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## 鄂尔多斯盆地西南缘洛河组下段含铀砂岩 锆石 U-Pb 年代学: 对岩石圈伸展作用的启示

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**摘要:**【研究目的】近年来, 砂岩型铀矿的研究成为地质学者讨论的热点, 对赋铀地层的研究尤为重要。【研究方法】文章以鄂尔多斯盆地西南缘镇原地区为例, 通过对下白垩统洛河组下段含铀砂岩碎屑锆石的 LA-ICP-MS U-Pb 测年分析, 探讨含铀砂岩的物质来源及其构造意义。【研究结果】结果表明, 下白垩统洛河组下段赋铀砂岩碎屑锆石 U-Pb 年龄主要分布在 166~370 Ma ( $n=49$ )、388~472 Ma ( $n=14$ )、1744~2150 Ma ( $n=14$ )、2241~2740 Ma ( $n=14$ )、615~1623 Ma ( $n=9$ ) 这 5 个年龄区间, 相应的峰值年龄为 272 Ma、427 Ma、1899 Ma 和 2493 Ma, 而 615~1623 Ma 无法确定峰值年龄。【结论】锆石 U-Pb 年龄谱对比分析表明, 研究区下白垩统洛河组下段含铀砂岩总体上具有来自北祁连造山带东段、北秦岭造山带、兴蒙造山带、贺兰山、华北板块基底及阿拉善地块 6 个物源区, 其中北秦岭造山带、北祁连造山带东段为主要物源区。碎屑锆石年龄显示, 奥陶纪鄂尔多斯盆地西南缘形成被动陆缘, 志留纪—泥盆纪转化为碰撞造山带, 石炭纪—二叠纪由造山带转化为沉积盆地, 侏罗纪—白垩纪形成稳定沉积地层为铀矿贮存提供有利空间。

**关键词:** 含铀砂岩; 下白垩统; 洛河组; 碎屑锆石 U-Pb; 镇原; 鄂尔多斯盆地; 矿产勘查工程

**创新点:** 通过研究发现, 研究区洛河组下段含铀砂岩的物源具有复杂性、多缘性, 其中北祁连造山带东段、北秦岭造山带为主要物源区。

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## Detrital zircon U-Pb geochronology of uranium-bearing sandstone in the lower member of Luohe Formation in the southwest margin of the Ordos Basin: Implications for the lithospheric extension

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**Abstract:** This paper is the result of mineral exploration engineering.

**[Objective]** In recent years, the study of sandstone-type uranium deposits has become a hotspot among geologists, especially the study of uranium-bearing strata. **[Methods]** Based on LA-ICP-MS U-Pb dating of detrital zircons from the lower part of the Lower Cretaceous Luohe Formation uranium-bearing sandstones in Zhenyuan area, southwestern Ordos Basin, the provenance and tectonic significance of uranium-bearing sandstones are discussed. **[Results]** The results show that the detrital zircon U-Pb ages of the lower member of the Lower Cretaceous Luohe Formation are mainly distributed in five age intervals: 166–370 Ma ( $n=49$ ), 388–472 Ma ( $n=14$ ), 1744–2150 Ma ( $n=14$ ), 2241–2740 Ma ( $n=14$ ) and 615–1623 Ma ( $n=9$ ), besides 615–1623 Ma whose peak age can not be determined, and the corresponding peak ages are 272 Ma, 427 Ma, 1899 Ma and 2493 Ma, respectively. **[Conclusions]** Comparative analysis of detrital zircon U-Pb age spectra shows that the lower member of Lower Cretaceous Luohe Formation in the study area has six source areas, including the eastern part of the North Qilian orogenic belt, the North Qinling orogenic belt, the Xingmeng orogenic belt, the Helan Mountain, the basement of the North China Plate and the Alxa Block, among which the eastern part of the North Qinling orogenic belt and the North Qilian orogenic belt are the main source areas. The detrital zircon ages show that the southwest margin of Ordos Basin formed a passive continental margin in Ordovician, the Silurian-Devonian transformed into a collisional orogenic belt, the Carboniferous-Permian transformed from an orogenic belt into a sedimentary basin, and the Jurassic-Cretaceous formed stable sedimentary strata to provide favorable space for uranium ore storage.

**Key words:** Uranium-bearing sandstone; Lower Cretaceous; Luohe formation; detrital zircon U-Pb; Zhenyuan; Ordos Basin; mineral exploration engineering

**Highlights:** The present study reveals that the source of uranium-bearing sandstones in the lower member of the Luohe Formation in the study area is complex and multi-marginal, with the eastern part of the North Qilian orogenic belt and the North Qinling orogenic belt as the main provenance.

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## 1 引 言

鄂尔多斯盆地西南部地区位于秦岭、祁连造山带结合部位,地理位置十分特殊,其不仅是古亚洲构造域和特提斯构造域交汇的重要区域,而且是“中亚造山带”东、西衔接处(罗顺社等,2017),其中生代的构造属性、原盆沉积范围、沉积物源等地质问题,一直是地质和能源工作者关注的热点(白云来等,2006;赵文智等,2006;赵俊峰等,2008;张珂等,2009;李子颖等,2022)。砂岩型铀矿作为新型能源矿产,近年来受到广大地质学者的关注。前人研究表明,鄂尔多斯东北部及准东地区主要的储铀层位于中生代侏罗纪碎屑岩中(焦养泉等,2005;王敏芳等,2006;张金带等,2010;金若时等,2014;何中波等,2018;冯晓曦等,2019;金若时和腾雪明,2022),金若时等(2014)与金若时和腾雪明(2022)认为侏罗纪时期的砂体利于 $U^{6+}$ 的还原、储集。而本次研究的鄂尔多斯盆地西南缘铀矿贮存在下白垩

统洛河组下段砂体中,同样具有铀元素储集成矿的有利条件(赵华雷等,2022)。笔者认为对鄂尔多斯盆地西南缘洛河组的物源及与周缘地体关系的研究具有重要意义。

研究表明,造山带和沉积盆地具有空间上互相依存、动力互换和物质互补的密切关系,均属大陆构造上的基本单元。造山带为盆地提供物源,而盆地中的碎屑沉积物不仅是盆地沉积、构造演化的重要标志,也是漫长地质历史中造山带、沉积盆地及周围环境相互作用的直接证据(陈世悦,2000)。碎屑锆石研究作为反映沉积物源的有效手段之一,通过碎屑锆石的成因及U-Pb年龄频率峰值能够示踪沉积物源区,揭示区域构造演化,为区域盆山演化研究提供证据(闫义等,2003)。

本文采用锆石LA-ICP-MS U-Pb测年分析技术,对鄂尔多斯盆地西南缘镇原地区下白垩统洛河组下段赋铀砂体碎屑锆石进行研究,结合前人对鄂尔多斯盆地西南缘大量的研究成果,分析锆石年龄分布特

征,建立鄂尔多斯盆地西南缘锆石年龄结构;对比周缘地质体,界定洛河组的沉积源区;试图恢复鄂尔多斯盆地西南缘古生代—中生代造山、沉积作用过程。

### 2 区域地质背景

鄂尔多斯盆地的地理范畴包括秦岭以北、阴山以南、吕梁山以西和贺兰山以东的广大地区。镇原

地区位于鄂尔多斯盆地西南缘,渭北隆起西段、伊陕斜坡和天环坳陷南部,处于北祁连造山带、秦岭—大别造山带与华北板块交互作用的地区;其西部以平凉为界,东至庆阳为界;南部以泾川为界,北部与环县相接(图1)。

