamma}$)与二叠系外接触带上。区域尺度上,北部有哈珠一雅干深断裂,南部有乌兰苏海深断裂,控制了区域内EW、NE、NW向断裂。复杂的断裂构造为区域岩浆侵位提供空间,从而为成矿元素的活化、迁移、富集提供热能。矿区尺度上,NE、NE、近EW向断裂既提供了含矿热液运移的通道,又提供了矿质沉淀的空间,如灰石山东北铌稀土矿(化)体主要分布于近EW向断层或其附近,辉森乌拉西发现的铌(金)矿体主要沉淀于NE断层或其附近。露头尺度上,异常发育的小断裂、节理裂隙亦为矿质沉淀提供了空间,目前发现的大多矿体都赋存于断层附近的次级裂隙内或挤压应力作用下形成的脆性变形虚脱空间内。

3 典型矿床特征

3.1 灰石山东北铌稀土矿

3.1.1 成矿地质特征

该矿区出露地层较为简单,主要分布下二叠统双堡塘组(P_{1sb})、上二叠统哈尔苏海组(P_3h)、上白垩统乌兰苏海组(K_2w),其中下二叠统双堡塘组(P_{1sb})是区内主要含铌稀土地层,岩性以含钙质变质粉砂岩夹薄层灰岩及砂质板岩为主,局部零星出露变质石英砂岩、角闪二长片麻岩和斑点板岩;区内构造极为发育,大致可分为近EW、NW、NE向三组,从矿体分布情况来看近EW、NE向断层是主要储矿构造,从图2中可看出,矿体主要呈EW、NE向分布,且均就位于断层内或其附近。NW向断层为成矿后构造,不同程度的破坏了矿体的连续性或使其错位。矿区岩浆岩分布较为简单,主要呈岩基、岩株状产出,岩体主要为二叠纪钾长花岗岩($P_{\zeta\gamma}$)、二长花岗岩($P_{\eta\gamma}$)、花岗闪长岩($P_{\gamma\delta}$),展布方向近EW—NE向,与区内主体构造方向一致。

区内正磁异常与矿体空间展布套合较好, ΔT 异常形态总体呈串珠状沿近EW—NE向分布,主要以条带状和串珠状异常为主,北西部的异常水平为最高,北东部异常次之,东南部的异常规模较小,以上异常均由几个椭圆状次级异常组成,性质相似。结合地质情况,该区正磁异常空间分布与铌稀土矿化对应,且I Nb矿体钻孔验证发现了磁铁矿,显示该区正异常均可能与含磁性矿物的铌稀土矿有关。

3.1.2 矿化蚀变带特征

区内目前发现了3条铌稀土多金属矿化蚀变带,主要以Nb为主,编号分别为I Nb、II Nb、III Nb(图2,表1),除I Nb为铌、稀土、磁铁复合矿化蚀变带外,其余两条均为独立铌矿化蚀变带。蚀变带地表主要以褐铁矿化为主(图3a),局部石英细网脉发育(图3b),岩性以钙质粉砂岩夹灰岩透镜体为主(图3c)。

I Nb矿化蚀变带长度约800 m,最宽约50 m,平均厚度22.38 m,蚀变带走向总体呈NEE向,其中中西部为EW向,东部为NE向;倾向总体NNW向,产状 $340^{\circ}\sim 350^{\circ}\angle 55^{\circ}\sim 70^{\circ}$ 。矿体呈似层状产出,地表品位 Nb_2O_5 0.02%~0.1%,平均0.09%;REO(稀土

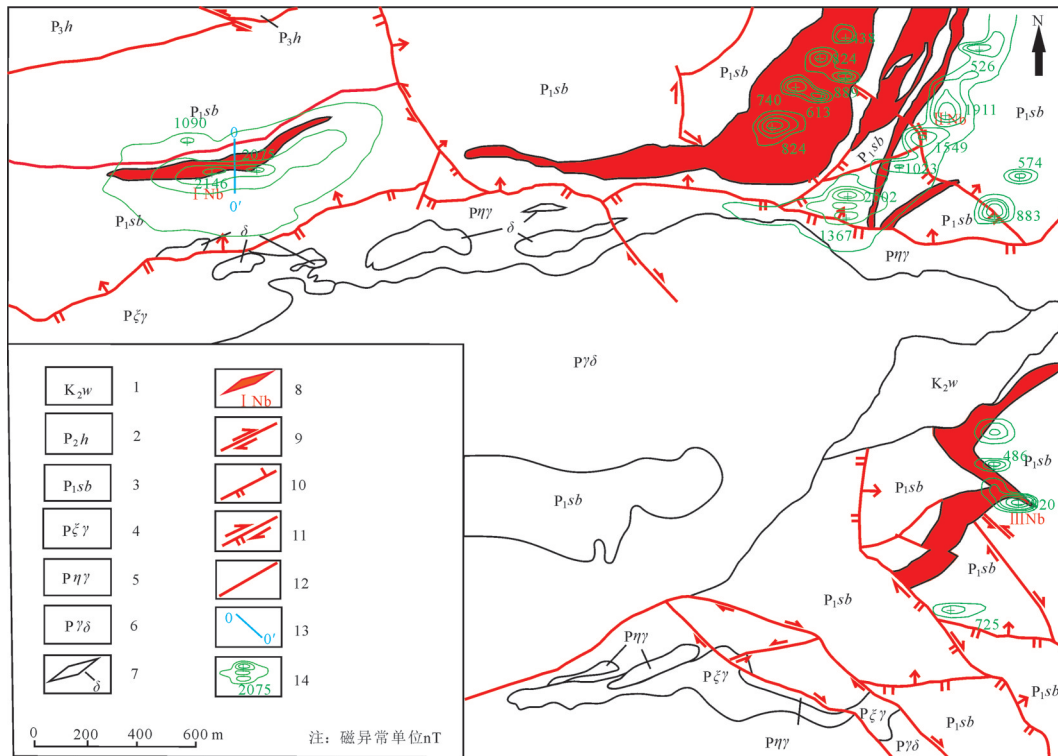


图2 灰石山东北矿区综合地质简图(据刘强等,2018修改)