鄂尔多斯盆地经历了早、中侏罗世内陆坳陷沉积之后,白垩纪盆地逐渐消亡。侏罗纪末期至早白垩世

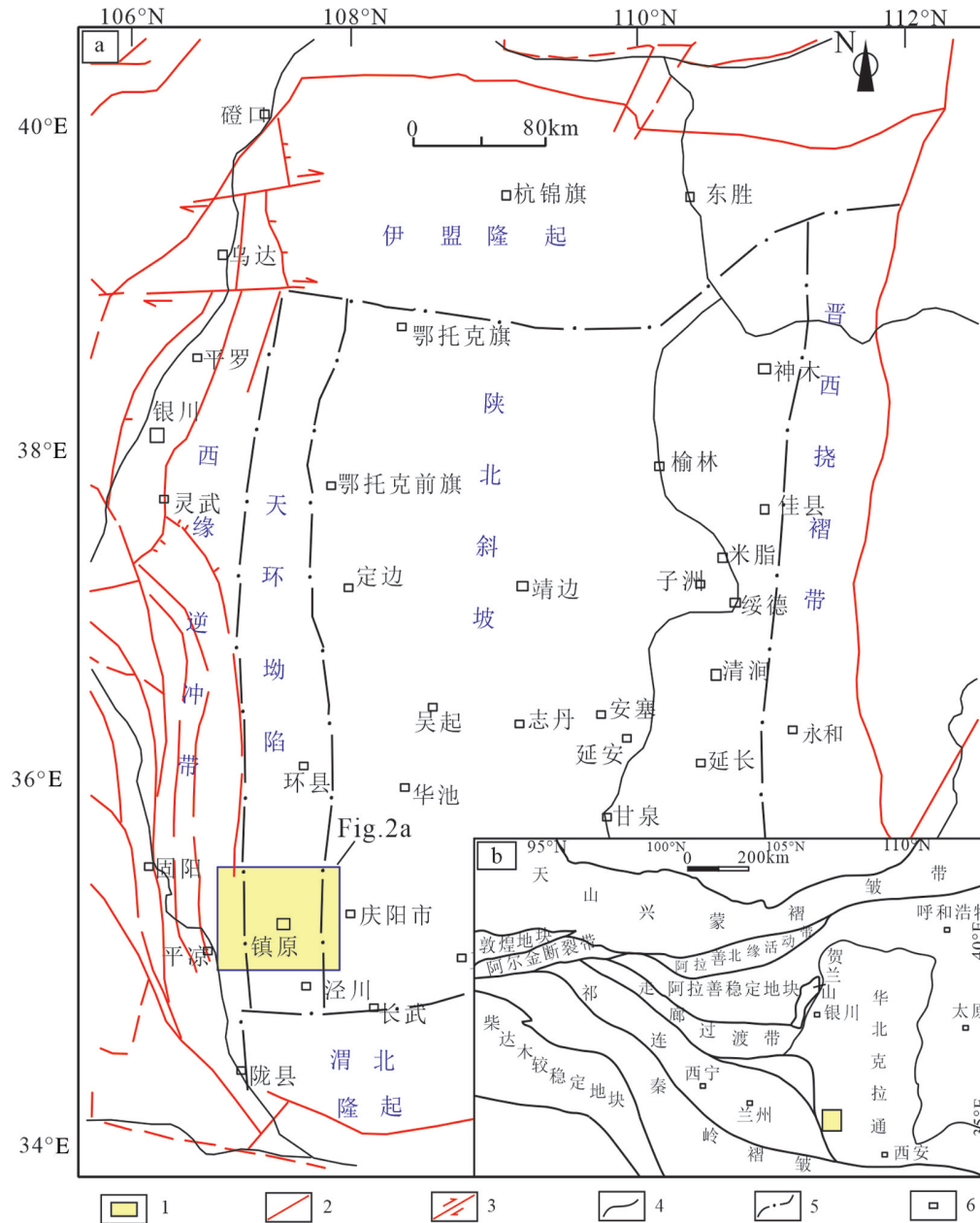


图1 鄂尔多斯盆地构造单元(a,据杨华等,2013)和鄂尔多斯盆地西南缘邻区大地构造位置(b,据刘池洋等,2006)  
 1—研究区;2—断层;3—走滑断层;4—一级构造分区线;5—二级构造分区线;6—地名  
 Fig1. Regional tectonic unit of Ordos Basin (a, after Yang Hua et al., 2013), geotectonic location in the southwest margin of Ordos Basin (b, after Liu Chiyang et al., 2006)

1—Research area; 2—Fault; 3—Strike slip fault; 4—First level construction zone line; 5—Secondary construction zone line; 6—Place name



受燕山运动的影响,盆地发生第三次拗陷。盆地内部构造相对简单,主要分为陕北斜坡(I)、伊盟隆起(II)、天环拗陷(III)三部分。天环拗陷为西陡东缓以镇原—环县—盐池—鄂托克旗为轴的不对称向斜结构的沉积拗陷区。这种区域构造特征,决定了白垩纪地层向西厚度增大、层位齐全的分布状况。

鄂尔多斯盆地白垩系自下而上划分为洛河组、环华池组、罗汉洞组和涇川组(郝诒纯等,1986)。洛

河组主要为一套风成沙漠相、河流湖泊相以及冲(洪)积物、泥石流和风化残积物沉积组合,主要岩石类型包括灰绿色中砂岩、细砂岩、青灰色粉砂质泥岩、泥岩和含砾砂岩以及砾岩,其中以砂岩为主。研究区第四系大面积覆盖,白垩系呈网脉状剥露于第四系冲洪积物之上,洛河组与涇川组呈不整合接触,与上覆环华池组呈整合接触(图2),古生代地层出露较少,主要出露于研究区西部。研究区西南部发育NE向逆

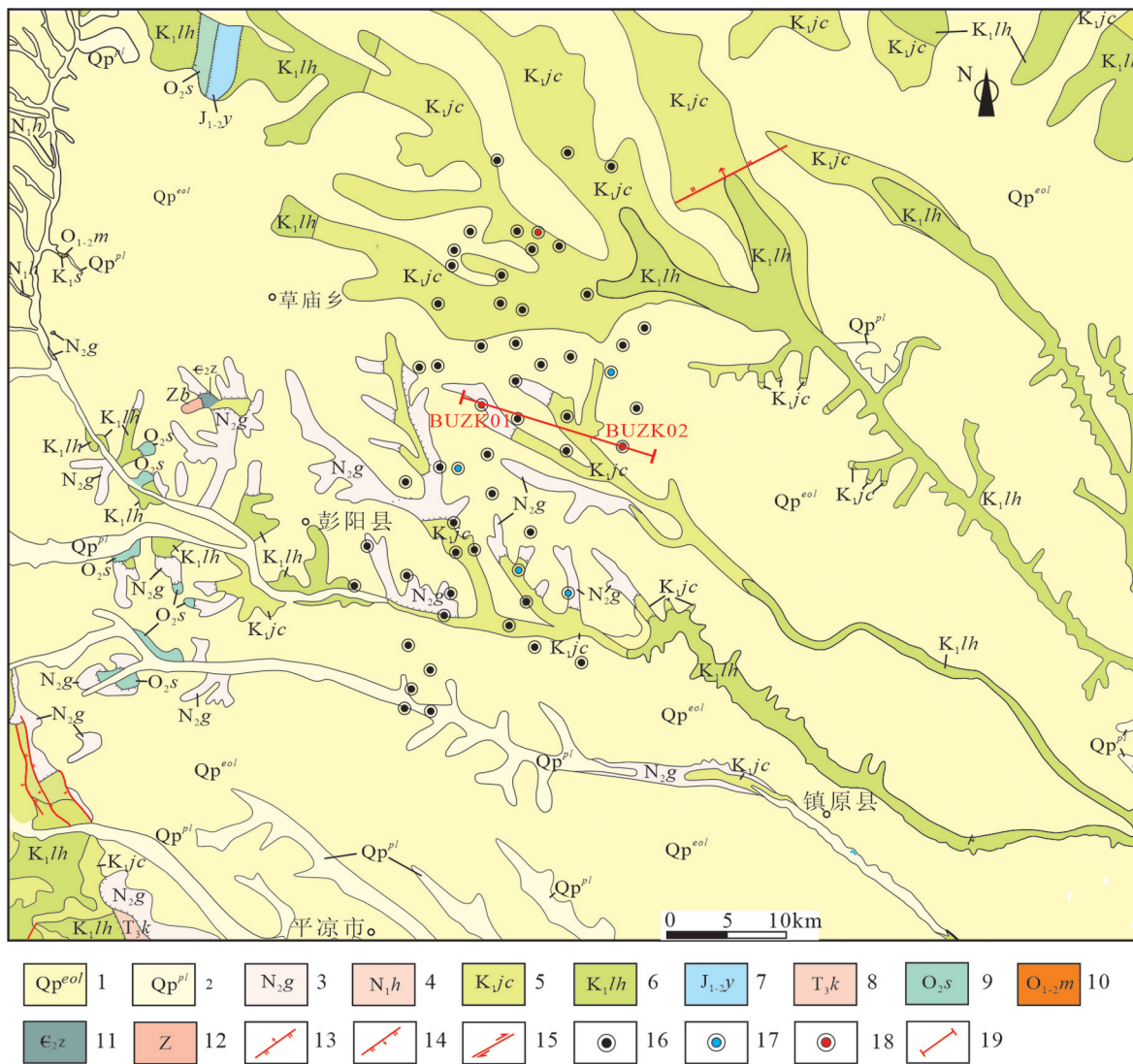


图2 研究区地质简图及采样井位

1—第四系风积砂;2—第四系冲洪积物;3—新近系干河沟组;4—新近系红柳河组;5—下白垩统涇川组;6—下白垩统洛河组;7—侏罗系延安组;8—上三叠统崆峒山组;9—中侏罗统三道沟组;10—中下奥陶统马家沟组;11—寒武系张夏组;12—震旦系;13—逆断层;14—正断层;15—走滑断层;16—石油钻孔;17—潜在铀矿化孔;18—铀矿工业钻孔;19—剖面线