1—白垩系上统乌兰苏海组;2—二叠系上统哈尔苏海组;3—二叠系下统双堡塘组;4—二叠纪钾长花岗岩;5—二叠纪二长花岗岩;6—二叠纪花岗岩闪长岩;7—闪长岩脉;8—铌矿体及编号;9—平移断层;10—逆断层;11—平移逆断层;12—性质不明断层;13—I Nb断面位置;14—1:1万高精度磁异常范围及极值(极值单位:nT)

Fig.2 Comprehensive geological sketch map of the northeast Huishi Mountain mining area(after Liu Qiang et al.,2018)

1—Upper Cretaceous Wulansuhai Formation; 2—Upper Permian Haersuhai Formation; 3—Lower Permian Shuangbaotang Formation; 4—Permian syenogranite; 5—Permian adamellite; 6—Permian adamellite; 7—Diorite dyke; 8—Niobium orebody and its serial number; 9—Parallel displacement fault; 10—Reverse fault; 11—Slip—reverse fault; 12—Unknown fault; 13— I Nb orebody section position; 14— 1:10000 high precision magnetic anomalies and maximum value (unit: nT)

氧化物)品位为0.52%~1.37%,平均0.93%。矿体深部层位较为稳定,一个钻孔见矿厚度41.00~154.27 m,平均85.96 m,品位Nb₂O₅ 0.04%~0.56%,平均0.11%; REO 0.13%~2.63%,平均0.93%,变化系数51.64%,属均匀型; TFe 7.55%~22.97%,平均14.09%; mFe 7.45%~12.74%,平均10.10%。铌—稀土矿体厚度3.50~94.27 m,平均厚度46.22 m。矿体埋深50 m(顶板)~200 m(底板),厚度变化系数46.41%,属稳定型,形态为似层状,复杂程度属简单型(图4)。

II Nb矿化蚀变带控制总长约1.50 km,平均厚度92.23 m,蚀变带走向总体NE向,向西扭转呈近EW,倾向NW,倾角40°~70°。该矿体是区内目前发现的规模最大的一条铌矿体,呈似层状、厚板状、具膨缩和分支复合现象,整体较为连续;矿体厚度变化系数76.41%,矿石品位Nb₂O₅ 0.11%~0.01%,平均0.03%,品位变化系数42.23%。

III Nb矿化蚀变带长约400 m,最宽处约16 m,平均厚度15 m;矿体走向总体呈北东向,倾向总体北西向,产状330°~335°∠68°~72°;矿体呈似层状产

表1 灰石山东北铌稀土矿化蚀变带特征简表

Table 1 Characteristic table of Nb-REE mineralized alteration zone in northeast Huishi Mountain

矿化蚀变带编号	蚀变带规模			产状			含矿岩石品位/%			
	长度/m	宽度/m	厚度/m	走向	倾向	倾角	Nb ₂ O ₅	REO	TFe	mFe
I Nb	800	5~50	3.5~94.27	75°~90°	340°~350°	55°~70°	0.02~0.56	0.13~2.63	7.55~22.97	7.45~12.74
II Nb	1500	15~500	92.23	60°~90°	330°~340°	40°~70°	0.01~0.11	\	\	\
III Nb	400	3~16	15	50°~70°	330°~335°	68°~72°	0.05~0.08	\	\	\

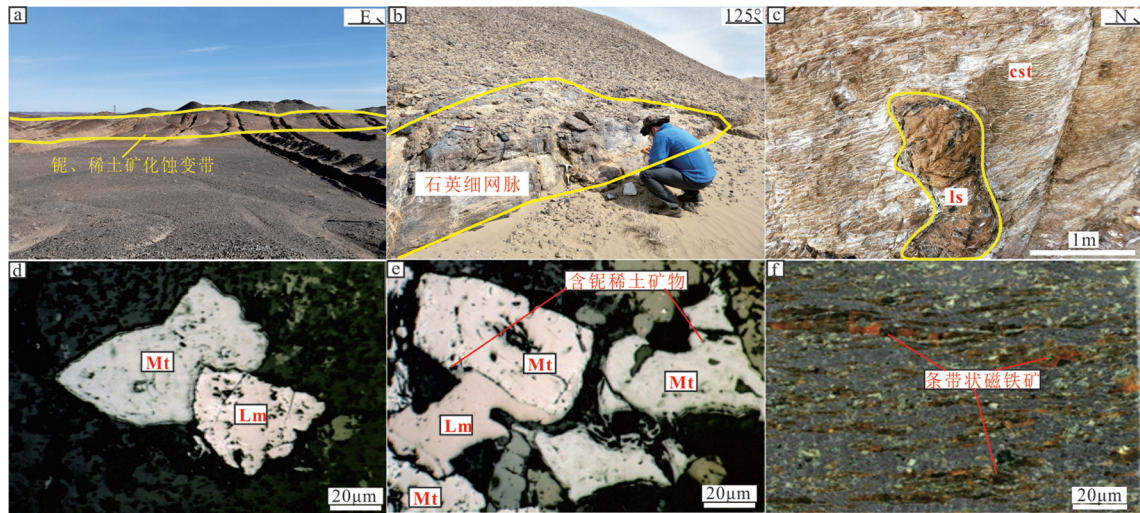


图3 灰石山东北铌稀土矿含矿岩石野外(a, b, c)及镜下(d, e, f)照片
 cst—钙质粉砂岩;ls—灰岩;Mt—磁铁矿;Lm—褐铁矿;a—ⅡNb矿化蚀变带;b—石英细网脉发育地段;c—二叠系下统双堡塘组钙质粉砂岩夹灰岩;d—f—镜下照片(光片)

Fig. 3 Field photos (a, b, c) and photomicrographs (b, c, d) of northeast Huishi Mountain Nb-REE ore-bearing rocks
 cst—Calcareous siltstone; ls—Limestone; Mt—Magnetite; Lm—Limonite; a—I Nb mineralization alteration zone; b—Quartz vein development area; c—Lower Permian Shuangbaotang Formation calcareous siltstone intercalated with limestone; d—f—Microscopic photos from flat light sheets

出;矿石品位 Nb_2O_5 0.05%~0.08%,平均0.06%,变化系数22.28%,属均匀型。矿体厚度与品位变化相对都不大,矿体顶板岩性为变质粉砂岩,底板岩性为砂质板岩。

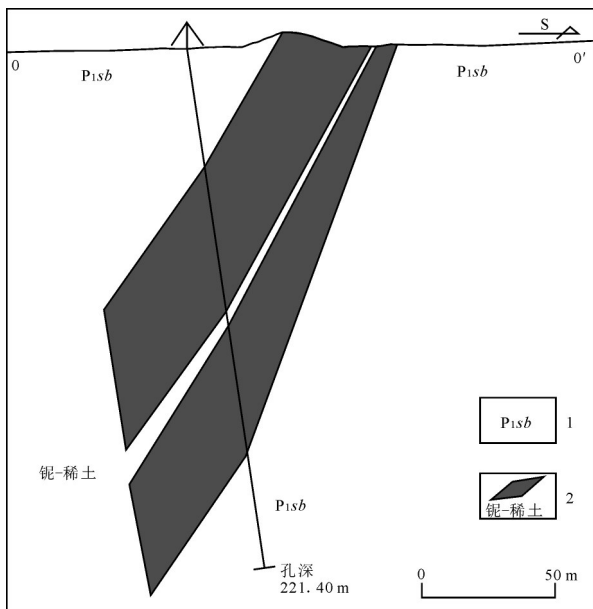


图4 灰石山东北 I Nb 矿体断面简图
 1—二叠系下统双堡塘组;2—铌-稀土矿体
 Fig.4 I Nb orebody section of the northeast Huishi Mountain mining area
 1—Lower Permian Shuangbaotang Formation; 2—Nb-REE orebody

3.1.3 含矿岩石矿物特征

结构:主要以半自形—他形晶粒状、变余半自形—他形粒状和不规则粒状结构为主,次为变余浑圆状结构、显微鳞(叶)片粒状变晶结构(图3d,e)。构造:以块状构造、板状构造、变余层状构造为主,次为星点状浸染构造、弱千枚状构造。含矿岩石中主要稀土矿物有氟碳铈矿、氟碳钙铈矿和磷灰石等,主要含铌矿物为铌铁矿、铌铁金红石、褐钇铌矿等;硫化物可见褐铁矿、条带状磁铁矿(图3f)、黄铁矿、磁黄铁矿、方铅矿、黄铜矿等;主要围岩矿物有斜长石、钾长石、石英、黑云母、绢云母、绿泥石、碳酸盐类矿物等。

3.1.4 围岩蚀变特征

围岩蚀变主要有硅化、绿泥石化、碳酸盐化、褐铁矿化等,次为绢云母化、磁铁矿化、磁黄铁矿化、黄铁矿化、方铅矿化等。其中硅化主要出现于二长花岗岩、灰岩、泥灰岩、砂岩内;绿泥石化出现于变质粉砂岩、砂岩内;碳酸盐化出现于变质粉砂岩;褐铁矿化出现于二长花岗岩、变质粉砂岩、泥灰岩、砂岩。

3.1.5 成因类型分析

矿(化)体主要赋存于二叠纪二长花岗岩外接触带的二叠系下统双堡塘组内的一套变碎屑岩内,并受NE—近EW向断裂控制,后期岩浆热液活动对

铌稀土矿化有进一步的叠加作用,刘强等(2018)认为该矿床属沉积变质-热液叠加型成因。

3.2 辉森乌拉西铌金矿

3.2.1 成矿地质特征

该矿区因矿体地表分布距离较远分为北、南两个子区(图1b),矿区地层主要为上二叠统哈尔苏海组(P_3h),呈NE—SN向展布,倾角 $60\sim 80^\circ$,呈现粗粒石英砂岩与细粒岩屑砂岩的沉积旋回特征(图5a),亦为一套滨浅海碎屑沉积岩,主要岩性为变质岩屑砂岩(P_3hss)夹薄层灰岩(P_3hls)、硅质灰岩(P_3hsils)。岩层中褶皱、片理发育。区内岩浆岩活动主要集中于海西晚期,岩石类型有深成侵入岩和脉岩等。深成侵入岩岩性主要为二叠纪花岗闪长岩($P\gamma\delta$)、二长花岗岩($P\eta\gamma$)。区内脉岩主要有石英脉(q)、闪长岩脉(δ)等,区内断裂非常发育,主要表现为片理化带、韧性变形带,主体方向为NNE。

3.2.2 矿化蚀变带特征

区内地表目前发现了3条矿化蚀变带(表2),编号分别为KH-1、KH-2(图6)和KH-3(图7),其中KH-3矿化蚀变带地表以金为主,经钻探验证在深部发现隐伏铌矿体。

KH-1矿化蚀变带地表长约1.6 km,宽约30 m,走向 $40\sim 50^\circ$ (图5b),倾向南东,倾角 70° 左右,金含

量 $0.11\sim 0.98$ g/t,铌含量 $0.01\%\sim 0.05\%$,矿化连续性较好。含金、铌矿化带主要赋存于哈尔苏海组变质岩屑石英砂岩中,岩石破碎强烈(图5c),围岩蚀变主要见绢云母化、绿帘石化、碳酸盐化现象,岩性为硅化、褐铁矿化变质石英砂岩,岩石主要具硅化、蜂窝状褐铁矿化、局部可见方铅矿化,呈团块状,石英细脉、微细脉穿插发育。岩石普遍片理化,挤压破碎现象,经化学样品分析验证,具千枚状片理化岩石均为围岩或夹石,含铌矿化带岩性为硅化变质岩屑石英砂岩,呈似层状带状连续产出,局部具膨大富集特征;含金矿化蚀变带总体位于含铌矿化带中,沿走向呈断续产出。

KH-2矿化蚀变带为地表拣块分析控制,走向平行于KH-1,宽约50 m,长约1.5 km,倾向 125° ,倾角 65° 左右。矿化带主要赋存于哈尔苏海组上统变质岩屑石英砂岩中,围岩蚀变主要见绢云母化、绿帘石化、碳酸盐化现象,岩性为硅化变质岩屑石英砂岩,主要具硅化、褐铁矿化,石英细脉、微细脉穿插发育,5件拣块样品中 Nb_2O_5 品位 $0.06\%\sim 0.1\%$,Au品位 $0.03\sim 0.13$ g/t,达到了原生铌矿床边界品位。