Fig.2 Geologic map of the study area with well locations of samples

1—Aeolian sand of Quaternary; 2—Alluvial diluvium of Quaternary; 3—Ganhegou Formation of Neogene; 4—Hongliuhegou Formation of Neogene; 5—Jingchuan Formation of Early Cretaceous; 6—Luohe Formation of Early Cretaceous; 7—Yan'an Formation of Jurassic; 8—Kongtongshan Formation of Upper Triassic; 9—Sandaogou Formation of Middle Jurassic; 10—Majiagou Formation of Middle lower Ordovician; 11—Zhangxia Formation of Cambrian; 12—Sinian system; 13—Reverse fault; 14—Normal fault; 15—Strike slip fault; 16—Oil drilling; 17—Potential uranium mineralization borehole; 18—Uranium industry drilling; 19—Section line

冲推覆构造,且地势较高,地下流体易向NE向流动。

洛河组上段主要为沙漠相风成沉积为主,主要岩性为棕红—黄褐色中细砂岩,岩心见高角度板状、楔状交错层理,分选较好,磨圆次棱角状等沉积特征。洛河组中、下段主要为河流相沉积,中段为曲流河相沉积,下段为辫状河相沉积。洛河组中段岩性主要为浅砖红色、灰色、灰白色细砂岩与薄层中砂岩互层夹粉砂质泥岩,具交错层理,磨圆分选较差,根据岩性柱状图(图3)可知,洛河组上段砂体具有正旋回沉积特征,具有曲流河二元结构。洛河组下段岩性主要为黄褐色、灰绿色、灰色中砂岩与细砂岩互层,粒度较粗,槽状交错层理为主,偶见块状和水平层理,分选磨圆较差,具砂包泥正旋回沉积特征,灰绿色中砂岩流通性较好。

### 3 样品特征与分析方法

#### 3.1 采样位置与样品特征

本次工作主要集中于鄂尔多斯西南缘镇原地区,选取铀矿孔(BUZK02)(图2)洛河组下段含矿层位砂岩样品(ZK02-1055)进行碎屑锆石U-Pb测年(图3),岩性为灰绿色中砂岩,分选中等呈次棱角状,具有辫状河河道沉积特征。

#### 3.2 分析方法

锆石分选由河北省廊坊诚信地质服务有限公司完成。首先进行粉碎分选,分选出的锆石在双目镜下挑选,选择透明度较高、晶形较完好且内部无裂隙具有代表性的锆石进行制靶,锆石阴极发光图像、制靶工作由北京锆年领航科技有限公司完成。通过反射光、透射光及阴极发光图像综合分析,选择环带清晰以及晶型好、浑圆状的锆石样品进行测试。锆石LA-ICP-MS测试在天津地质调查中心实验室完成,利用激光剥蚀等离子体质谱仪(LA-ICP-MS)进行锆石U-Pb同位素测试。

激光剥蚀系统为New Wave UP213, ICP-MS为布鲁克M90。激光剥蚀过程中采用氦气作载气、氩气为补偿气以调节灵敏度,二者在进入ICP之前通过一个Y型接头混合。每个时间分辨分析数据包括20~30 s的空白信号和50 s的样品信号。对分析数据的离线处理(包括对样品和空白信号的选择、仪器灵敏度漂移校正、元素含量及U-Th-Pb同位素比值和年龄计算)采用软件ICPMSDataCal完成

(Liu et al., 2010)。锆石样品的U-Pb年龄谱和图绘制和年龄权重平均计算均采用Isoplot/Ex\_ver3完成(Ludwig, 2003)。本次测试剥蚀直径根据实际情况选择25  $\mu\text{m}$ 。

## 4 分析结果

本次测试BUZK02井含铀砂岩层位获得104颗碎屑锆石,谐和度>90%的一共有100颗,利用Isoplot软件处理这些谐和度较高的数据,并绘制锆石U-Pb谐和图及年龄分布直方图(图4)。

#### 4.1 锆石形貌学特征

锆石阴极发光(CL)图像(图5)显示,锆石粒径一般在60~100  $\mu\text{m}$ ,个别可达150  $\mu\text{m}$ ,晶形完整锆石较碎片状锆石略多,磨圆度从无磨圆到较好磨圆均有。根据锆石颗粒复杂多样的内部结构,可将其分为四类:第一类具清晰震荡环带,自形程度较好,表明其原始成因类型为典型岩浆锆石(如5、28等),部分锆石具有残留的核部,为继承核或捕获核(如13、16、43等),仍为岩浆成因锆石,约占40%;第二类为边部具或窄或宽的浅色增生边,为继承性变质锆石,是后期变质作用的产物,年龄值较大(如47、53、60、83等),约占20%;第三类占总数35%左右,无分带结构或具弱分带结构,或呈溶蚀结构,内部结构复杂不清,成因不明(如1、10、32、52、75等);第四类为颜色较深,无分带结构且内部浑浊不清,可能来自古老基底(如33、63、77等),约占10%;复杂多样的锆石形态,反映了不同成因的锆石类型,同时也反映了物源的复杂性。

#### 4.2 锆石年龄特征

镇原地区BUZK02-1055的碎屑锆石Pb、Th、U含量和Th/U比值及锆石U-Pb年龄见表1,锆石的Th含量为 $2 \times 10^{-6}$ ~ $1037 \times 10^{-6}$ , U含量为 $9 \times 10^{-6}$ ~ $966 \times 10^{-6}$ , Th/U比值为0.05~1.66,平均为0.6。一般情况下, Th/U > 0.4表现为岩浆成因, Th/U < 0.1表现为变质成因(Belousova et al., 2002; Crofu et al., 2003)。本次研究镇原地区BUZK02-1055碎屑锆石Th/U比值表明其主体为岩浆成因,部分锆石后期可能受变质事件改造的影响,少部分锆石表现为变质成因。根据锆石年龄谐和图和直方图(图4)可知,样品中锆石年龄分为166~370 Ma、388~472 Ma两个主峰年龄区间和1744~2150 Ma、2241~2740 Ma两个次峰年龄区



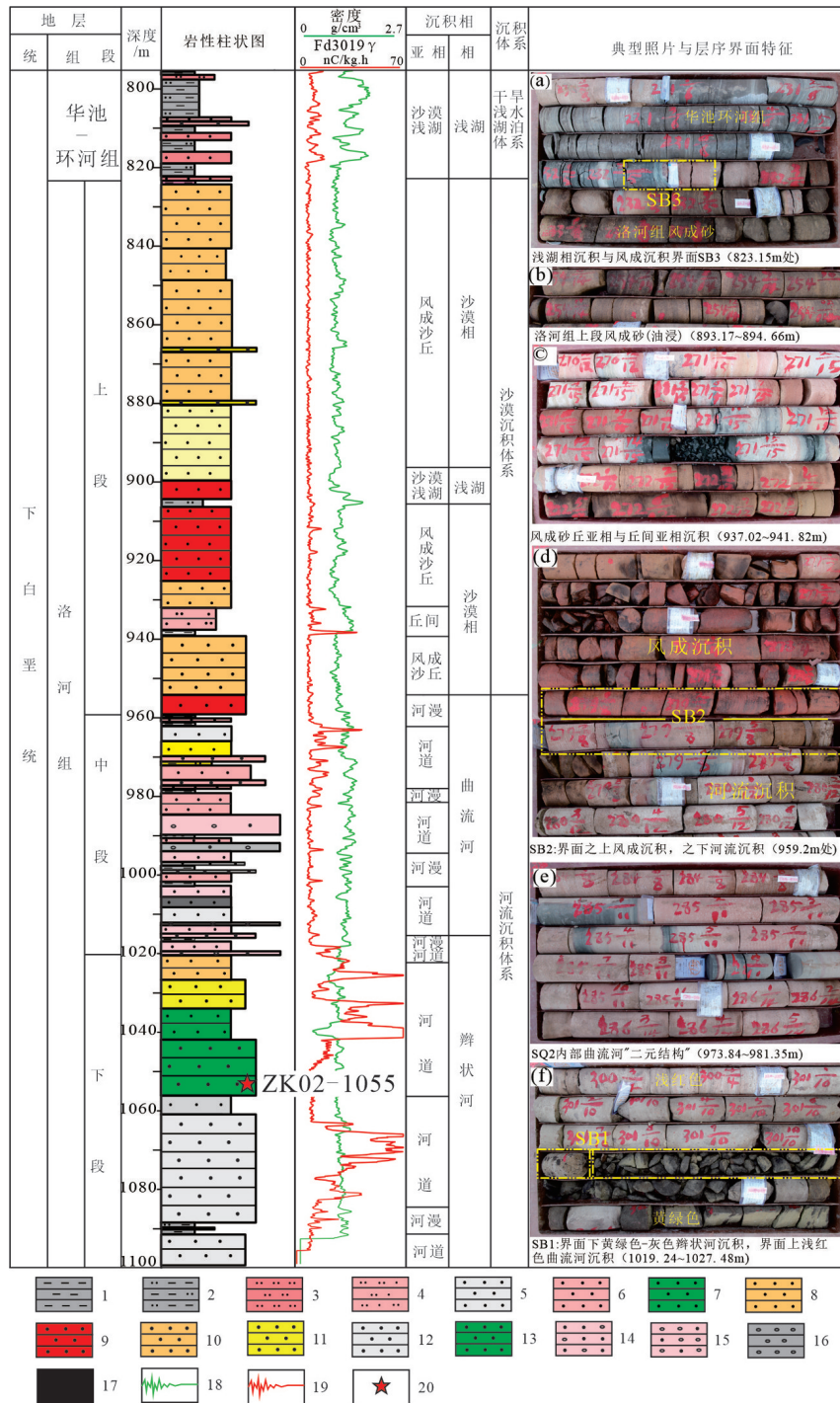


图3 鄂尔多斯盆地西南缘镇原地区洛河组综合柱状图

1—灰黑色泥岩;2—灰黑色粉砂质泥岩;3—砖红色粉砂岩;4—肉红色粉细砂岩;5—灰白色细砂岩;6—肉红色细砂岩;7—灰绿色细砂岩;8—黄褐色细砂岩;9—红色中砂岩;10—黄褐色中砂岩;11—黄色中砂岩;12—灰白色粗砂岩;13—灰绿色粗砂岩;14—肉红色含砾砂岩;15—肉红色含砂砾岩;16—灰黑色砾岩;17—煤层;18—自然电位;19—自然伽马;20—采样位置

Fig.3 Comprehensive column of Luohe Formation in Zhenyuan area in southwestern Ordos Basin

1—Grey black mudstone; 2—Grey black silty mudstone; 3—Brick red siltstone; 4—Flesh red siltstone; 5—Grey white fine sandstone; 6—Flesh red fine sandstone; 7—Grey green fine sandstone; 8—Yellowish brown fine sandstone; 9—Red medium sandstone; 10—Yellowish brown medium sandstone; 11—Yellow medium sandstone; 12—Gray white coarse sandstone; 13—Grey green coarse sandstone; 14—Flesh red gravelly sandstone; 15—Flesh red sandy conglomerate; 16—Grey black conglomerate; 17—Coal seam; 18—Spontaneous potential; 19—Natural gamma ray; 20—Sampling location

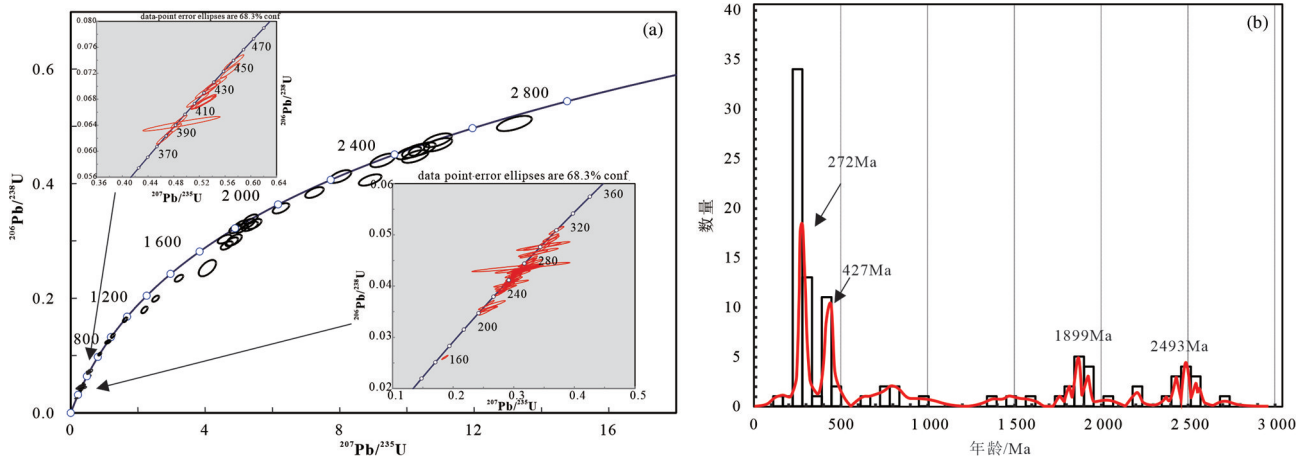


图4 (a) BUZK02-1055 碎屑锆石U-Pb年龄谐和图;(b)锆石年龄谱  
Fig.4 (a) U-Pb concordia diagrams of detrital zircons of BUZK02-1055;(b) Zircon age spectrum

间,相应的峰值年龄为272 Ma、427 Ma、1899 Ma和2493 Ma,另外还有9颗中新元古代的锆石。样品存在一颗95 Ma的锆石,推断其为混入较晚期的碎屑,不具代表性,此处不做重点研究。

## 5 讨论

### 5.1 物源分析

鄂尔多斯盆地西南缘处于多个块体结合区域,沉积-构造演化与邻区造山带的构造运动密切相关。研究区相邻的北祁连造山带、北秦岭造山带、阿拉善地块古陆以及华北克拉通北缘都有可能为研究区提供物源。

166~370 Ma:峰值年龄为272 Ma,该区间锆石颗粒49颗,占总数的49%,Th/U比值为0.19~1.66,表现为岩浆成因锆石,部分后期经历变质事件改

造,对应中侏罗世一晚泥盆世的岩浆-构造热事件。由于该年龄区间的锆石数量较多,以作为本次研究的重点。西伯利亚板块南缘和华北板块北缘的索伦缝合带的最终闭合时间可能在230~310 Ma(陈斌等,2001);兴蒙造山带平泉地区存在大量250~330 Ma年龄的锆石(马收先等,2011);陈岳龙等(2012)也对兴蒙造山带中的碎屑岩、变质岩、变岩浆岩等进行了锆石U-Pb测年,结果显示250~350 Ma形成的锆石普遍存在(图6f);位于狼山岩体内部发育多个海西期花岗岩体(彭润民等,2007),陈登超等(2010)测其K-Ar年龄为267~302 Ma;东升庙岩体中二长花岗岩体年龄为(259.4±3.3)Ma(吴亚飞等,2013);霍各乞岩体中的辉长-闪长岩年龄为(273.9±1.2)Ma(皮桥辉等,2010);邹雷等(2019)获得东阿拉善波罗斯坦庙杂岩U-Pb年龄为



图5 镇原地区BUZK02井洛河组下段含铀砂岩典型碎屑锆石阴极发光照片(年龄单位:Ma)

Fig.5 CL images of detrital zircons from uranium-bearing sandstone from the lower part of Luohe Formation of BUZK02 in Zhenyuan area (age unit: Ma)

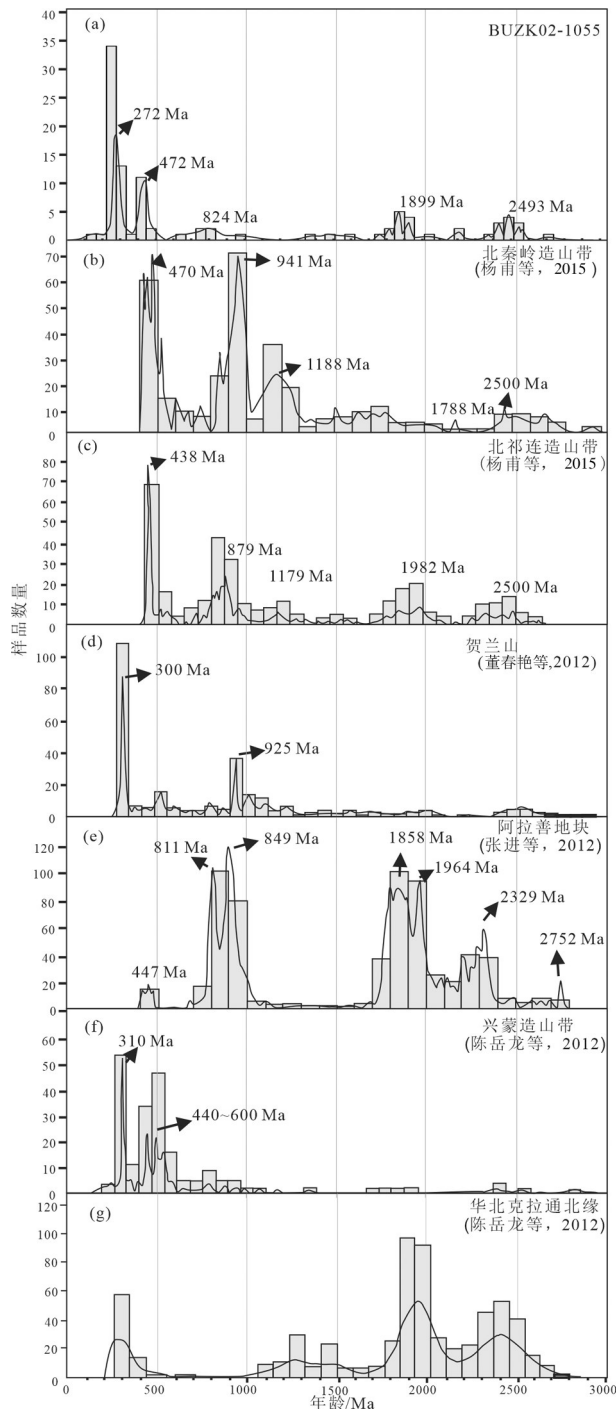


图6 鄂尔多斯盆地西南缘下白垩统碎屑锆石U-Pb年龄谱与周缘构造单元结晶岩体晚古生代以前岩体对比图

Fig.6 Comparison of U-Pb ages spectra of detrital zircons of the Lower Cretaceous sandstones in the southwestern Ordos Basin and the pre-Late Paleozoic crystallization rocks of the adjacent potential sources