KH-3矿化蚀变带地表断续长约2.6 km,宽 $1.00\sim 14.30$ m,倾向西,倾角 $70^\circ\sim 80^\circ$,该矿化带受

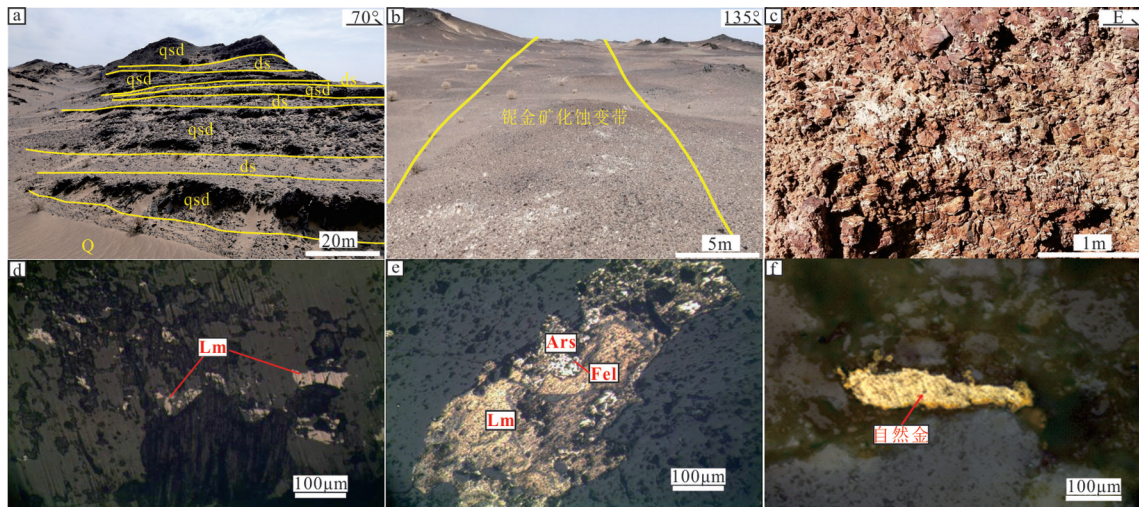


图5 辉森乌拉西铌金矿含矿岩石野外(a, b, c)及镜下(d, e, f)照片

qsd—石英砂岩;ds—岩屑砂岩;Lm—褐铁矿;Ars—毒砂;Fel—铌铁矿;a—二叠系上统哈尔苏海组岩屑砂岩与石英砂岩互层;b—KH-1矿化蚀变带;c—含铌金矿岩石地表破碎特征;d—f—镜下照片(光片)

Fig.5 Field photos (a, b, c) and photomicrographs (b, c, d) of west Huisenwula Nb-Au ore-bearing rocks
qsd—Quartz sandstone; ds—Lithic sandstone; Lm—Limonite; Ars—Arsenopyrite; Fel—Columbite; a—Interbedding of lithic sandstone and quartz sandstone of Upper Permian Wulansuhai Formation; b—KH-1 Nb-Au mineralization alteration zone; c—Nb-Au ore-bearing rocks characteristics of surface fragmentation; d—f—Microscopic photos from flat light sheets

表2 辉森乌拉西铌金矿化蚀变带特征简表

Table 2 Characteristic table of Nb-Au mineralized alteration zone in west Huisenwula

矿化蚀变带编号	蚀变带规模		产状			含矿岩石品位	
	长度/m	宽度/m	走向	倾向	倾角	Nb ₂ O ₅ /%	Au/(g/t)
KH-1	1600	39~142	40°~50°	SE	70°	0.01~0.05	0.11~0.98
KH-2	1500	1~50	40°~50°	125°	65°	0.06~0.1	0.03~0.13
KH-3	2600	1~14.3	90°~120°	W	70°~80°	0.04~0.1	0.1~0.99

韧-脆性变形带控制,为构造挤压形成,主体岩性为褐铁矿化、硅化变质石英砂岩,岩石具粒状、蜂窝状褐铁矿化、硅化,局部可见方铅矿物集合体,并在裂隙面具黄绿色硫化物氧化膜,石英细脉、微细脉穿插发育,岩石片理化,挤压破碎强烈,整体呈韧-脆性变形特征。含金矿化蚀变带总体呈断续产出,局部具膨胀收缩、分枝复合现象,Au品位0.10~0.99 g/t。在7勘探线上经钻探工程控制一条深部隐伏铌矿体,斜深94.43~114.80 m,假厚19.37 m,真厚度14.17 m,Nb₂O₅品位0.05%~0.10%,平均品位0.06%,矿体均呈层状赋存于变质岩屑石英砂岩中,矿体与围岩界线较不清晰,矿体倾向西,在

斜深25.65~34.00 m,假厚8.35 m,Nb₂O₅品位0.04%~0.05%,品位较高(图8)。

3.2.3 含矿岩石矿物特征

结构主要为片状他形变余结构,交代残留结构,包含结构,碎裂结构。构造主要有片状构造,稀疏浸染状构造、细脉浸染状构造、网脉状构造。主要含铌矿物为铌铁矿、毒砂等(图5d,e),金主要为裂隙金(图5f)和包裹金,主要硫化物为黄铁矿、方铅矿、磁黄铁矿等,主要脉石矿物为石英、黑云母、斜长石、绿泥石及方解石等。