242~284 Ma(表2);这些年龄与该组峰值年龄相似,推断索伦缝合带闭合过程中,兴蒙造山带中大量火山岩浆产物随着造山运动而剥蚀搬运,向研究区提供物源。一直以来,二叠纪花岗岩在秦岭造山带中的记载较少,研究发现西秦岭江里沟花岗岩存在( $264.0\pm 1.4$ )Ma的加权平均年龄(孙小攀等,2013),属中二叠世晚期,证明了在中二叠世晚期存在地壳增厚背景下的构造岩浆活动事件。秦岭造山带在石炭纪末—二叠纪进入碰撞阶段,北秦岭构造带有较大幅度的隆升,其北侧相邻地区演变为沉积盆地(陈世悦,2000)。碰撞造山引发的一系列岩浆事件与本组锆石年龄具有较好的对应性。综上所述,该时期鄂尔多斯盆地西南缘物源主要来自兴蒙造山带,不排除北秦岭造山带提供物源的可能。

223~250 Ma的锆石可能来自秦岭造山带内广泛发育的印支期岩浆热事件(秦江峰等,2010;骆必继等,2013),西秦岭糜署岭花岗岩年龄为237 Ma(李永军等,2004),草关地区花岗岩年龄为205 Ma(李永军等,2004),西秦岭江里沟复式岩体年龄为( $229.1\pm 1.8$ )Ma(路东宇等,2017),夏河地区岩体年龄为238 Ma(金惟浚等,2005),秦岭造山带黑沟峡火山岩Sm-Nd等时线年龄为242 Ma,Rb-Sr等时线年龄为221 Ma,认为其火山岩遭受碰撞-变质时间为221~242 Ma(李曙光等,1996)。

388~472 Ma:峰值年龄为427 Ma,该区间锆石颗粒14颗,占总数的14%,该组锆石年龄较为集中,Th/U比值为0.30~1.07,表现出典型岩浆锆石成因,锆石具清晰的韵律环带结构,对应中奥陶世—晚志留世早古生代构造岩浆热事件。早古生代是秦岭和祁连造山带最主要的洋陆转化阶段(夏林圻等,1996,2001;张国伟等,2001),鄂尔多斯西缘米钵山组的物源来自北祁连造山带(黄喜峰等,2009),表明北祁连造山带向阿拉善地块和华北板块俯冲碰撞提供物源(张进等,2012;程先钰等,2019)。北秦岭罗汉寺岩群辉长岩脉成岩年龄为( $475\pm 4$ )Ma(刘军锋等,2007),秦岭、祁连山造山带结合部位广泛出露呈近东西向产出的有川草铺花岗岩体( $434\pm 10$ )Ma(Zhang et al., 2006)、阎家店闪长岩体( $440.2\pm 0.92$ )Ma(裴先治等,2007)、( $441\pm 10$ )Ma(Zhang et al.,



**表1 鄂尔多斯盆地西南缘镇原地区洛河组下段BUZK02-1055碎屑锆石LA-ICP-MS U-Pb年龄测试结果**  
**Table1 Detrital zircon LA-ICP-MS U-Pb dating results of the lower part of Luohe Formation of BUZK02-1055 in Zhenyuan area in southwestern Ordos Basin**

样品号	含量/10 <sup>6</sup>			Th/U	同位素比值						表面年龄/Ma					
	Pb	Th	U		<sup>207</sup> Pb/ <sup>206</sup> Pb		<sup>207</sup> Pb/ <sup>235</sup> U		<sup>206</sup> Pb/ <sup>238</sup> U		<sup>207</sup> Pb/ <sup>206</sup> Pb		<sup>207</sup> Pb/ <sup>235</sup> U		<sup>206</sup> Pb/ <sup>238</sup> U	
					1σ	1σ	1σ	1σ	1σ	1σ	1σ	1σ				
A1055.1	2	34	43	0.78	0.0527	0.0031	0.3468	0.0207	0.0477	0.0005	315	136	302	18	301	3
A1055.2	8	90	183	0.49	0.0527	0.0013	0.3249	0.0085	0.0447	0.0005	315	58	286	8	282	3
A1055.3	107	197	296	0.66	0.1140	0.0016	5.0721	0.0806	0.3227	0.0034	1864	26	1831	29	1803	19
A1055.4	322	463	600	0.77	0.1626	0.0023	10.2575	0.1662	0.4576	0.0050	2483	24	2458	40	2429	26
A1055.5	23	444	485	0.92	0.0530	0.0034	0.2957	0.0189	0.0405	0.0004	328	146	263	17	256	3
A1055.6	12	117	161	0.72	0.0567	0.0010	0.5284	0.0104	0.0676	0.0007	478	41	431	8	422	4
A1055.7	19	178	248	0.72	0.0561	0.0011	0.5447	0.0116	0.0704	0.0007	458	45	442	9	438	5
A1055.8	32	142	178	0.80	0.0721	0.0011	1.6161	0.0271	0.1626	0.0017	989	31	976	16	971	10
A1055.9	21	48	48	1.02	0.1269	0.0018	6.2505	0.1005	0.3572	0.0037	2056	26	2012	32	1969	20
A1055.10	6	104	138	0.75	0.0529	0.0014	0.2873	0.0078	0.0394	0.0004	323	61	256	7	249	3
A1055.11	4	54	87	0.63	0.0515	0.0010	0.2958	0.0059	0.0417	0.0004	262	44	263	5	263	3
A1055.12	130	324	325	1.00	0.1186	0.0020	5.3296	0.0966	0.3259	0.0034	1936	29	1874	34	1818	19
A1055.13	1	10	19	0.50	0.0553	0.0011	0.5173	0.0113	0.0679	0.0007	423	46	423	9	423	5
A1055.14	16	144	194	0.74	0.0569	0.0009	0.5681	0.0098	0.0724	0.0007	487	35	457	8	451	5
A1055.15	4	37	97	0.38	0.0522	0.0008	0.2815	0.0047	0.0391	0.0004	294	35	252	4	247	3
A1055.16	3	43	62	0.70	0.0534	0.0026	0.3199	0.0160	0.0434	0.0005	347	112	282	14	274	3
A1055.17	12	236	271	0.87	0.0528	0.0023	0.2566	0.0095	0.0353	0.0006	318	101	232	9	223	4
A1055.18	10	65	133	0.49	0.0565	0.0009	0.5738	0.0101	0.0737	0.0007	470	37	460	8	458	5
A1055.19	4	52	93	0.56	0.0517	0.0019	0.2853	0.0107	0.0400	0.0004	273	83	255	10	253	3
A1055.20	36	83	185	0.45	0.2073	0.0047	3.9428	0.1023	0.1379	0.0021	2885	37	1623	42	833	13
A1055.21	4	39	75	0.52	0.0529	0.0044	0.3487	0.0289	0.0478	0.0006	324	187	304	25	301	4
A1055.22	25	499	547	0.91	0.0527	0.0020	0.2875	0.0112	0.0396	0.0004	315	87	257	10	250	3
A1055.23	20	235	451	0.52	0.0522	0.0012	0.3025	0.0072	0.0420	0.0004	296	51	268	6	265	3
A1055.24	18	45	68	0.66	0.0992	0.0014	3.2165	0.0515	0.2351	0.0025	1610	27	1461	23	1361	14
A1055.25	19	324	362	0.89	0.0522	0.0023	0.3331	0.0150	0.0462	0.0005	296	101	292	13	291	3
A1055.26	34	134	726	0.19	0.0539	0.0009	0.3646	0.0067	0.0490	0.0005	368	39	316	6	309	3
A1055.27	5	25	33	0.77	0.0654	0.0011	1.1282	0.0201	0.1252	0.0013	786	35	767	14	760	8
A1055.28	4	26	57	0.46	0.0566	0.0013	0.5264	0.0130	0.0675	0.0007	474	51	429	11	421	4
A1055.29	37	1037	689	1.50	0.0532	0.0014	0.3021	0.0078	0.0412	0.0004	339	58	268	7	260	3
A1055.30	5	53	96	0.56	0.0532	0.0014	0.3568	0.0097	0.0487	0.0005	337	58	310	8	306	3
A1055.31	8	55	115	0.48	0.0555	0.0011	0.5304	0.0110	0.0693	0.0007	434	42	432	9	432	5
A1055.32	12	226	231	0.98	0.0519	0.0044	0.3094	0.0264	0.0433	0.0005	279	194	274	23	273	3
A1055.33	79	78	133	0.58	0.1893	0.0027	13.1883	0.2123	0.5053	0.0054	2736	23	2693	43	2636	28
A1055.34	10	133	264	0.50	0.0516	0.0008	0.2610	0.0044	0.0367	0.0004	269	36	236	4	232	2
A1055.35	74	131	197	0.66	0.1155	0.0017	5.2634	0.0851	0.3306	0.0034	1887	26	1863	30	1841	19
A1055.36	7	82	145	0.57	0.0531	0.0018	0.3105	0.0110	0.0424	0.0004	334	78	275	10	268	3
A1055.37	70	95	133	0.72	0.1659	0.0023	10.2427	0.1592	0.4477	0.0045	2517	24	2457	38	2385	24
A1055.38	13	169	296	0.57	0.0517	0.0017	0.3017	0.0105	0.0423	0.0004	274	77	268	9	267	3
A1055.39	22	243	288	0.84	0.0564	0.0012	0.5264	0.0130	0.0676	0.0007	470	48	429	11	422	4
A1055.40	6	69	139	0.50	0.0532	0.0012	0.3159	0.0076	0.0431	0.0004	337	53	279	7	272	3
A1055.41	10	13	28	0.48	0.1119	0.0016	5.0582	0.0804	0.3278	0.0034	1831	26	1829	29	1827	19
A1055.42	231	355	483	0.73	0.1588	0.0022	8.9015	0.1375	0.4067	0.0040	2442	24	2328	36	2200	22
A1055.43	11	65	178	0.37	0.0544	0.0009	0.4656	0.0081	0.0621	0.0007	388	36	388	7	388	4
A1055.44	14	106	328	0.32	0.0529	0.0012	0.3228	0.0076	0.0442	0.0005	327	51	284	7	279	3
A1055.45	182	70	602	0.12	0.1167	0.0017	4.8564	0.0935	0.3019	0.0040	1906	26	1795	35	1701	23
A1055.46	23	467	357	1.31	0.0528	0.0011	0.3705	0.0081	0.0509	0.0005	321	46	320	7	320	3
A1055.47	248	158	966	0.16	0.1166	0.0017	4.0622	0.1052	0.2528	0.0058	1904	25	1647	43	1453	34
A1055.48	6	106	129	0.82	0.0517	0.0096	0.3118	0.0533	0.0437	0.0006	274	427	276	47	276	4
A1055.49	127	197	536	0.37	0.1712	0.0024	4.8102	0.1181	0.2037	0.0042	2570	24	1787	44	1195	25
A1055.50	10	50	76	0.67	0.0646	0.0018	1.1044	0.0325	0.1240	0.0013	761	60	755	22	754	8

续表1

样品号	含量/ $10^{-6}$			Th/U	同位素比值						表面年龄/Ma					
	Pb	Th	U		$^{207}\text{Pb}/^{206}\text{Pb}$	$1\sigma$	$^{207}\text{Pb}/^{235}\text{U}$	$1\sigma$	$^{206}\text{Pb}/^{238}\text{U}$	$1\sigma$	$^{207}\text{Pb}/^{206}\text{Pb}$	$1\sigma$	$^{207}\text{Pb}/^{235}\text{U}$	$1\sigma$	$^{206}\text{Pb}/^{238}\text{U}$	$1\sigma$
A1055.51	20	110	288	0.38	0.0561	0.0008	0.5392	0.0088	0.0697	0.0007	455	33	438	7	435	5
A1055.52	3	36	96	0.37	0.0518	0.0009	0.1863	0.0035	0.0261	0.0003	277	39	174	3	166	2
A1055.53	75	139	356	0.39	0.0917	0.0013	2.5230	0.0420	0.1994	0.0022	1462	27	1279	21	1172	13
A1055.54	7	74	158	0.47	0.0525	0.0015	0.2949	0.0090	0.0407	0.0004	307	67	262	8	257	3
A1055.55	180	167	369	0.45	0.1526	0.0022	9.2759	0.1513	0.4408	0.0048	2376	24	2366	39	2354	26
A1055.56	6	75	119	0.63	0.0526	0.0014	0.3121	0.0086	0.0430	0.0005	312	59	276	8	272	3
A1055.57	1	7	9	0.73	0.0673	0.0011	1.2497	0.0226	0.1347	0.0015	847	34	823	15	814	9
A1055.58	28	28	62	0.45	0.1408	0.0020	8.0122	0.1334	0.4127	0.0045	2237	25	2232	37	2227	25
A1055.59	16	157	394	0.40	0.0522	0.0010	0.2784	0.0058	0.0387	0.0004	294	45	249	5	245	3
A1055.60	50	47	121	0.39	0.1366	0.0019	7.2465	0.1124	0.3848	0.0039	2184	25	2142	33	2099	21
A1055.61	4	41	99	0.42	0.0518	0.0010	0.3052	0.0062	0.0427	0.0005	279	43	270	6	270	3
A1055.62	133	187	234	0.80	0.1661	0.0024	10.9030	0.1735	0.4760	0.0049	2519	24	2515	40	2510	26
A1055.63	213	256	396	0.65	0.1645	0.0023	10.4488	0.1639	0.4607	0.0048	2503	24	2475	39	2442	25
A1055.64	60	391	808	0.48	0.0567	0.0016	0.5497	0.0177	0.0703	0.0007	480	63	445	14	438	5
A1055.65	20	288	424	0.68	0.0517	0.0010	0.2930	0.0061	0.0411	0.0005	272	45	261	5	260	3
A1055.66	11	159	275	0.58	0.0529	0.0019	0.2587	0.0096	0.0355	0.0004	323	83	234	9	225	2
A1055.67	51	15	299	0.05	0.0883	0.0012	2.1903	0.0386	0.1799	0.0024	1389	27	1178	21	1067	14
A1055.68	33	482	692	0.70	0.0530	0.0029	0.3153	0.0174	0.0431	0.0005	330	125	278	15	272	3
A1055.69	5	63	119	0.53	0.0526	0.0014	0.3131	0.0086	0.0432	0.0004	312	60	277	8	272	3
A1055.70	4	34	83	0.41	0.0529	0.0010	0.3023	0.0062	0.0414	0.0004	326	43	268	6	262	3
A1055.71	4	72	75	0.96	0.0516	0.0009	0.2933	0.0054	0.0412	0.0004	270	39	261	5	260	3
A1055.72	19	152	393	0.39	0.0530	0.0033	0.3405	0.0216	0.0466	0.0005	330	141	298	19	293	3
A1055.73	4	52	78	0.66	0.0530	0.0012	0.3611	0.0086	0.0494	0.0005	329	52	313	7	311	3
A1055.74	4	95	79	1.20	0.0527	0.0012	0.3256	0.0075	0.0448	0.0005	316	50	286	7	283	3
A1055.75	9	44	145	0.30	0.0549	0.0009	0.4793	0.0085	0.0633	0.0006	410	37	398	7	396	4
A1055.76	86	142	240	0.59	0.1105	0.0016	4.9002	0.0775	0.3217	0.0033	1807	26	1802	29	1798	18
A1055.77	106	180	190	0.95	0.1606	0.0023	10.0495	0.1544	0.4537	0.0045	2462	24	2439	37	2412	24
A1055.78	26	532	438	1.22	0.0535	0.0008	0.3508	0.0060	0.0476	0.0005	349	36	305	5	300	3
A1055.79	4	77	85	0.91	0.0519	0.0010	0.3064	0.0061	0.0428	0.0005	281	43	271	5	270	3
A1055.80	9	88	186	0.47	0.0529	0.0009	0.3293	0.0063	0.0451	0.0005	327	39	289	6	284	3
A1055.81	59	49	168	0.29	0.1148	0.0017	5.3499	0.0849	0.3381	0.0035	1876	26	1877	30	1877	19
A1055.82	5	46	96	0.48	0.0534	0.0022	0.3275	0.0135	0.0445	0.0005	346	92	288	12	281	3
A1055.83	30	49	100	0.49	0.1689	0.0024	5.8079	0.1244	0.2495	0.0043	2546	24	1948	42	1436	24
A1055.84	15	2	47	0.04	0.1200	0.0017	5.4568	0.0874	0.3299	0.0036	1956	25	1894	30	1838	20
A1055.85	1	24	23	1.03	0.0525	0.0010	0.2935	0.0058	0.0405	0.0005	307	43	261	5	256	3
A1055.86	107	113	413	0.27	0.1607	0.0023	5.1064	0.0907	0.2305	0.0030	2463	24	1837	33	1337	18
A1055.87	8	153	498	0.31	0.0569	0.0009	0.1160	0.0020	0.0148	0.0002	488	34	111	2	95	1
A1055.88	75	63	142	0.45	0.1692	0.0024	10.9201	0.1713	0.4681	0.0047	2550	24	2516	39	2475	25
A1055.89	14	239	263	0.91	0.0523	0.0010	0.3115	0.0062	0.0432	0.0005	300	43	275	6	273	3
A1055.90	8	152	164	0.93	0.0522	0.0013	0.3016	0.0077	0.0419	0.0004	296	55	268	7	264	3
A1055.91	7	190	115	1.66	0.0529	0.0026	0.3174	0.0161	0.0435	0.0005	325	111	280	14	274	3
A1055.92	168	259	501	0.52	0.1093	0.0019	4.5714	0.0836	0.3034	0.0031	1788	31	1744	32	1708	18
A1055.93	13	140	280	0.50	0.0528	0.0009	0.3114	0.0058	0.0428	0.0004	319	40	275	5	270	3
A1055.94	98	543	830	0.65	0.0607	0.0013	0.8613	0.0189	0.1029	0.0011	629	45	631	14	631	7
A1055.95	35	72	103	0.70	0.1151	0.0016	4.6373	0.0728	0.2923	0.0030	1881	26	1756	28	1653	17
A1055.96	1	15	14	1.07	0.0556	0.0045	0.4920	0.0397	0.0642	0.0007	435	180	406	33	401	5
A1055.97	3	45	78	0.58	0.0528	0.0028	0.2639	0.0146	0.0363	0.0006	318	120	238	13	230	4
A1055.98	8	104	180	0.58	0.0514	0.0009	0.3004	0.0058	0.0424	0.0004	257	42	267	5	268	3
A1055.99	143	63	476	0.13	0.1162	0.0017	4.7868	0.0766	0.2987	0.0032	1899	26	1783	29	1685	18
A1055.100	8	46	124	0.38	0.0548	0.0008	0.4882	0.0082	0.0646	0.0007	406	34	404	7	403	4