3.2.4 围岩蚀变特征

围岩蚀变主要有有硅化、碳酸盐化,局部见绿

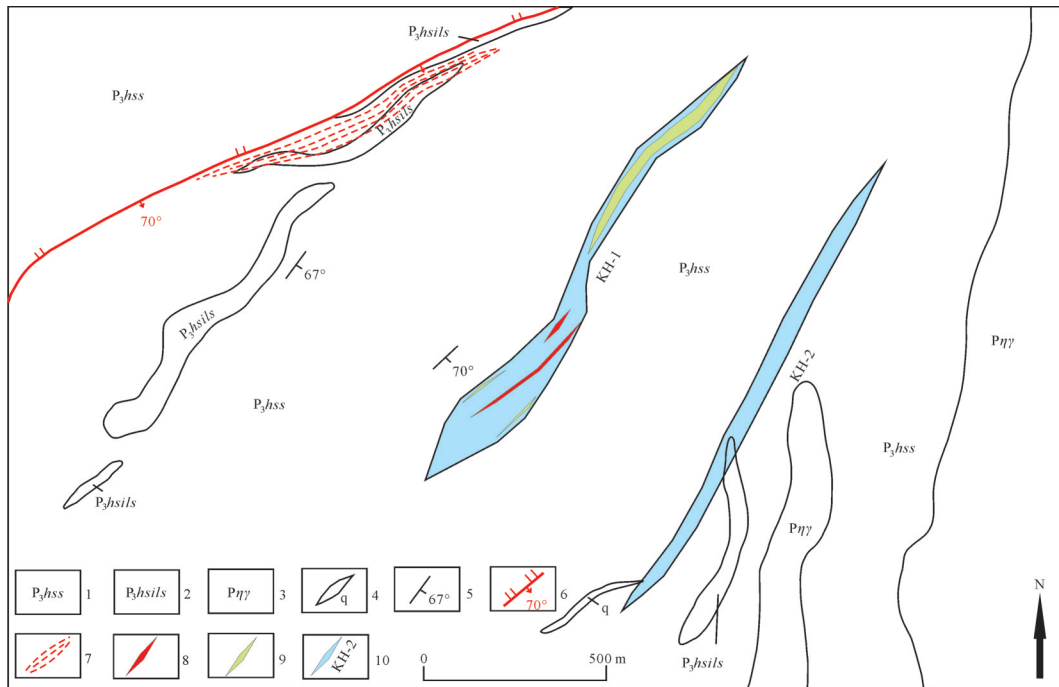


图6 辉森乌拉西铌金矿地质简图(北矿区)

1—上二叠统哈尔苏海组变质岩屑砂岩;2—上二叠统哈尔苏海组硅质灰岩;3—二叠纪二长花岗岩;4—石英脉;5—地层产状;6—逆断层及倾角;7—韧脆性变形带;8—铌矿体;9—铌(金)矿体;10—矿化蚀变带

Fig. 6 Geological sketch map of the west Huisenwula niobium gold deposit (north mining area)

1—Metamorphic lithic sandstone of Upper Permian Wulansuhai Formation; 2—Siliceous limestone of Upper Permian Wulansuhai Formation; 3—Permian adamellite; 4—Quartz vein; 5—Attitude of strata; 6—Reverse fault and dip; 7—Ductile-brittle deformation zone; 8—Niobium orebody; 9—Niobium (gold) orebody; 10—Mineralization alteration zone

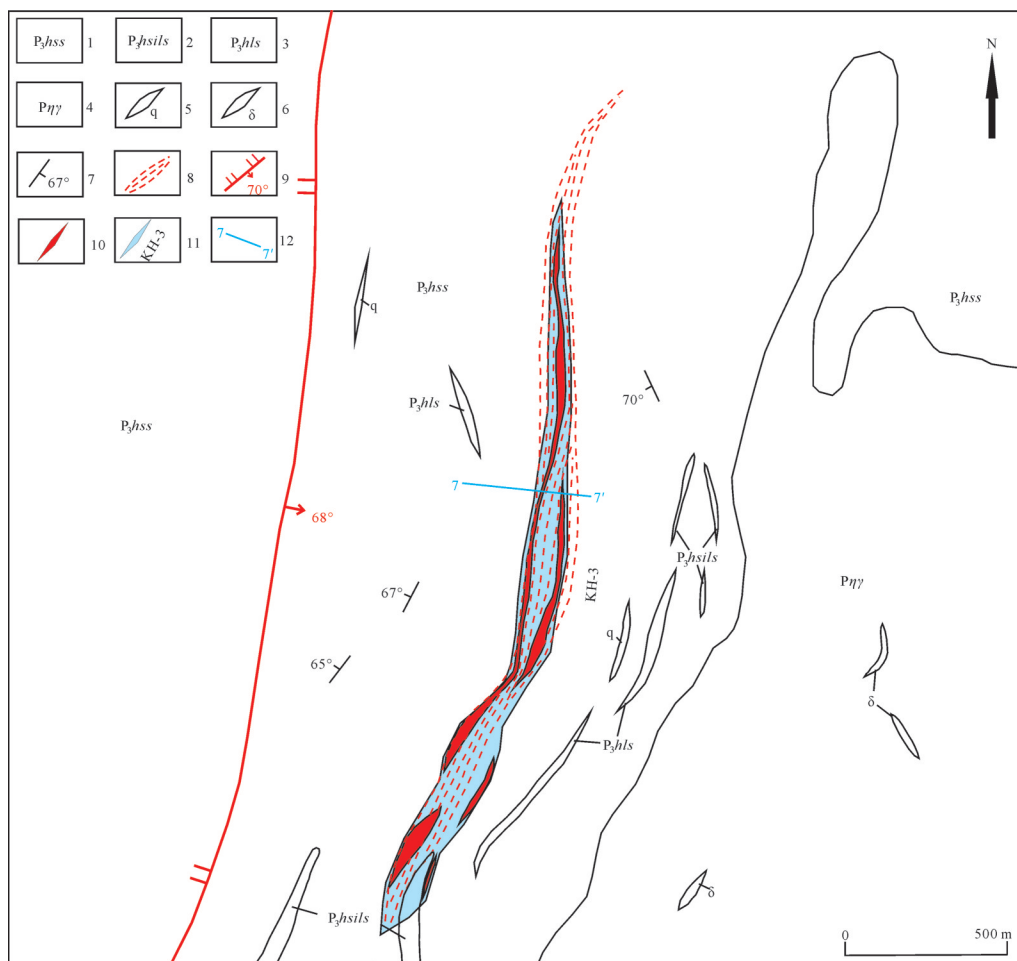


图7 辉森乌拉西铌金矿地质简图(南矿区)