2006); 祁连山东段的陇山岩群内发育较多早古生代辉绿岩墙,成岩年龄为 440.9 Ma,变质年龄为 413 Ma; 祁连造山带靖远地区清凹山石英闪长岩 K-Ar 年龄为 402~445 Ma(王金荣, 2006); 李猛等(2015)在对北祁连造山带肃南地区阴沟群粗砂岩碎屑锆石进行 LA-ICP-MS U-Pb 年龄测定时,分别获得最小谐和年龄(425±2)Ma 和(425±5)Ma; 北秦岭造山带同样经历了该期岩浆热事件(图 6b),北秦岭西段和北祁连东端均发育 391~450 Ma 俯冲碰撞型花岗岩(徐学义等, 2008b),孟祥舒等(2017)获得秦—祁结合部糜棱岩化闪长岩 U-Pb 年龄为(418±3.2)Ma(表 2)。据此笔者推断鄂尔多斯盆地西南缘早古生代的物源主要来自北祁连造山带和北秦岭造山带。

1744~2150 Ma: 峰值年龄为 1899 Ma, 该区间锆石颗粒 14 颗, 占总数的 14%, Th/U 比值为 0.04~1.02, 主体表现为岩浆成因锆石, 部分为变质成因锆石。研究认为华北板块发生有 1900~2300 Ma 的陆内造山, 记录了华北克拉通裂谷形成到闭合的演化过程(翟明国, 2004), 表明了华北板块在中元古代可能为周边地区提供物源。华北克拉通西北缘孔兹岩系大量发育该期岩体(陈岳龙等, 2012), 其同碰撞期的花岗岩锆石 U-Pb 年龄为(1958±34)Ma(李正辉等, 2013; Dan et al., 2014), 千里山孔兹岩系年龄集中在 2000~2300 Ma(Yin et al., 2009)。阿拉善地块同样发育古元古代构造热事件(图 6e)(耿元生等, 2007; 周喜文和耿元生, 2010), 李俊建等(2004)获得巴彦乌拉片麻状花岗闪长岩 U-Pb 年龄为 2080 Ma, 古元古代岩体是阿拉善地块重要的基底(张进等, 2012), 北祁连造山带陇县白家沟花岗岩存在(1846±32)Ma 的加权平均年龄(尤佳等, 2014)(表 2); 杨甫等(2015)总结北祁连造山带古元古代晚期存在 1982 Ma 的峰值年龄等(图 6c); 徐欢(2019)获得华北板块西南缘陇县钾长花岗岩 U-Pb 年龄为(2018±16)Ma。综上所述, 阿拉善地块、华北板块和北祁连造山带均有可能为鄂尔多斯盆地西南缘提供古元古代碎屑物源。

2241~2740 Ma: 峰值年龄为 2493 Ma, 该区间锆石颗粒 14 颗, 占总数的 14%, Th/U 比值为 0.27~0.95, 表现为岩浆成因锆石, 后期经历变质事件, 锆石颜色较深, 内部结构浑浊不清, 且锆石磨圆度较

好, 体现了古老锆石多期变质, 多次搬运的特征。华北地块内部 2500 Ma 时期的构造热事件较为发育(林少泽等, 2019; 李立兴等, 2022), 沈其韩等(2005)在华北板块内部发现了古元古代 2500~2550 Ma 的峰值年龄; 张维杰等(2000)在固阳获得英云闪长岩锆石 U-Pb 年龄为(2440±35)Ma; 陶继雄和胡凤翔(2002)在固阳西红山乡获得石英闪长岩单颗粒锆石 U-Pb 年龄为 2575~2676 Ma, 祁连地区同样存在 1900 Ma、2350 Ma、2500 Ma 的构造岩浆峰值年龄(何艳红等, 2005)(表 2), 宫江华等(2012)获得了阿拉善地块北大山岩体(2496±11)Ma 的谐和年龄, 张进等(2012)获得了阿拉善地块内部古元古代 2329 Ma 的峰值年龄(图 6e), 且存在与本次碎屑锆石中最老年龄相似的锆石年龄(耿元生等, 2006)。综上所述, 华北板块基底结晶岩系、北祁连造山带和阿拉善地块的岩浆热事件与本次研究的古元古代—新太古代的年龄具有较好的对应性, 推测新太古代华北板块老基底、祁连造山带以及阿拉善地块古老基底为中生代鄂尔多斯盆地西南缘提供物源。

本次研究存在 9 颗 615~1623 Ma 的锆石颗粒, 与格林威尔造山运动具有较好的对应性。新元古代岩浆热事件在鄂尔多斯西南缘周缘地区均有发生(周喜文和耿元生, 2010)。

秦岭、祁连造山带结合部位的新元古代岩体较多, 如木其滩岩组斜长角闪岩(762.5±4.6)Ma(张张国等, 2011)、祁连山东段兴隆山群火山岩 723~824 Ma(徐学义等, 2008a)、西秦岭新阳花岗质片麻岩(935.5±3.1)Ma(刘会彬等, 2006)(表 2)。杨甫等(2015)(图 6b、c)提出北秦岭、北祁连造山带均发育中元古代锆石, 此外有报道称贺兰山地区同样可以为研究区提供中—新元古代锆石(图 6d)(董春艳等, 2012)。由于锆石颗粒较少, 且较为分散, 推测物源区主要为秦祁造山带, 同时不排除贺兰山地区向研究区提供物源的可能。

## 5.2 盆山演化过程

本次研究的鄂尔多斯盆地西南缘洛河组碎屑锆石年龄与周缘地区发生的构造运动事件有着较好的对应性, 不同的年龄区间对应着不同的地质事件。笔者认为研究区晚三叠世—晚泥盆世沉积物主要来源于兴蒙造山带以及北秦岭造山带印支期



表2 鄂尔多斯盆地西南缘周缘代表性地质体新太古代—古生代年龄统计

Table 2 Statistics of ages for Neoproterozoic to Paleozoic of representative geological bodies in the adjacent areas of southwestern Ordos Basin