1—上二叠统哈尔苏海组变质岩屑砂岩;2—上二叠统哈尔苏海组硅质灰岩;3—上二叠统哈尔苏海组薄层灰岩;4—二叠纪二长花岗岩;5—石英脉;6—闪长岩脉;7—地层产状;8—韧—脆性变形带;9—逆断层及倾向;10—铌金矿(化)体;11—矿化蚀变带;12—7勘探线剖面位置

Fig. 7 Geological sketch map of the west Huisenwula niobium gold deposit(south mining area)

1— Metamorphic lithic sandstone of Upper Permian Wulansuhai Formation; 2— Siliceous limestone of Upper Permian Wulansuhai Formation; 3— Thin limestone of Upper Permian Wulansuhai Formation; 4— Permian adamellite; 5— Quartz vein; 6— Diorite dyke; 7— Attitude of strata; 8— Ductile-brittle deformation zone; 9— Reverse fault and dip; 10— Niobium and gold mineralized body; 11— Mineralization alteration zone; 12— Geological section along No. 7 exploration line

泥石化、钾化、高岭土化等,其中铌金矿化带与硅化、黄铁矿化、褐铁矿化等蚀变关系密切。

3.2.5 成因类型分析

区内金铌矿(化)体严格受控于二叠系上统哈尔苏海组褐铁矿化变质石英砂岩蚀变带内,在褐铁矿化变质石英砂岩蚀变带中形成次级挤压破碎带,为矿床的形成提供了有利的成矿空间,受二叠纪二长花岗岩及后期热液活动影响,含矿热液沿地层裂隙充填形成了多条铌金矿体,初步认为中低温热液型。

4 成矿潜力浅析

呼伦西白地区铌稀土矿床(点)是内蒙古北山

地区近年来地质勘查的新发现,从前文介绍来看,二叠系滨浅海相碎屑岩—碳酸盐岩地层是铌稀土主要含矿层位,矿体形态基本受控于主断裂裂隙系统内,后期岩浆热液活动对铌稀土具叠加作用,显示多期次成矿特征,尤其是灰石山东北 I Nb 矿化蚀变带内铌—稀土含矿岩石品位富且矿体形态稳定,深部还发现了磁铁矿,而辉森乌拉西矿区深部钻探验证发现了隐伏铌矿体,已知矿床(点)均具特征的放射性—磁异常。笔者将该矿带内铌稀土矿与白云鄂博超大型矿床进行了类比分析,发现两者之间具众多相似性(表3),认为该地区铌稀土多金属矿成矿潜力巨大。

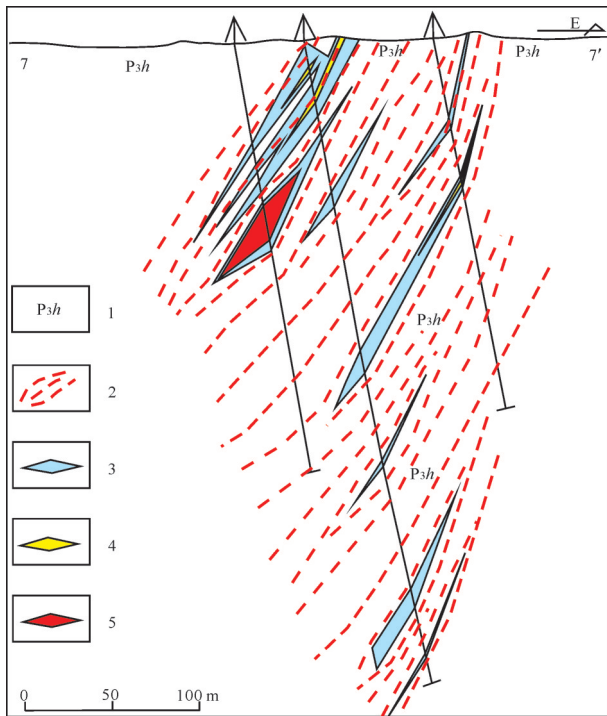


图8 辉森乌拉西铌金矿7号勘探线剖面图
1—二叠系上统哈尔苏海组;2—韧—脆性变形带;3—铌金矿(化)体;4—金矿体;5—铌矿体

Fig. 8 Geological section along No.7 exploration line of the west Huisenwula niobium gold deposit
1- Upper Permian Wulansuhai Formation; 2- Ductile- brittle deformation zone; 3- Niobium and gold mineralized body; 4- Gold orebody; 5- Niobium orebody

5 找矿标志

5.1 二叠系滨浅海相碳酸盐岩-变碎屑岩地层

区域上双堡塘组由下到上,砾岩逐渐减少消失,灰岩层数增多,层厚变厚,反映当时的海进环境,灰岩中含有大量生物化石,发育斜层理,说明其沉积环境为滨浅海相;哈尔苏海组下部以细碎屑岩为主,中部粒度变粗,上部粒度变细,呈现一粗一细的沉积旋回,中部出现含砾砂岩,向上有变粗趋势,是海退层序,也为滨浅海沉积特征。区内与稀有稀土多金属成矿元素有关的地层主要与下二叠统双堡塘组、上二叠统哈尔苏海组碎屑岩有关,尤其是褐铁矿化、硅化蚀变地段为铌稀土多金属矿主要含矿层位。

5.2 构造挤压形成的虚脱空间或韧—脆性转换部位

挤压应力条件下形成的虚脱空间往往是成矿元素有利沉淀部位,构造变形引发的动力分异作用形成的变质动力热液是成矿流体的主要来源,而韧—脆性剪切构造既是唯一的赋矿构造,也对矿化体(带)的形态、产状、规模及分布起着决定性作用(翟裕生等,1993;陈柏林等;2002)。产于韧性剪切带中的成矿元素通常有逐渐富集的过程,在糜棱岩化过程中,剪切作用可以使原岩中含量很低的元素发生迁移、而在糜棱岩阶段之后叠加于其上的脆性变形阶段发生时,由于劈理、片理和裂隙的发育导致

表3 呼伦贝尔西白地区铌稀土矿与白云鄂博超大型矿床对比

Table 3 Comparison of the Nb-REE deposits (ore spots) in Hulunxibai area and Nb-REE-Fe superlarge deposits in Bayan Obo

矿床	大地构造位置	成矿地质背景	航磁航放特征	矿床地质特征	矿床成因
白云鄂博超大型矿床	华北陆块北缘狼山—白云鄂博裂陷槽东段白云鄂博褶断束多组断裂复合部位(李江海等,2006)	矿体赋存于中元古界白云鄂博群哈拉霍疙特岩组H8白云岩,矿床经历多期成矿作用,后期岩浆热液活动对铌稀土具叠加富集作用(翟裕生,2003),矿体受向斜构造的控制	具明显的航磁异常和航放总计量、钽、铀异常,航放钾异常高背景或梯度带与主矿体对应(王继春等,2016)	矿带长14 km,宽1~2 km,自西向东有西矿、主矿及东矿、东介勒格勒和都拉哈拉5个矿体,铁的含量以块状铁矿石最高,TFe含量平均58.78%,稀土元素主要为钽族元素,占稀土总量的97%,Nb ₂ O ₅ 含量平均0.16%。放射性元素ThO ₂ 含量0.029%~0.051%,稀散元素有Ga、In、Sc、Rb、Cs及Zr和Hf。稀土矿物有独居石、氟碳铈矿、氟碳铈钨矿等18种。铌矿物有铌铁金红石、铌铁矿、易解石、烧绿石等19种(张培善等,2001)	多成因特征,主要有沉积-热液叠加、火成碳酸岩型、后生热液交代型、混合型(朱祥坤等,2012)
呼伦贝尔西白地区铌稀土矿	塔里木板块之敦煌—玉门构造区之雅干南晚二叠世拉分盆地褶皱带(左国朝等,2003)	矿床(点)呈弧形分布二叠纪二长花岗岩外接触带的二叠系滨浅海相碎屑岩-碳酸盐岩内,矿体形态受主断裂裂隙控制,后期岩浆热液对铌稀土有叠加富集作用	已知矿床(点)及矿化蚀变位于航放钽、铀正负异常梯度带和带状航磁异常位置	灰石山东北铌稀土矿区发现了三条铌稀土多金属矿化蚀变带,主要以Nb为主,编号分别为I Nb、II Nb、IIINb, Nb ₂ O ₅ 品位在0.02%~0.56%、平均0.11%,REO品位在0.13%~2.63%、平均0.93%;辉森乌拉西铌金矿区发现了三条铌金矿化蚀变带,编号分别为KH-1、KH-2和KH-3, Nb ₂ O ₅ 品位在0.01%~0.1%。稀土矿物有氟碳铈矿、氟碳钙铈矿和磷灰石等,主要含铌矿物为铌铁矿、铌铁金红石、褐钨铌矿等	灰石山东北铌稀土矿为沉积变质-热液叠加型(刘强等,2018),辉森乌拉西具明显的后期热液特征,为中低温热液型