事件	地质体	年龄/Ma	测试方法	资料来源
新太古代末—古元古代初期	华北板块西红山乡获得石英闪长岩	2575~2676	LA-ICP-MS	陶继雄等,2002
	华北板块固阳英云闪长岩	2440±35	LA-ICP-MS	张维杰等,2000
	北祁连东段陇山杂岩长英质片麻岩	1900、2350、2500	LA-ICP-MS	何艳红等,2005
	阿拉善地块叠布斯格岩群	1926	LA-ICP-MS	耿元生等,2010
	阿拉善地块哈拉陶勒盖角闪斜长片麻岩	2700	SHRIMP	耿元生等,2006
古元古代中—晚期	阿拉善地块北大山岩浆岩	2496±11	LA-ICP-MS	宫江华等,2012
	华北板块西北缘花岗岩	1958±34	LA-ICP-MS	李正辉等,2013
	华北板块西北缘千里山孔兹岩系	2000~2300	LA-ICP-MS	Yin et al., 2009
	阿拉善地块巴彦乌拉片麻状花岗岩闪长岩	2080	LA-ICP-MS	李俊建,2004
	阿拉善地块东缘	1900~1950	LA-ICP-MS	耿元生等,2010
中—新元古代	北祁连造山带白家沟花岗岩	1846±32	LA-ICP-MS	尤佳等,2014
	华北板块西南缘陇县钾长花岗岩	2018±16	LA-ICP-MS	徐欢,2019
	秦—祁造山带结合部位木其滩岩组斜长角闪岩	762.5±4.6	LA-ICP-MS	张志国等,2011
	祁连山东段兴隆山群火山岩	723~824	LA-ICP-MS	徐学义等,2008a
	西秦岭新阳花岗质片麻岩	935.5±3.1	LA-ICP-MS	刘会彬等,2006
加里东期	秦—祁造山带结合部位陇山岩群	929	LA-ICP-MS	徐可心等,2018
	北秦岭罗汉寺岩群辉长岩脉	475±4	LA-ICP-MS	刘军锋等,2007
	秦、祁造山带结合部位川草铺花岗岩	434±10	LA-ICP-MS	Zhang et al., 2006
	秦、祁造山带结合部位阎家店闪长岩	440.2±0.92	LA-ICP-MS	裴先治等,2007
	秦、祁造山带结合部位阎家店闪长岩	441±10	LA-ICP-MS	Zhang et al., 2006
	祁连山东段陇山岩群辉绿岩墙	440.9	LA-ICP-MS	徐学义等,2008b
	祁连东段清凹山石英闪长岩	402~445	K-Ar	王金荣等,2006
	北祁连造山带阴沟群粗砂岩	425±2	LA-ICP-MS	李猛等,2015
	北秦岭西段与北祁连东段花岗岩	391~450	LA-ICP-MS	徐学义等,2008a
	北秦岭早古生代岩浆岩	420~450	LA-ICP-MS	张成立等,2013
海西期	秦—祁结合部糜楞岩化闪长岩	418±3.2	LA-ICP-MS	孟祥舒等,2017
	兴蒙造山带平泉地区砂岩	250~330	LA-ICP-MS	马收先等,2011
	兴蒙造山带碎屑岩、变质岩、变岩浆岩	250~350	LA-ICP-MS	陈岳龙等,2012
	兴蒙造山带狼山岩体	267~302	K-Ar	陈登超等,2010
	兴蒙造山带东升庙二长花岗岩体	259.4±3.3	LA-ICP-MS	吴亚飞等,2013
	阿拉善地块东北缘霍各乞辉长—闪长岩	273.9±1.2	LA-ICP-MS	皮桥辉等,2010
	西秦岭江里沟花岗岩	264.0±1.4	LA-ICP-MS	孙小攀等,2013
	东阿拉善波罗斯坦庙杂岩	242~284	LA-ICP-MS	邹雷等,2019
	北秦岭花岗岩	256±4	LA-ICP-MS	李侃等,2015
	印支期	秦岭造山带黑沟峡火山岩	221、242	Rb-Sr、Sm-Nd
西秦岭糜署岭花岗岩		237	LA-ICP-MS	李永军等,2004
草关地区花岗岩		205	LA-ICP-MS	李永军等,2004
夏河地区岩体		238	LA-ICP-MS	金惟浚等,2005
西秦岭厂坝花岗岩闪长岩		209~215	LA-ICP-MS	魏然等,2017
西秦岭江里沟复式岩体		229.1±1.8	LA-ICP-MS	路东宇等,2017

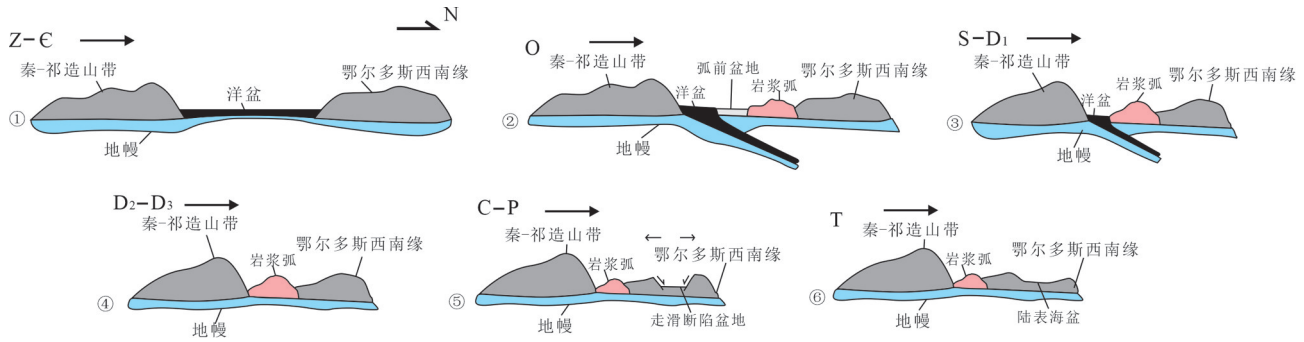


图7 盆山演化过程模式简图(据陈世悦,2000;王银川,2013;罗顺社等,2017)

Fig.7 The model of basin mountain evolution(after Chen Shiyue., 2000; Wang Yinchuan., 2013; Luo Shunshu et al., 2017)

俯冲碰撞的岩浆活动;中奥陶世—晚泥盆世的沉积物主要受秦—祁造山带加里东期的俯冲造山运动控制;中—新元古代沉积物可能来自秦祁造山带和贺兰山地区;古元古代晚期的物源来自阿拉善地块、北祁连造山带和华北板块变质褶皱基底;新太古代的碎屑物源则来自华北板块老基底、祁连造山带以及阿拉善地块古老基底。

结合本次研究认为,研究区碎屑锆石在古生代—中生代年龄区间为 166~472 Ma,碎屑锆石颗粒一共 63 颗,占总数的 63%,反映了周临地区在古生代存在较活跃的构造岩浆活动并为研究区提供主要的物质来源。据记载,震旦纪—寒武纪时期鄂尔多斯盆地西南部与秦—祁造山带间发育洋盆,奥陶纪秦—祁造山带由南西向北东俯冲碰撞(王银川,2013),岩浆弧及弧前盆地随之形成,早古生代晚期—晚古生代早期,洋盆逐渐消亡,弧前盆地随之消亡,并开始碰撞造山(陈世悦,2000)。秦—祁造山带与华北板块西南缘在晚古生代中期发生碰撞,此时秦—祁地区洋陆转化逐渐完成,至二叠纪,秦—祁造山带全面接触碰撞,其北侧鄂尔多斯西南部地区由于走滑拉伸作用,演变为山间断陷盆地,标志着一次盆山转换的完成(罗顺社等,2017;陈世悦,2000)(图7)。鄂尔多斯盆地受古亚洲洋闭合的影响,由石炭纪海陆交互环境转变为陆表海环境,盆地西南缘受到秦—祁造山带印支期的碰撞影响,共同进入了陆内稳定环境(任军锋,2004)。

## 6 结论

(1)鄂尔多斯盆地西南缘镇原地区早白垩世洛河组下段含铀砂岩的碎屑锆石年龄分为 166~370

Ma、388~472 Ma 两个主峰年龄区间和 1744~2150 Ma、2241~2740 Ma 次峰年龄区间以及数据较少且较为离散的中—新元古代年龄 615~1623 Ma。

(2)通过与周缘造山带构造岩浆事件及地层锆石对比研究发现,研究区洛河组下段含铀砂岩的物源较为复杂,分别来自北秦岭造山带、北祁连造山带东段、阿拉善地块、兴蒙造山带、贺兰山孔兹岩系和华北板块基底 6 个物源区,其中北祁连造山带东段、北秦岭造山带为主要物源区。

(3)本次研究发现洛河组下段含铀砂岩古生代年龄数据较多且较集中,标志着奥陶纪秦—祁造山带由西南向北东俯冲;志留纪—早泥盆世洋盆逐渐消亡,中—晚泥盆世秦—祁造山带与鄂尔多斯盆地西南缘共同形成陆—陆碰撞造山带;石炭纪—二叠纪造山带逐渐演变为沉积盆地;经印支期秦—祁造山事件后,鄂尔多斯盆地西南缘沉积环境逐渐稳定。

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