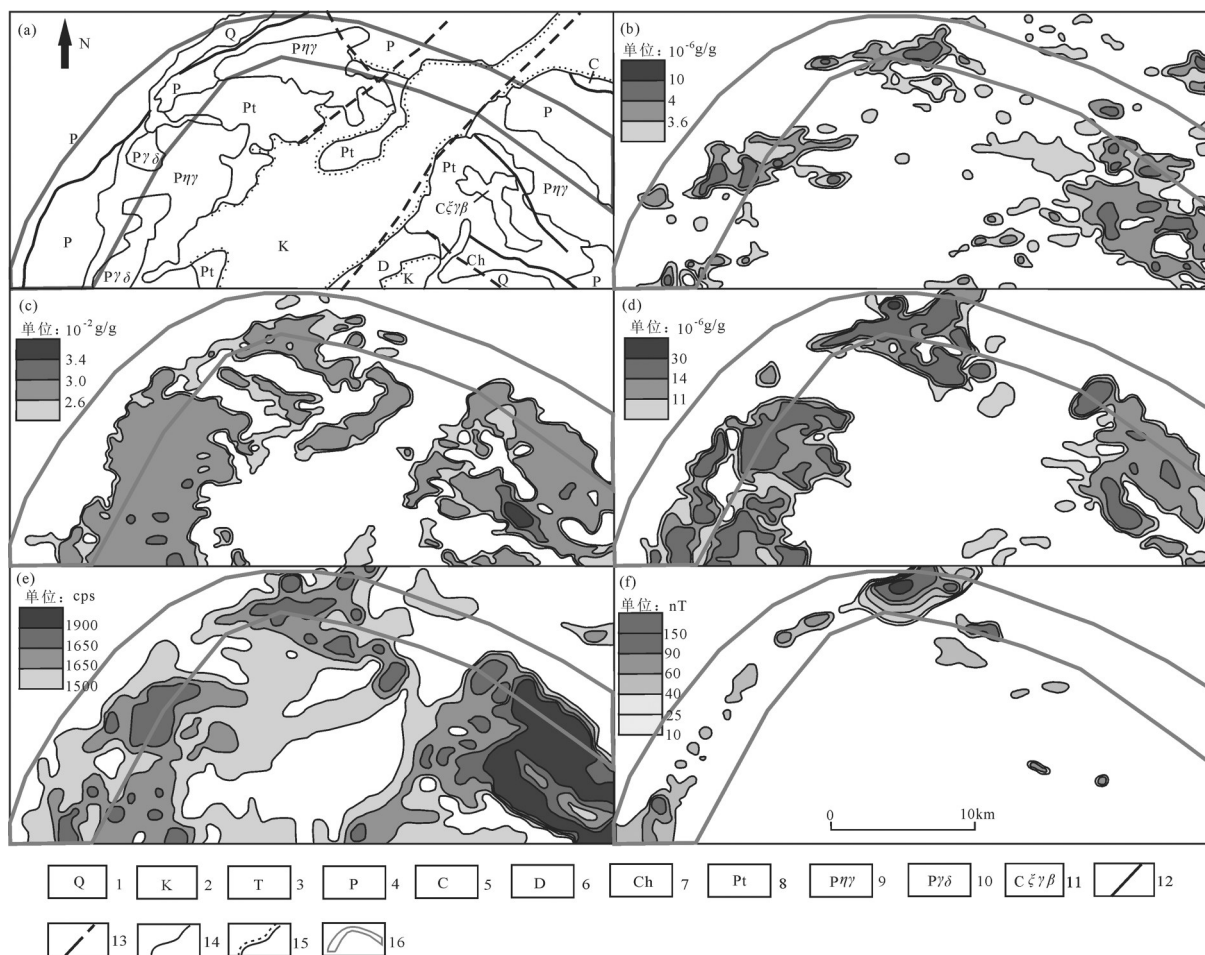


图9 呼伦西白地区1:5万航磁、航放特征简图

a—呼伦西白地区地质简图；b—航放铀等值线异常图；c—航放钾等值线异常图；d—航放钍等值线异常图；e—航放计数率等值线异常图；f—航磁 ΔT 等值线异常图；1—第四系；2—白垩系；3—三叠系；4—二叠系；5—石炭系；6—泥盆系；7—长城系；8—元古界；9—二叠纪二长花岗岩；10—二叠纪花岗岩闪长岩；11—石炭纪黑云钾长花岗岩；12—断层；13—推断断层；14—地质界线；15—不整合界线；16—铌稀土多金属矿化蚀变带

Fig. 9 1:50000 aeromagnetic and aeroradiometric features sketch map of Hulunxibai area

a—Geological sketch map of Hulunxibai area; b—Aeroradiometric uranium anomaly contour map; c—Aeroradiometric potassium anomaly contour map; d—Aeroradiometric thorium anomaly contour map; e—Aeroradiometric total counting rate contour map; f—Aeromagnetic ΔT anomaly contour map; 1—Quaternary; 2—Cretaceous; 3—Triassic; 4—Permian; 5—Carboniferous; 6—Devonian; 7—Changcheng system; 8—Proterozoic; 9—Permian adamellite; 10—Permian granodiorite; 11—Carboniferous biotite K-feldspar granite; 12—Fault; 13—Inferred faults; 14—Geological boundary; 15—Unconformity; 16—Niobium and rare earth polymetallic mineralization zone

岩石孔隙度的增大,有利于后期热液的活动和矿质的析出,因此在此阶段可形成脉状的富矿体(翟裕生等,1993)。区内已知的铌金-铌稀土矿(化)体往往产于构造挤压形成的虚脱空间内或韧性变形带内劈理、片理和裂隙发育地段,即韧-脆性转换部位。

5.3 二叠纪二长花岗岩外接触带

北山地区内生矿产的成矿作用与岩浆侵入和期后热液活动关系密切(江思宏等,2001,2004;苗来成等,2014),岩浆不仅是成矿元素的源泉,由于岩浆热液的活动,给成矿物质的搬运、富集成矿创造了有利条件(翟裕生等,1993,2011)。区内已知

的稀有稀土矿(化)体、矿点均分布于二叠纪二长花岗岩外接触带上,反映了该部位成矿元素的沉淀与岩浆-构造活动密不可分,是铌-稀土多金属矿有利成矿部位。

5.4 放射性异常梯度带-磁异常位置

区内航磁、航放高背景及异常主要呈弧形分布于二叠纪岩体与二叠系内、外接触带上,主要呈带状、串珠状分布,图9显示航放计数率、钾、钍、铀高背景或异常均分布于海西晚期侵入岩,已知铌稀土多金属矿化蚀变带主要分布于高、低背景梯度带上,亦反映了岩体与地层接触位置是目标矿体产出

有利位置;而磁异常区则反映了铈稀土多金属矿成矿有利位置,如灰石山东北发现的铈稀土矿位于1:5万航磁异常最显著地段,辉森乌拉西铈金矿也发育于局部航磁异常区内,图2显示灰石山东北1:1万高精度磁法异常与矿体位置套合较好。从以上特征来看,放射性异常梯度带、磁异常为寻找铈稀土多金属矿的直接地球物理标志,放射性梯度带与磁异常串珠状展布特征则代表了矿化带、矿(化)体的空间分布位置。

5.5 岩石蚀变标志

铈稀土多金属矿主要产于二叠纪二长花岗岩外接触带的二叠系滨浅海相地层中,矿化蚀变带地表岩石明显具硅化、褐铁矿化、碳酸盐化,基本呈带状与周边断层方位一致。钻孔深部往往可见微细浸染状黄铁矿化,微细石英脉、方解石脉,局部可见方铅矿化。因此,与铈稀土多金属矿有关的蚀变主要为发育于断层和二叠纪酸性岩体接触附近的硅化、褐铁矿化、碳酸盐化蚀变带。

5.6 石英细网脉发育地段

已知的铈稀土多金属矿(化)体、矿点分布部位往往石英细网脉较为发育,脉体与矿体走向基本一致,由于单脉宽度基本不大于5 cm,脉体不连续,地表风化破碎容易致其分布凌乱,野外工作时极易被忽略。前人研究显示脉岩群往往是大量流体迁移的反映(罗照华等,2008),在小型张性裂隙内或应力分化转换处极易富集成矿(邱俊挺等,2011;张善明等,2018),因此在找矿过程中石英细网脉发育地段须高度重视。

6 结 论

(1)内蒙古呼伦西白地区铈稀土多金属矿产于二叠纪二长花岗岩外接触带上,矿体赋存于二叠系滨浅海相碳酸盐岩-变碎屑岩内,其形态、产状、规模同时受构造-岩浆活动控制,已知矿床(点)都具有特征航磁航放异常,与白云鄂博Nb-REE-Fe超大型矿床特征类似。

(2)区内铈稀土多金属矿的找矿标志为:二叠系滨浅海相碳酸盐岩-变碎屑岩地层、构造挤压形成的虚脱空间或韧-脆性转换部位、二叠纪二长花岗岩外接触带、放射性异常梯度带-磁异常位置、岩石蚀变(褐铁矿化、硅化、碳酸盐化等)和石英细网

脉发育地段。

致谢: 内蒙古第一地勘院苏永辉高级工程师、内蒙古地勘院袁晓鹏工程师在成文过程中给予了典型矿床成矿特征的指导;审稿专家及编辑部老师对论文提出了宝贵修改意见。在此一并致以诚挚的谢意!

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